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Scientific Cruise Report of the Arctic Expedition ARK-XX/3 of RV "Polarstern" in 2004: Fram Strait, Yermak Plateau and East Greenland Continental Margin

Wissenschaftlicher Fahrtbericht über die Arktis-Expedition ARK-XX/3 von 2004 mit FS "Polarstern": Framstraße, Yermak-Plateau und ostgrönländischer Kontinentalrand

Edited by Ruediger Stein with contributions of the participants

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# 1. Objectives and Summary R. Stein

The scientific program of the "Polar-stern" ARK-XX/3 expedition, which started in Tromsø on August 31 and ended in Bremerhaven on October 03, 2004, was concentrating on geophysical and geological aspects in the Yermak Plateau, Fram Strait and East Greenland continental margin areas (Fig. 1.1). Major part of these geoscientific studies has to be seen in the context of a drilling proposal submitted to the Integrated Ocean Drilling Pro-

gram (IODP) (Jokat, Stein et al., 2004). In addition, a pilot study devoted to the lost "Deutsche Arktische Expedition 1912" was included in the expedition program. As undone work taken-over from the preceding ARK-XX/2 oceanography program, three PIES (Pressure Inverted Echo Sounder) were deployed west of Spitsbergen.

The seismic network concentrated on the Yermak Plateau area and additional site survey boxes along the East Greenland margin. In total, 2920 km of new seismic data were acquired. These data will allow determining in the optimum way the locations of IODP drill sites for our IODP proposal. Parallel to the seismic profiling across the Yermak Plateau an intensive aeromagnetic survey was performed by helicopter, and more than 7250 km were flown.

Overall goal of the geophysics program was to enlarge the geophysical database in the Fram Strait/Yermak Plateau area to provide further constraints on the deeper structure and the tectonic evolution of this region. In detail the following problems should be addressed:

- Seismic investigations across the northern Yermak Plateau to identify the continent-ocean transition.
- Deep seismic investigations to identify the deeper structure of the Yermak Plateau. With the wide-angle reflection data it should be possible to make estimates on the magmatic underplating, if any is present, of the Yermak Plateau.
- Try to sample basement rocks in order to date and constrain the evolution of the plateau.

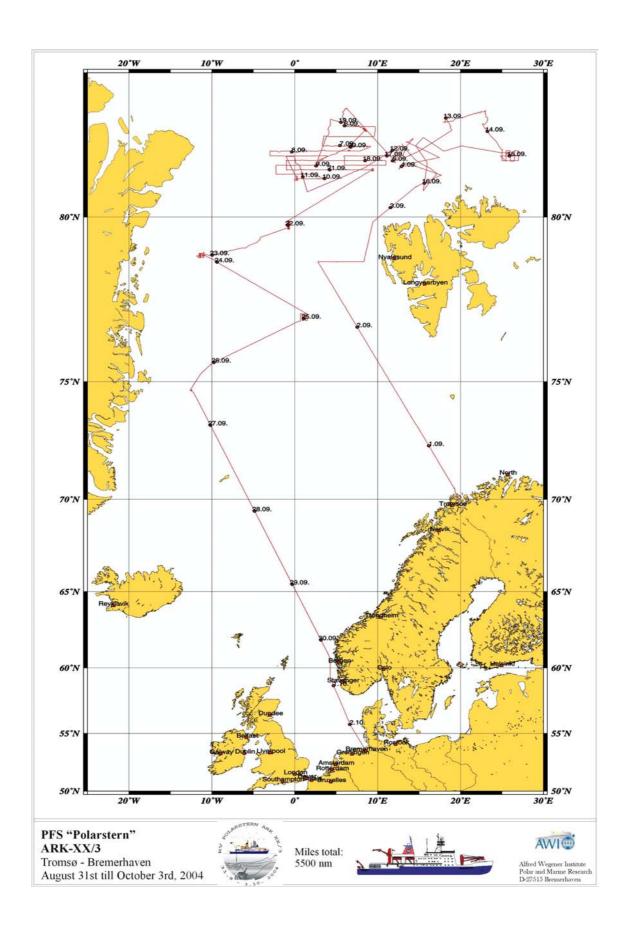


Fig. 1.1: Cruise track of RV "Polarstern" during ARK-XX/3.

One of the key marine-geology objectives was the study of sediment dynamics of megaslides along the Svalbard continental margin (as part of the ESF Program "EUROMARGINS"), using Hydrosweep (bathymetry) and PARASOUND profiling data, multi-sensor-core-logging data, and long sediment cores. In this context, the characterization of sediment facies within debris flow and turbidite sequences as well as undisturbed pelagic sequences, the dating of sediment mass flows and estimates of sedimentary budgets are of interest. Furthermore, detailed multidisciplinary (i.e., sedimentological, mineralogical, micropaleontological, and geochemical) studies will be performed on the sediment cores for high-resolution reconstructions of paleoclimate, paleooceanic circulation patterns, paleosea-ice cover, and paleoproductivity and their variability during late Quaternary times. In order to reach these goals, an intensive sampling program was performed using large-volume box corer, multicorer, gravity corer, and kastenlot corer. In addition, gravity cores were taken from the central part of the Yermak Plateau as well as the East Greenland continental margin as part of the IODP Site survey work. During the entire period of working within the study area, a Hydrosweep and PARASOUND survey was performed. In total, 7800 km of new Hydrosweep and PARASOUND profiles were obtained.

Along the ice-edge area in the Yermak-Plateau area, a sea-ice sampling program was carried-out. In total, 9 stations were sampled for snow, sea ice, and sea-ice sediments.

A pilot study was devoted to the lost Deutsche Arktische Expedition (DAE) 1912, also known as Schroeder Stranz Expedition, a trial expedition to the North East Land of Spitsbergen thought to be a forerunner of the ambitious plan to cross the North East Passage. During "Polarstern" cruise ARK-XX/3 two reported places of historical interest were visited by helicopter to locate and to document: (a) the landing place of the DAE in the Duvefjord and (b) the landing place of the Lerner relief expedition 1913 in search for Schroeder Stranz. The expedition vessel of Lerner got entrapped in the pack ice and sunk off the Beverlysund near the North Cape. With the help of historical stereo-photographs and ARK-XX/3 field observations, the wreck site could be identified more accurately. The wreck is the most northern wreck of the world, and it is planned to investigate it for various research purposes during a submersible expedition in 2006.

In summary, a huge amount of new geoscientific data could be selected. These data will be an important basis for the more concrete planning of our IODP drilling proposal.

#### Acknowledgement

The Scientific Party of "Polarstern" Cruise ARK-XX/3 gratefully thank captain Udo Domke and his crew, as the success of our expedition was substantially supported by their excellent cooperation and efforts.

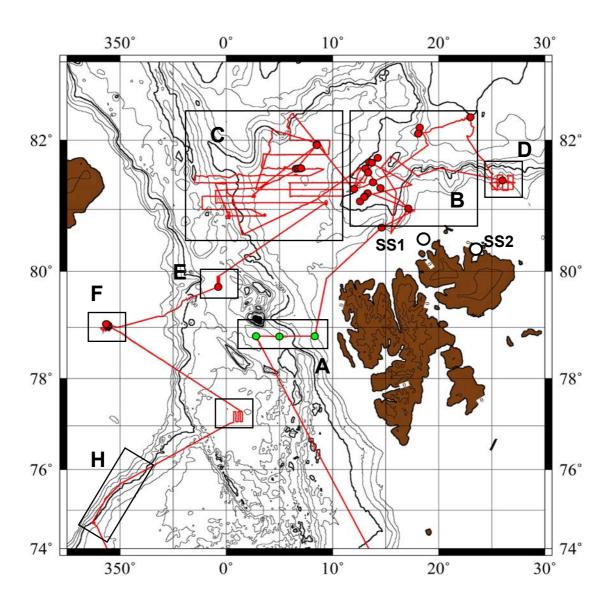
# 2. Itinerary

R. Stein

RV "Polarstern" left the port of Tromsø/Norway on August 31, 2004.

After a 52h transit we reached the area of PIES deployment west of Spitsbergen along 78°50′N (Fig. 2.1, Area A) on September 02, 22.00 UTC. As undone work taken-over from the preceding ARK-XX/2 oceanography program, three PIES (Pressure Inverted Echo Sounder) were deployed within a 6h work program:

PIES 71	78°50.25′N, 02°48.20′E	2496 m water depth
<b>PIES 141</b>	78°49.87′N, 05°00.93′E	2708 m water depth
PIES 62	78°50.03′N, 08°19.91′E	793 m water depth



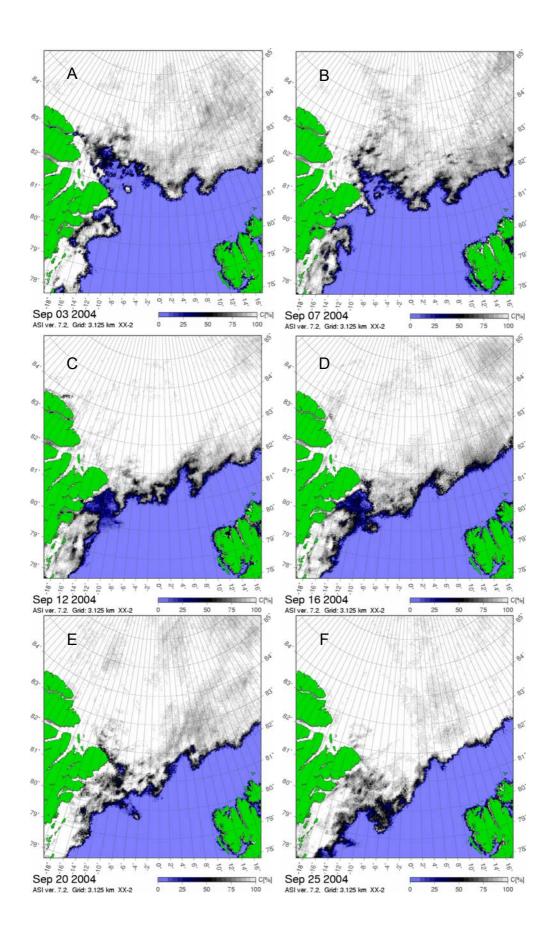
**Fig. 2.1**: Working areas (A to H) of Cruise ARK-XX/3. Cruise track and stations are shown. In addition, the two locations on North East Land of Spitsbergen, devoted to the Schroeder Stranz Expedition 1912, are marked (SS1 and SS2).

On September 03, 15.00h UTC, we arrived at the first geological station (PS66/304; 80°43.36′N, 14°39.39′E) and started with our own research program in the area north of Spitsbergen (Fig. 2.1, Area B) under optimum ice-free conditions. The ice edge was retreated as far north as 82°N (Fig. 2.2). An intense sampling of near-surface sediments and sediment cores was carried out between September 03 and 05, using giant box corer (GKG), multicorer (MUC), gravity corer (SL), and kastenlot corer (KAL). On September 03 a first helicopter flight, devoted to the Schroeder-Strantz program, was performed to Cape Rubin (Fig. 2.1, area around location SS1).

The seismic profiling program on Yermak Plateau started in the morning of September 05 (Fig.2.1, eastern part of Area C). On September 06, 05.20 UTC, we had to stop seismic profiling near 82°25′N, 06°E (our northernmost location of this expedition) due to too strong ice conditions (Fig. 2.2). We steamed towards 82°N, 06°E where seismic profiling was continued until the afternoon of September 12. In the following two days, the eastern flank of the Yermak Plateau (Fig. 2.1, Area B) was sampled by means of GKG, MUC, SL, and KAL coring systems. Due to strong pack ice, we were proceeding slowly to our northeast-ernmost station (PS66/323; 82°03.52′N, 22°38.30′E). After having finished the station work at that site, we were steaming towards the SE.

In the early afternoon of September 14, we reached the continental slope north of North East Land at about 81°30′N, 25°E (Fig. 2.1, Area D) where a detailed seismic site survey for a proposed IODP drilling location was carried out. Parallel to the seismic profiling, a helicopter flight was performed to the Duvefjord area where remains of the Schroeder-Stranz Expedition were found in 1937 (Fig. 2.1, area around SS2). We finished the work in Area D with a geological sampling at site PS66/325 (81°26.18′N, 25°59.86′E) in the afternoon of September 15. Until the early morning of September 17, we completed our geological sampling program in Area B at further four stations.

We continued our work program with a seismic profiling on Yermak Plateau (Area C) during the following two days. On top of the central Yermak Plateau where it is assumed that older sediments are cropping out, a geological sampling was carried out by means of GKG, SL and the sediment dredge on September 20 (Fig. 2.1, Area C). We finished our working program on Yermak Plateau Area C with a seismic profiling lasting until September 21, 20.00 UTC.



**Fig. 2.2**: Sea-ice condition during ARK-XX/3 <a href="http://iup.physik.uni-bremen.de:8084/amsr/amsre.html">http://iup.physik.uni-bremen.de:8084/amsr/amsre.html</a>

Parallel to the seismic profiling and geological sampling across the Yermak Plateau and the Svalbard continental margin, an aeromagnetic survey was performed by helicopter. Within nine days and 20 flights most of the survey area was successfully completed and more than 7250 km were flown.

Furthermore, a sea-ice sampling program was carried out along the ice-edge area in the Yermak-Plateau area (see Fig. 2.2) between September 07 and 19, using the helicopter. In total, 9 stations were sampled for snow, sea ice, and sea-ice sediments. These samples will be investigated for mineralogy, grain size, organic carbon content and composition, and/or major and minor element composition.

On September 21, 22.00 UTC, we steamed towards the south and reached Area E on September 22 (Fig. 2.1) where a seismic site survey for a proposed IODP drilling location was carried-out. In addition, near-surface sediments were sampled by means of gravity corer (Station PS66/341; 79°44.07′N, 00°45′W).

In the early morning of September 23, we arrived on the East Greenland continental shelf (Fig. 2.1, Area F). Several big icebergs were observed here, some of them contained large amount of sediment (see Chapter 4). In Area F, seismic profiling work was done. In this area, seismic data reveal the presence of a prominent salt province and suggest that Mesozoic strata are cropping-out at the surface (Schmitz and Jokat, 2005). In order to sample these old strata, 10 short gravity cores were taken. In one of the gravity cores (PS66/345-2), a large crystal of gypsum was found in the core catcher, which may be related to Permian salt deposits (see Chapter 9.6.4).

After having finished this station work, we steamed towards the SW. During the last two days of our working program (September 25 and 26), seismic site survey profiling was continued in the central southern Fram Strait (Boreas Basin) (Fig. 2.1, Area G) and along the East Greenland continental slope between about 76°N and 74°30′N (Fig. 2.1., Area H).

On September 27, 05.30h UTC, we started our way back to Bremerhaven. "Polarstern" arrived in Bremerhaven on October 03 early in the morning, bringing to an end a scientifically very successful cruise.

## 3. Meteorological Conditions

M. Gebauer, K. Buldt

During departure on 31 August 2004 the weather was fine. An anticyclone was centered near the Greenland Sea, only a shallow low moved from Svalbard in direction of Novaja Semlja. The following days the high pressure influence presented moderate easterly winds until we reached Fram Strait. During arrival in this area the wind changed to a southerly direction and its force increased a little bit. Due to warm air advection in front of a cyclone that had developed near Iceland, there was occasional drizzle und poor visibility.

On September 03, "Polarstern" passed the eastern coast of Svalbard and arrived at the exploration area near Yermak Plateau during the next day. The wind speed increased up to 6 Bft (Fig.3.1), but there were some fog patches due to southerly winds that brought moist and rather mild air over the undercooled water near ice fields. Flight weather was problematic with some restrictions of visibility and ceiling conditions (Figs. 3.2 and 3.3). In the meantime the Iceland low moved across the Greenland Sea to the Svalbard area (Fig. 3.4).

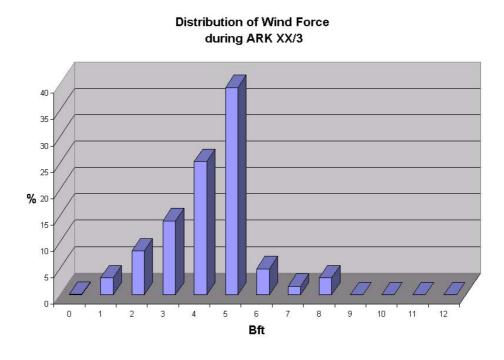


Fig. 3.1: Distribution of wind speed according to normal climatic conditions in September

#### Ceiling during ARK XX/3

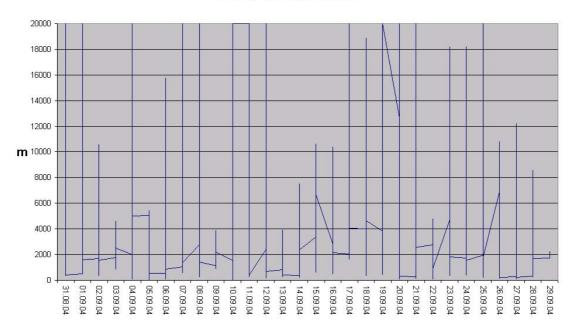


Fig. 3.2: Ceiling: Measurements per 10 minutes; cloud conditions changed very fast.

#### Visibility during ARK XX/3

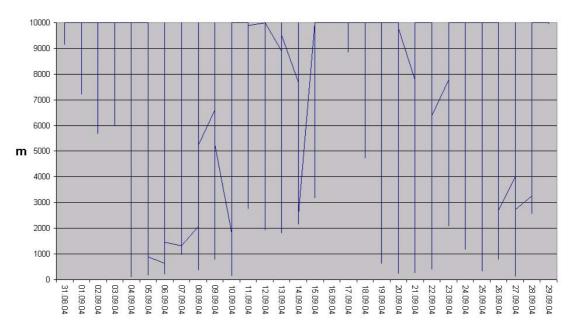


Fig. 3.3: Visibility: Measurements per 10 minutes; conditions changed also very fast.

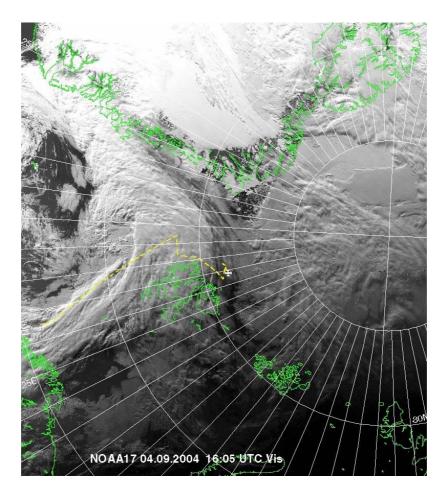


Fig. 3.4: A new low arriving; dashed line indicates course of "Polarstern", cross her position.

The following weather period was determined by moist and rather mild air masses. The wind came from easterly directions and then, when the low had moved off, from northerly directions. It reached for a short time a wind force 8 to 9 Bft, accompanied by intermittent rain or drizzle with a temperature only barely exceeding the freezing point and thus inducing danger of icing - helicopter flights were not possible

From September 07 onwards, a new weather situation with cold air advection began. Snow was falling, and the air temperatures smashed below –5 °C, later –10 °C. Friendly periods with good visibility and convenient flight weather changed temporarily with some disturbances and shallow lows that arrived from northwest. Intermediately there fell snow.

When the temperature finally rose, the weather conditions for helicopter flights were difficult with a change of snow, rain and also freezing drizzle. The wind blew from northeast with 4 to 5 Bft (Fig. 3.1).

During the middle of the month the high pressure influence rose from Greenland. It was very cold, and due to the low condensation level and very low clouds, the ceiling of less than 300 ft was sometimes not sufficient for flights with the magnetic sensor more than 30 m below the helicopter.

The last part of the month was determined by cyclones, which moved slowly from Iceland to Scandinavia. The working area of "Polarstern" was between these cyclones and high pressure influence over Greenland and over the islands Svalbard, Franz Joseph Land and Novaja Semlja. Only shallow lows were able to influence us temporarily with low clouds, wet snow and poor visibility. Mostly it was cold with -3 to -9 °C. The wind blew from different directions and varied between 2 and 5 Bft.

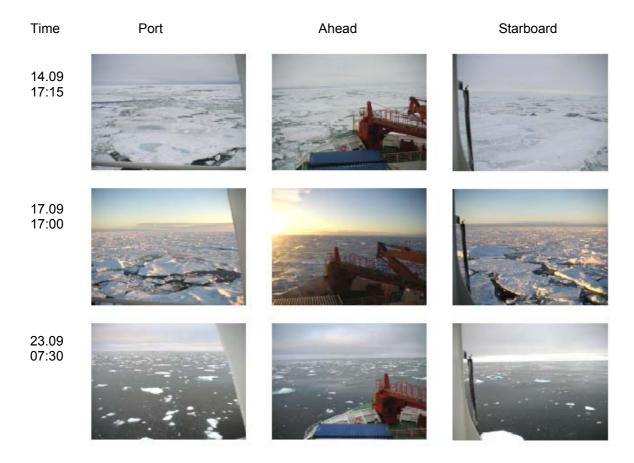
In the last week of September a new cyclone, once it had been the hurricane "Karl", moved from Iceland to the middle of Norway. North of it "Polarstern" moved southeast, accompanied by mostly northerly, later on westerly winds blowing with 5 Bft. After passing the post cold-frontal weather, a short high pressure influence followed, then a short period during that "Polarstern" had to work against southerly winds 6 Bft and its waves.

"Polarstern" came back to Bremerhaven on October 03, 2004.

#### 4. Sea-ice Observations

V. Shevchenko

"Polarstern" met sea ice on September 5 evening at 81°49.8'N, 9°12.4'E. During seismic profiling, "Polarstern" came to the area with strong ice conditions in the morning of September 6 at 82°20'N, 6°E. After that she came back to mostly open water at 82°N, 6°E and continued seismic studies. In the morning of September 7, regular visual ice observations from the ship's bridge began. They continued till September 24. These observations were not performed during night time and during the work in ice-free waters. Variables like ice concentration, ice thickness, floe size, melt-pond coverage, ridge height and age, as well as the occurrence of ice-rafted sediments ("dirty ice") and icebergs were recorded, representing ice conditions in an area of 500 to 1000 m around the ship similarly as it was done during the ARK XVII/2 expedition of "Polarstern" in August—September 2001 (Haas and Lieser, 2003). For most typical situations photos were made from the bridge (Fig. 4.1). In total 89 observations were carried out. Results of these observations are presented in the Annex (Chapter 12.3).



**Fig. 4.1:** Photos from the bridge of "Polarstern" (performed by R. Stein)

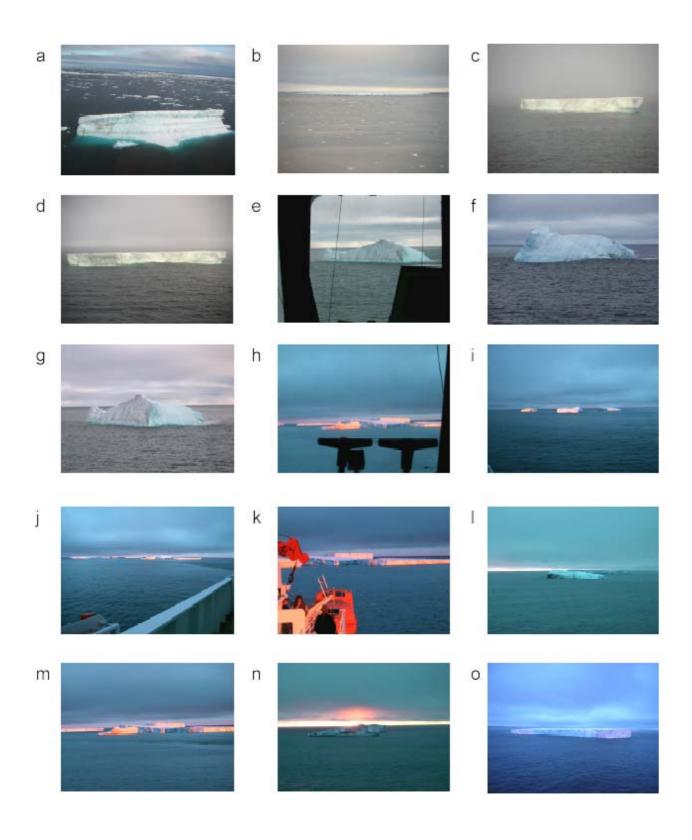
In the area of Yermak Plateau and Fram Strait we mostly worked in the ice-free zone or in the marginal ice zone (region within a hundred kilometres or so of the ice edge). In this region the ice tended to resemble a granular medium with relatively small floes, often densely packed (Hibler, 1989). The typical sizes of ice floes varied from 5 to 20 m, sometimes they were up to 50 m. The typical thickness of ice floes varied from 0.5 to 1 m, snow depth varied from 0 to 15 cm.

Melt ponds were not clearly expressed on the ice floes, probably due to recent snow cover. As it was shown earlier, with the beginning of snowfalls in September the identification of melt ponds by aerial photography and visual observations became difficult or impossible (Bareiss and Haas, 2002).

Ice-rafted sediments were very rare (for more details see Chapter 5). Even their search during the helicopter flights sometimes was not successful. It is possible, that part of these dirty patches on sea ice was covered by recent snow.

Often (see Annex, Chapter 12.3) brown ice was observed on the bottom of ice floes when "Polarstern" reversed them. This discoloration of ice is the result of active development of ice algae (ice algal bloom), that is typical for marginal ice zones (Smith, 1987).

During the expedition icebergs were observed. Photos of some icebergs are presented in Figure 4.2. Along the ship route in the Yermak Plateau area we observed few small icebergs (Fig. 4.2a). In the western part of the Fram Strait (NW part of the Greenland Sea) icebergs were more abundant. In the vicinity of point 79°42' N, 0°48' W three icebergs having sizes from 100 to 300 m, were registered. In the morning of September 23, we crossed a narrow field of old destroyed ice floes (79°12.10' N, 6°11.69' W) and observed three icebergs (largest of them had a size of 500 m to 600 m and 25 m height; Fig. 4.2b). Moving to the west we arrived in the area of detailed seismic studies on the East Greenland shelf, and at 79°02.67'N, 10°55.6'W we were close to an iceberg of 100 m length. This iceberg was reversed to its side, and layering was clearly seen. Dark layers are enriched by sediments (Fig. 4.2e–g). In the vicinity of point 79°03' N, 10°56' W, we observed 16 icebergs ranging in size from 100 to 2000 m and in height from 15 to 30 m (Fig. 4.2h–o).



**Fig. 4.2:** Photos of icebergs (from the archive of the expedition): a — September 13, 17.00, 82°05.3'N, 18°03.0'E; b — September 23, 07:45, 79°12.1'N, 6°27.0'W; c–d — September 23, 12:00, 79°0.32'N, 10°12.2'W; e–g — September 23, 13:00, 79°02.67'N, 10°55.6'W; h–o — group of icebergs on September 23, 18:30 around location 78°55.6'N, 11°17.5'W.

## 5. Sea-ice Sedimentology and Snow Chemistry

V. Shevchenko, N. Kukina

#### 5.1. Ice-rafted Sediments

V.Shevchenko, N. Kukina

Sediment transport via sea ice is expected to contribute significantly to deep-sea sedimentation at least in regions of ice ablation (Pfirman et al., 1990; Nürnberg et al., 1994; Lisitzin, 2002) with potential importance also for the dispersal of pollutants (Pfirman et al., 1995; Meese et al., 1997). In the Arctic Ocean, sediments are transported by ice drifting mainly in two major surface current systems – the anticyclonic Beaufort Gyre in the Amerasian Basin and the Transpolar Drift in the Eurasian Basin.

As shown in many works, large amounts of sediments can be entrained into the ice cover through several processes: scavenging of suspended material from the water column by frazil-ice crystals, uplift of material by anchor ice, river spilling on sea ice and aeolian transport (e.g., Mulen et al., 1972; Reimnitz et al., 1987; Dethleth et al., 2000; Eicken, 2003). Now there are many evidences that much of its particulate load originates from the Siberian and specifically the Laptev Sea shelf (Pfirman et al., 1990; Wolenburg, 1993; Nürnberg et al., 1994; Eicken et al., 1997; Lindemann, 1998; Hölemann et al., 1999; Dethleff et al., 2000; Eicken, 2003). During its drift multi-year sea ice is exposed to the extensive surface melting which takes place each summer. Surface melting affects the distribution of particles in sediment-laden multi-year sea ice. Seasonal surface ablation of snow and sea ice and freezing of new ice at the base will result in surface accumulation of particles from the entire melted snow and sea-ice column, while "clean" ice is added underneath each winter that ice-floes are in the central Arctic (Pfirman, 1990).

Surface ablation of sediment-laden ice may form cryoconites (e.g. Pfirman et al., 1990). Cryoconite formation results in aggregation of particles at the base of cylindrical holes in the ice surface. With extensive freeze-thaw cycling, formation of more or less cohesive pellets may occur (Barnes et al., 1988). This process may shield particle accumulation from melt water wash-off and run-off and thus delay particle release until the bulk of the floe disintegrates. When released from the floe, cohesive cryoconite pellets may settle intact to the sea floor. Similar pellets found in sediment traps in the Fram Strait have been attributed to this process (Berner and Wefer, 1990).

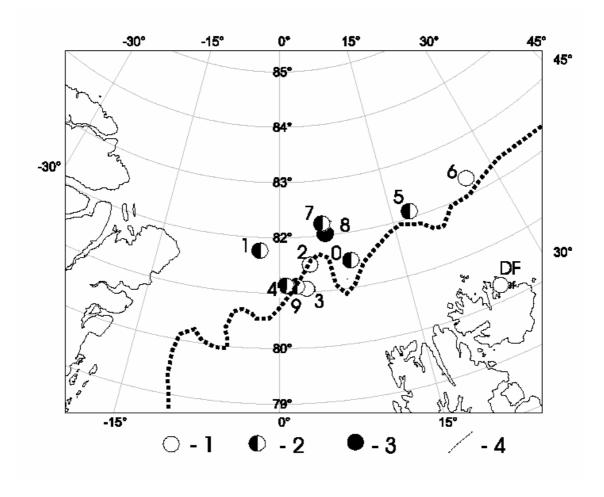
Because incorporated material is concentrated mostly at the ice surface in multi-year floes, it appears likely that deposition on the sea floor will occur primarily along the drift path of the ice-floes as it finally disintegrates. Thus, the greatest potential for sediment deposition is expected to occur along the ice margin, near polynyas and large leads. Much of the ice-rafted sediments transported by the Transpolar Drift may be deposited along the axis of the East Greenland Current (Pfirman et al., 1989), with some inflow into the Barents Sea (Elverhøi et al., 1989; Dethleff, 1997).

Determination of the source is very difficult question. The main purpose of studies of ice-rafted sediments is to reveal the main sources of sediments in the area of Yermak Plateau and Fram Strait.

Table 5.1: Time and location of sampling of snow and ice-rafted sediments (IRS) in the Yermak Plateau area.

Date	Investigators	Ice	Coordinates		No. of sam	ples
		station	Lat. (N)	Longitude	Snow	IRS
07.09.2004	Shevchenko V.	0	81°29.96'	08°43.61 'E	1	1**
08.09.2004	Shevchenko V. Schäfer C.	1	81°46.0'	02°27.0' W	2*	
09.09.2004	Shevchenko V. Winkelmann D. Kukina N.	2	81°30.20'	03°39.12' E	3*	
10.09.2004	Shevchenko V. Schoster F. Mende I.	3	81°03.73'	03°09.22' E	4*	2
11.09.2004	Shevchenko V. Schneider J. Penshorn D.	4	81°07.25'	00°54.96' E	5	3**
13.09.2004	Shevchenko V. Yanina Ya. Saarswat R.	5	82°08.22'	17°20.31' E	6	
14.09.2004	Shevchenko V. Bahr B. Noffke H.	6	82°18.21'	25°56.27' E	7	4**
14.09.2004	Fricke N. Mende I.	DF	80°16.8'	24°24.4' E	8	
19.09.2004	Shevchenko V.	7	82°12.23'	05°41.76' E	9	5**
	Nam S. Puhr A.	8	82°02.41'	05°55.67' E		6**
21.09.2004	Shevchenko V. Winkler A. Mende I.	9	81°05.70'	02°00.60' E	10	7**

<sup>\*</sup>Snow sampling for oxygen isotopy was also collected \*\*Subsamples of IRS for biomarker analysis were taken



**Fig. 5.1:** Position of sampling sites: 1 – only snow; 2 – snow and ice-rafted sediments; 3 – only ice-rafted sediments; 4 – ice edge on September 15, 2004 according to satellite data (from www-site of Bremen University).

During the ARK-XX/3 Expedition, we collected samples of ice-rafted sediments on ice floes for multidisciplinary studies. Date, coordinates and short description of sampling sites are given in Table 5.1. Positions of sampling sites are shown in Figure 5.1. All sampling sites were accessed by helicopter. At six stations we found ice-rafted sediments in the form of cryoconite pellets 1–10 cm long dispersed in the 1–2 cm ice layer on the surface of ice ridges. Only at ice station 4, sediments formed a 2–5 mm thick crust on the surface of smooth ice ridges (1 m in height) (Fig. 5.2). Sampling of ice-rafted sediments was carried out by stainless knife and plastic shovel in plastic bottles for trace metal analysis and in stainless tanks for organic geochemistry (Fig. 5.3). Onboard "Polarstern", ice enriched by sediment pellets was melted at +4°C, and after decantation subsamples of sedimentary material for biomarker analysis were stored at –30°C and for other analyses at +4°C.



**Fig. 5.2**: Grey sediment-laden ice ridges at an ice floe in the Yermak Plateau area (Station 4,  $81^{\circ}07.25^{\circ}N$ ,  $0^{\circ}54.96^{\circ}E$ ).



Fig. 5.3: Sampling of sediment crust on the ice ridge (Station 4).

Smear slide analyses carried out on sea-ice sediments from seven stations (Table 5.1, Fig.5.1), were performed to estimate the grain size distribution and mineralogical composition of the material. In general, the sea-ice sediments are composed of fine-grained material with partly more than 80 % in the fraction less 63 µm (Table 5.2). Only from one station (Station 1) located in the western part of the Yermak Plateau, coarse-grained sea-ice sediments were sampled that consist of 18% of the sand fraction. The sandy material points to entrainment of surface deposits through anchor ice formation or bottom adfreezing. At other six ice floes we found sediments that more probably were entrained into the ice through scavenging of suspended material from the water column by frazil-ice crystals.

Table 5.2: Grain size distribution ice-rafted sediments (IRS) in %.

No. of IRS sample	Sand	Silt	Clay	Sediment type
1	18	46	36	sandy clayey silt
2	1.2	34.1	64.7	silty clay
3	2.1	41.9	56	silty clay
4	8.6	62	29.4	clayey silt
5	3.6	64	32.4	clayey silt
6	3.9	55	41.1	clayey silt
7	1.1	30.5	68.4	silty clay

Based on smear slide analyses, terrigenous particles dominate in the sea-ice sediments during ARK-XX/3. The principal minerals include quartz, feldspar, mica and clay minerals (Table 5.3). Quartz contents ranged from 16.9% (Station 2) to 36.6% (Station 6). Feldspar contents were up to 21%. The highest amounts of clay minerals (>20%) were observed in sea-ice sediments of stations 2, 3, and 7. The sea-ice sediments of western part of the Yermak Plateau (Stations 1, 2, 4, and 7) obtained less than 1.9% of volcanic glasses. The content of organic remains generally represented by diatoms, spores and plant debris, was between 3.7% and 15.9%.

Table 5.3 : Results of smear slide analysis: Mineral composition of ice-rafted sediments in %

Number*	Quartz	Feldspars	Mica	Volc. glass
1	31	13,2	1,2	1,9
2	16,9	10,2	2,3	1,8
3	30,2	20,1	9,5	0
4	30,2	20,6	7,2	1,2
5	30,9	20,4	5,2	0
6	36,6	20,3	9,2	0
7	30,2	14,2	7,1	1,9
Number*	Clay minerals	Org. remains	Fe-hydroxide	Pyroxene
1	18	3,7	9,4	7,6
2	29,1	15,9	5,3	4,5
3	20	9,2	0	3,1
4	7,3	6,8	3,5	6,8
5	9,2	9,6	2,1	3,5
6	7,2	15,9	1,5	1,4
7	20,3	14,2	2,4	1,8
Number*	Amphibole	Garnet	Epidote	Black ores
1	2	1,8	4	5
2	2,6	3,5	1,1	3,5
3	1,6	2,1	2,2	1
4	5,2	3,5	2,1	3,1
5	7,9	9,2	0,1	1,3
6	1,3	3,2	0,2	0,9
7	2,2	2,8	0,8	1
Number*	Turmaline	Rutile	Pyrite	Opaque
1	0,1	0	0	1,1
2	0,7	1,1	0,4	1,1
3	0	0	0	1
4	0	1,8	0	0,7
5	0	0	0	0,6
6	0,4	0,6	0,5	0,8
7	0,4	0,2	0,2	0,3
*Number of II	RS sample (see T	able 5.1).		

Heavy minerals include pyroxenes, amphiboles, epidote, garnets, Fehydroxides and black ores. Pyroxenes and amphiboles dominate the spectrum although in sea-ice sediments of some stations (Stations 1 and 2) Fehydroxides dominate the association.

Further laboratory investigation of ice-rafted sediments will include determination of sand/silt/clay percentages (wet sieving and Atterberg separation), silt grain size distribution (Sedigraph), clay mineral assemblages (X-ray diffractometry), total and organic carbon and nitrogen content, biomarkers and elemental composition (instrumental neutron activation, atomic absorption, synchrotron radiation X-ray fluorescence analyses) as well as diatom analysis. Analytical work will be carried out in at the Alfred Wegener Institute for Polar and Marine Research Bremerhaven, the P.P. Shirshov Institute of Oceanology Moscow, the Murmansk Marine Biological Institute, and the United Institute of Geology, Geophysics and Mineralogy SB RAS Novosibirsk, and diatoms will be studied in Moscow State University.

At ice station 5 (82°08.22'N, 17°20.31'E) 1.5 kg of visually clean ice were collected from 2 m high ice ridge as reference sample of Arctic multi-year ice for comparative study of carbon and sulfur isotopes composition with tabular ground ice in the Russian Arctic. The sample is stored in refrigerator at –30°C, and it will be analysed in Institute of Oceanology RAS and P.P. Shirshov Institute of Oceanology RAS, Moscow.

#### 5.2 Snow Chemistry

V. Shevchenko

Aeolian transport of particulate matter onto the Arctic sea ice is one of the sources of sedimentary material in the Arctic (Shevchenko et al., 2000; Shevchenko, 2003). Its role in sedimentation in the Arctic Ocean has been studied insufficiently. The drifting ice in the Arctic is a giant natural accumulator of the aeolian material and attendant pollutants, which first are deposited onto the ice and are transformed into cryosols. When the ice melts, often many thousands kilometres away from the places of their fallout, they are released into the water (mostly in Fram Strait and the Greenland Sea).

The main purpose of our work is to study composition of snow covering drifting ice in the remote arctic area to find sources of dissolved and particulate components. Most snow samples were collected on ice floes. One sample was taken on the shore of the Duvefiord, Spitsbergen. Date, coordinates and short description of sampling sites are given in Table 5.1, position of sampling site is shown in Figure 5.1. All sampling sites were accessed by helicopter at the distance more than 2 km upwind from the RV "Polarstern".

At most sites we collected upper 3-4 cm of snow, reworked by winds. Only on September 10, the upper 3 cm of really fresh snow were taken on an ice floe (Sample No. 4). Snow was collected in pre-cleaned plastic bags using a plastic shovel. Samples are stored in the refrigerator at -30°C till the analysis in the land laboratory.

In the laboratory snow will be melted at +18°C. After that pH will be measured, and ion chromatography will be carried out. Snow melt water will be filtered

through Whatman GF/F and pre-weighted Nuclepore (pore size of 0.45  $\mu$ m) filters to determine concentrations of total particulate matter, particulate organic carbon, black carbon, particulate nitrogen. Qualitative composition of particulate matter will be studied by scanning electron microscopy. Most laboratory work will be carried out at the Alfred Wegener Institute for Polar and Marine Research, Bremerhaven and at the P.P. Shirshov Institute of Oceanology RAS, Moscow. Contents of black carbon will be determined at the Institute of Chemical Kinetics and Combustion SB RAS, Novosibirsk. The obtained data will be used for an estimation of the aeolian input of particulate matter and chemical components to the drifting ice.

At three ice floes snow for oxygen isotope composition analysis was collected (Table 5.1). After melting at +4°C water subsamples were mixed and combined. In total 30 I of water was obtained. It will be analysed at the Technical University of Freiberg, Germany.

#### 6. Secchi Depth Measurements

V. Shevchenko

The Secchi disc was used for qualitative estimates of the concentration of suspended particulate matter (turbidity). The Secchi disc is a flat circular plate, 30 cm in diameter, all white. It is lowered through the water column in a horizontal attitude until it is observed just to disappear. The depth, at which this happens, is called the Secchi depth, and it depends on the turbidity of the water. The Secchi disc is both cheap and easily made, and it has been used by oceanographers for over a century as a rapid means of assessing water clarity (Seawater, 1995).

In our expedition nine measurements of the Secchi depth were carried out (Table 6.1). It varied from 13 to 18 m with an average of 14.7 m. It shows that the concentration of suspended particulate matter (SPM) in water is very low. The SPM concentration in the Yermak Plateau area in September 2004 was estimated using Secchi depth data and fitting power equation calculated from data of simultaneous SPM and Secchi disk measurements in the White Sea in April 2003 (V. Shevchenko and A. Filippov, unpublished data). Accordingly to these qualitative estimates SPM concentration in the upper 5 m of water column in the studied area varied from 0.23 to 0.37 mg/l (0.31 mg/l in average). For comparison, in the Kara Sea the 5-m transparency isoline shows the border of the main area of influence of the fresh turbid waters from Ob, Yenisey and other rivers (Lisitzin et al., 2000). Only in the outer part of the Kara Sea where the concentration of suspended matter is <0.5 mg/l the Secchi depth increases up to 15 m. In the Gakkel Ridge area in August–September 2001, the Secchi depth varied from 16 to 30 m with an average of 21 m (Shevchenko, 2002).

**Table 6.1** Time, locations and results of Secchi depth measurements.

			Coordinates								
Station	Date	Time			Secchi						
		(UTC)	Latitude (N)	Longitude (E)	depth (m)						
PS 66/318	12/09/2004	16:40	81°18.52'	12°2.55'	14						
PS 66/321	13/09/2004	08:45	82°10.38'	18°15.02'	14						
PS 66/322	13/09/2004	14:04	82°5.36'	18°5.27'	13						
PS 66/325	15/09/2004	15:35	81°26.19'	25°59.85'	14						
PS 66/326	16/09/2004	14:00	80°43.23'	15°52.85'	16						
PS 66/327	16/09/2004	16:32	81°0.99'	17°9.54'	15						
PS 66/333	19/09/2004	17:29	81°55.19'	8°35.81'	18						
PS 66/336	20/09/2004	08:09	81°36.35'	6°47.74'	14						
PS 66/337	20/09/2004	09:30	81°36.33'	7°1.32'	14						

# 7. Bathymetric Investigations

B. Platten, R. Rathlau, A. Winkler

The main task of the bathymetric working group on this cruise was the support of the geological and geophysical working groups by creating bathymetric charts of their research areas by means of the Hydrosweep system developed by STN Atlas.

The deep-water sounding system Hydrosweep D2 (Fig. 7.1) on RV "Polarstern" DS2 stands out through a 90°/120° coverage angle in which the seafloor is recorded with 59 specific values for water depths (hard beams) perpendicular to the ship's long axis. The sonar beam has an accuracy of approximately 1% deviation from the real waterdepth after the correction of an automated crossfan calibration. The mean sound velocity over the vertical water column is determined by regular transmission and measurement of a sweep profile in the ship's longitudinal direction and comparison of the slant beams with the vertical beam. These data are used for the depth computation

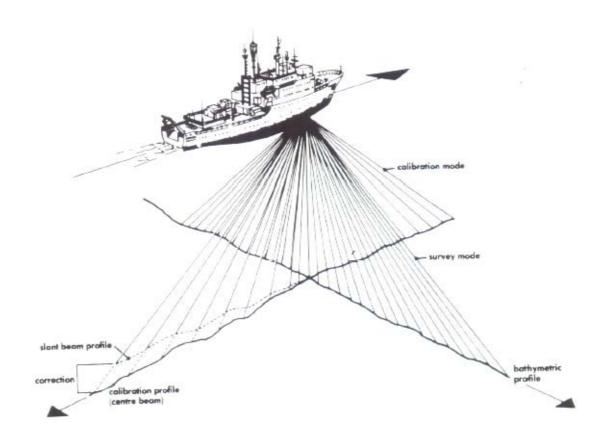


Fig. 7.1: Multibeam sonar system

The operation of the sonar system and the data processing made up the main part of the work. 12 days of the 27 days lasting survey were done under thick ice conditions. In order to prevent abnormal functioning and false measurements due to the thick ice conditions the mostly automatically working multibeam sonar system had to be observed continuously. The changing of the ship's speed and direction due to ice breaking as well as hydroacoustic disturbances caused faulty measurements. Even post-processing became difficult because of the ice conditions, so it was mandatory to check the positions, depths and the ship's attitude data continuously for outlier- or blunder-values. The corrections of the data as well as the general post-processing tasks were done with the software CARIS HIPS and SIPS v. 5.3.

Despite these difficulties the bathymetry working group managed to process all data so far that they now can be used for further examination directly at the AWI.

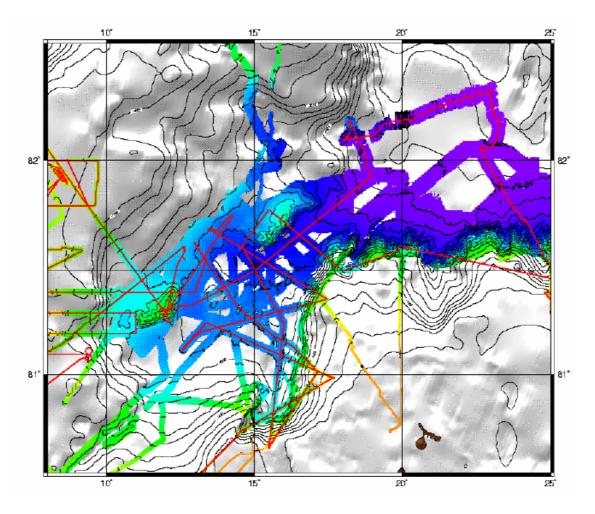
Beside the sonar operation and the data editing the analysis of multibeam data, the preparation of quick-look track plots and the creation of preliminary bathymetric charts were made. Further presentations of the sea bottom topography grids were calculated out of the edited data with the Generic Mapping Tool (GMT) software.

Based on the grids, contour line maps with a contour interval of 100 m and scales up to 1:150'000 were produced with colour coded depth ranges. These

maps allowed a vivid display of the oceanic ridge topography. The IBCAO Vol. 1 (grid spacing 2500 m) data set was used as background bathymetric data in the contour maps. But since the IBCAO in that region is based on sporadic single beam sonar tracks and bathymetry estimated from satellite data, there were substantial differences between the systematically surveyed maps and the IBCAO data. By the combination of previously recorded data with the new data from this cruise the working group bathymetry was able to close some major gaps in the existing data as well as to discover new examination areas.

Bathymetric issues have been very important for the geological investigations on this cruise as shown in three examples.

(1) The eastern slope Yermak Plateau and the North East Land shelf of Svalbard was of special interest due to a suspected megaslide in this area (See Chapter 9). The shelf and its eastern and western slopes (Fig. 7.2) could be surveyed by using the multibeam sonar system Hydrosweep. Altogether the sea-ice conditions were good, so that reasonable profile measurements could be done on the Yermak Plateau.



**Fig. 7.2**: Hydroweep profiles in the Sophia Basin, southeast of the Yermak Plateau, and at the northern Svalbard continental margin

(2) On top of the central Yermak Plateau, probably out-cropped older sedimentary rocks should be sampled by means of the sediment dredge. The localization of petrologic sampling stations was accomplished based on the bathymetric maps (Fig. 7.3).

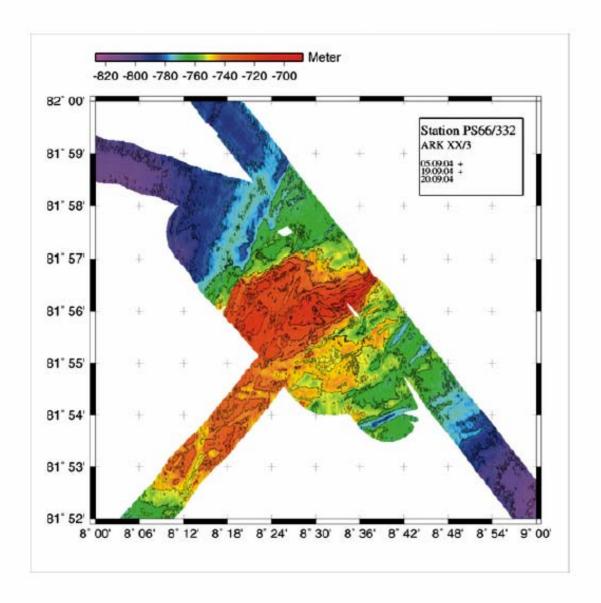
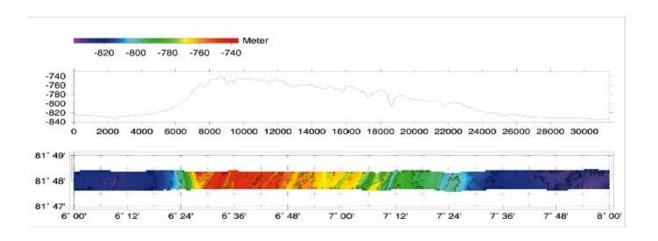


Fig. 7.3: Station PS 66/332 on the central Yermak Plateau, selected based on a bathymetric chart

(3) With systematic seismic profiles it was possible to determine the depth and course of iceberg plough marks on the Yermak plateau from the recorded bathymetric depth data (Fig. 7.4). On the western slope of the plateau the transition from the plateau area in Fram Strait-direction to the deep sea could be documented successfully.



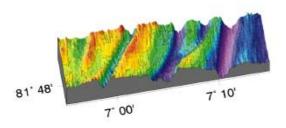


Fig. 7.4: Iceberg plough marks on Yermak Plateau

# 8. Marine Geophysics

W. Jokat, Y. Behr, H. Birnstiel, A. Gebauer, K. Gößling, D. Günther, H. Martens, N. Lensch, W. Raabe, M. Schmidt-Aursch, M. Schroeder, Th. Spengler

The only deep water connection of the Arctic Ocean with the world oceans is the Fram Strait. The deepest part of the Fram Strait hosts an active mid-ocean ridge, the Lena Trough. The spreading in the Lena Trough is the latest stage of the separation of Greenland and Svalbard, which started with large strike-slip movements some 50 Myr ago. According to published geodynamic models a deep-water connection for the Arctic Ocean was established some 10 to 15 Ma. This is inferred from reconstructions of the northern Atlantic and the relevant magnetic spreading anomalies in the North Atlantic and the Eurasia Basin. Details on the opening history of the Lena Trough, however, are not available, since spreading anomalies in the Fram Strait and across the southern basins (Greenland-Spitzbergen Sill, Boreas Basin) are missing.

The mid-ocean ridge in the Lena Trough is bounded by two prominent plateaus to the east and west: Yermak Plateau and Morris Jessup Rise. Based on the current knowledge these rises were formed some 35 Ma, when the relative

movements between Svalbard and Greenland became significant, and as a consequence, volcanic material erupted beneath the rises and/or formed them. There is no information available to constrain the amount of magmatic material that erupted during this time period. Speculations suggest that the north-eastern Yermak Plateau is oceanic in origin, while its southern part might consist of stretched continental crust. The same interpretations are valid for the Morris Jessup Rise. Most of these hypotheses base on aeromagnetic investigations, which show strong up to 1000 nT magnetic anomalies at both rises. Such strong anomalies are believed to be related to magmatism during continental break-up. However, no seismic data exist to constrain the interpretations in defining the nature of the crust.

The main scientific objective of this cruise was to enlarge the geophysical database in the Fram Strait to provide further constraints on the deeper structure and the evolution of this region. In detail the following problems should be addressed:

- Seismic investigations across the northern Yermak Plateau to identify the continent-ocean transition. The multichannel seismic network should be designed in a way that within a recent IODP preprosal the location of the sites can be determined in an optimum way.
- Deep seismic investigations to identify the deeper structure of the Yermak Plateau. With the wide-angle reflection data it should be possible to make estimates on the magmatic underplating, if any are present, of the Yermak Plateau.
- Try to sample basement rocks in order to date and constrain the evolution of the plateau.

Within the proposed IODP drilling proposal the following goals should be achieved:

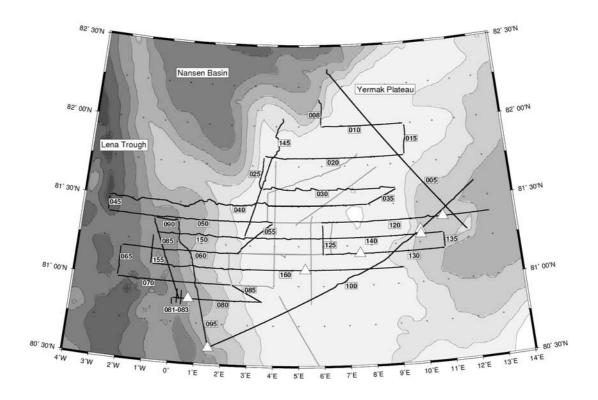
- To find suitable locations to drill the basement in order to resolve the origin of the plateau and to understand the origin of the strong magnetic anomalies along the northern rim of the plateau.
- To find suitable locations to reach old drift sediments in the Fram Strait in order to resolve the palaeoceanographic history of the Arctic Gateway.

#### Experiment set-up, constraints and statistics

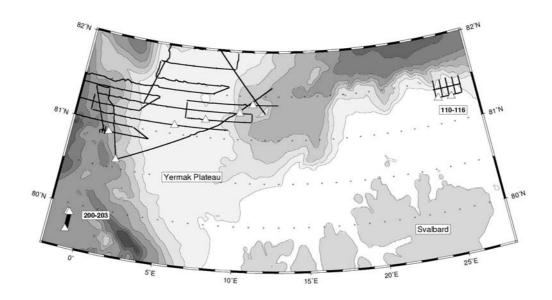
Before the expedition available ice maps showed already that during this season the ice edge in the Fram Strait retreated as far north as 82°N. Arriving in the research area of the north-eastern plateau this observation was more or less confirmed with the exception that compact ice fields drifting in the entire area were making seismic investigations difficult. To conduct deep seismic experiments in ice we needed ice floes with a sufficient snow cover to dig the geophones into it. Such kind of floes was completely missing. Almost no snow was left on the floes. Thus, the deep seismic experiments were not conducted since there was no chance to dig the geophones into the naked ice, and shelter them

against wind. As a consequence the geophysical experiments were concentrated on acquiring multi-channel seismic data with an 800 m long streamer. As source we used an airgun array with a total volume of 24 l. In addition, 20 long-range sonobuoys were deployed and successfully recovered (Table 8.2). Their signal ranges vary between 10 and 50 km.

The geophysical ship time can be split into three large portions a) the seismic network on the Yermak Plateau (Fig. 8.1, 260 h), b) dredging of basement highs on the Yermak Plateau (20 h), and c) the four site survey boxes along the East Greenland margin (Figs. 8. 2 and 8.3; 110 h). This sums up to 17 days of ship time including transit and seismic profiling. During this period in total 2920 km of new seismic data were acquired in the areas of investigation (Table 8.1). Despite an 5-10/10 ice cover the data quality is good, and there is a fair chance to suppress the seafloor multiple in order to image the basement of the Yermak Plateau along some profiles. The spacing of the W-E trending lines on the north-western part of the Yermak Plateau is approximately 11 km. Existing seismic networks acquired in 2002 and by Norwegian institutions were supplemented wherever possible. Areas south of 81°N were not investigated, since they are also accessible during normal sea ice seasons. Here, using a longer streamer of up to 3000 km is more appropriate, since a strong water bottom multiple has to be removed in order to image the thicker sediment packages.



**Fig.8.1**: Location of the seismic profiles on the Yermak Plateau. The small triangles indicate the location of sonobuoy deployment.



**Fig. 8.2**: Location of IODP site surveys on the Spitzbergen Sill and along the margin of Northern Svalbard

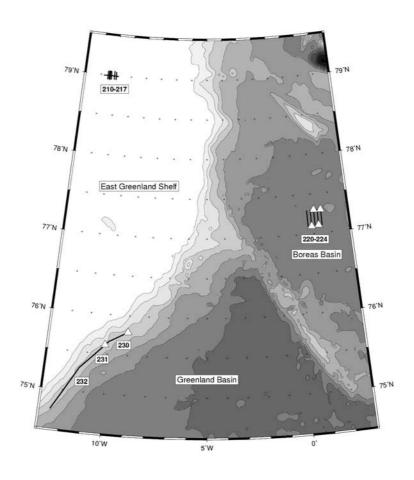


Fig. 8.3: Location of IODP site surveys along the East Greenland margin and on the East Greenland Shelf

Table 8.1: Profile and shot statistics

Profil	Beginn		End		Start		End		Shots	Length	Sonobuoys	Delay	Streamer	Lead In	Airgun	Chan	dx Chan
	Date	Time	Date	Time	Lon	Lat	Lon	Lat		[km]	Running Number	[s]	[m]	[m]			
20040005	05.09.2004	13:50:00	06.09.2004	04:35:00	81,3029	11,9847	82,3577	6,3005	3516	150,4	none	0	600	30	8x3l	96	6,25
20040008	06.09.2004	10:05:00	06.09.2004	11:57:00	82,1674	5,9172	82,0087	5,9910	446	18,7	none	0	600	30	8x3l	96	6,25
20040010	06.09.2004	12:07:00	06.09.2004	17:28:00	81,9977	6,0521	82,0005	9,6548	1276	56,35	none	0	600	30	8x3l	96	6,25
20040015	06.09.2004	17:38:00	06.09.2004	19:34:00	81,9890	9,7165	81,8071	9,6239	462	20,75	none	0	600	30	8x3l	96	6,25
20040020	06.09.2004	19:40:00	07.09.2004	05:14:00	81,8017	9,5782	81,8127	3,5446	2281	97,19	none	0	600	30	8x3l	96	6,25
20040025	07.09.2004	05:24:00	07.09.2004	07:30:00	81,8023	3,4816	81,6088	3,3753	502	22,4	none	1	600	30	8x3l	96	6,25
20040030	07.09.2004	07:41:00	07.09.2004	18:05:00	81,6045	3,4817	81,6003	9,1208	2246	98,11	none	1	600	30	8x3l	96	6,25
20040035	07.09.2004	18:16:00	07.09.2004	20:23:00	81,5908	9,1385	81,5055	7,9214	506	22,25	none	1	600	30	8x3l	96	6,25
20040040	07.09.2004	20:26:00	08.09.2004	16:25:00	81,5043	7,8922	81,5011	-2,9565	4436	191	none	3	600	30	8x3l	96	6,25
20040045	08.09.2004	16:37:00	08.09.2004	17:30:00	81,4869	-3,0025	81,4073	-2,9932	270	8,96	none	3	600	30	8x3l	96	6,25
20040050	08.09.2004	17:40:00	09.09.2004	05:16:00	81,3994	-2,9245	81,3979	3,9000	2765	116,1	none	4	600	30	8x3l	96	6,25
20040055	09.09.2004	05:25:00	09.09.2004	08:54:00	81,3881	3,9026	81,1968	2,4869	740	32,81	none	4	600	30	8x3l	96	6,25
20040060	09.09.2004	13:03:00	09.09.2004	21:27:00	81,1990	2,5337	81,1970	-2,2092	2003	82,28	none	0	600	30	8x3l	96	6,25
20040065	09.09.2004	21:36:00	09.09.2004	23:35:00	81,1883	-2,2743	81,0093	-2,2717	473	20,23	none	0	600	30	8x3l	96	6,25
20040070	09.09.2004	23:44:00	10.09.2004	07:28:00	80,9997	-2,2322	81,0003	2,2710	1842	79,48	none	0	600	30	8x3l	96	6,25
20040075	10.09.2004	07:35:00	10.09.2004	09:52:00	80,9974	2,3334	80,8976	3,5299	544	23,93	none	1	600	30	8x3l	96	6,25
20040080	10.09.2004	12:05:00	10.09.2004	18:30:00	80,9004	3,4686	80,8998	0,0110	1406	61,98	sb0401	0	600	30	8x3l	96	6,25
20040081	10.09.2004	18:36:00	10.09.2004	19:01:00	80,8978	-0,0406	80,8594	-0,0140	101	4,34	none	0	600	30	8x3l	96	6,25
20040082	10.09.2004	19:25:00	10.09.2004	20:25:00	80,8668	0,1661	80,9571	0,1667	231	10,17	none	0	600	30	8x3l	96	6,25
20040083	10.09.2004	20:50:00	10.09.2004	21:51:00	80,9498	0,3336	80,8596	0,3305	243	10,14	none	0	600	30	8x3l	96	6,25
20040085	10.09.2004	22:27:00	11.09.2004	04:30:00	80,8932	0,2574	81,3848	-1,0284	1436	60,34	none	0	600	30	8x3l	96	6,25
20040090	11.09.2004	04:39:00	11.09.2004	06:17:00	81,3904	-0,9523	81,3865	-0,0368	390	15,94	none	1	600	30	8x3l	96	6,25
20040095	11.09.2004	06:24:00	11.09.2004	15:33:00	81,3776	-0,0082	80,6089	1,4884	2085	93,54	sb0402	1	600	30	8x3l	96	6,25
20040100	11.09.2004	15:40:00	12.09.2004	13:22:00	80,6021	1,5312	81,5997	12,4949	5161	223,9	sb0402 / sb0403	2	600	30	8x3l	96	6,25
20040110	14.09.2004	15:45:00	14.09.2004	17:50:00	81,5010	25,0007	81,3009	25,0030	498	22,44	none	0	600	30	8x3l	96	6,25
20040111	14.09.2004	17:56:00	14.09.2004	20:59:00	81,2972	25,0379	81,3008	26,9711	727	32,83	sb0404	1	600	30	8x3l	96	6,25
20040112	14.09.2004	21:03:00	14.09.2004	23:07:00	81,3052	27,0002	81,4998	26,9996	494	21,83	sb0404	1	600	30	8x3l	96	6,25
20040113	15.09.2004	00:03:00	15.09.2004	02:07:00	81,5003	26,4956	81,3000	26,5033	494	22,47	sb0404	1	600	30	8x3l	96	6,25
20040114	15.09.2004	03:09:00	15.09.2004	05:23:00	81,2998	26,0019	81,5074	26,0028	533	23,29	sb0404 / sb0405	1	600	30	8x3l	96	6,25

Profil	Beginn		End		Start		End		Shots	Length	Sonobuoys	Delay	Streamer	Lead In	Airgun	Chan	dx Chan
	Date	Time	Date	Time	Lon	Lat	Lon	Lat		[km]	Running Number	[s]	[m]	[m]			
20040115	15.09.2004	06:14:00	15.09.2004	08:26:00	81,502	25,4988	81,2951	25,4972	525	23,21	sb0404 / sb0405	1	600	30	8x3l	96	6,25
20040116	15.09.2004	10:18:00	15.09.2004	13:38:14	81,3996	24,8971	81,4001	27,0641	749	36,34	none	2	600	30	8x3l	96	6,25
20040120	17.09.2004	08:54:00	17.09.2004	20:09:45	81,4002	12,9789	81,3995	5,9591	2679	118,93	sb0406	0	600	30	8x3l	96	6,25
20040125	17.09.2004	20:32:00	17.09.2004	22:32:00	81,3805	5,9898	81,1946	6,0015	477	20,88	sb0406	1	600	30	8x3l	96	6,25
20040130	17.09.2004	22:45:00	18.09.2004	06:40:00	81,1984	6,1144	81,2006	10,972	1865	83,79	sb0406 / sb0407	1	600	30	8x3l	96	6,25
20040135	18.09.2004	06:47:00	18.09.2004	07:43:00	81,2079	11,0016	81,2970	10,9967	221	10,02	sb0406 / sb0407	1	600	30	8x3l	96	6,25
20040140	18.09.2004	07:49:00	18.09.2004	21:33:00	81,3000	10,9570	81,2993	2,8190	3192	139,11	sb0406 / sb0407 / sb0408	2	600	30	8x3l	96	6,25
20040145	18.09.2004	21:44:00	19.09.2004	07:25:00	81,3099	2,7736	82,0513	4,2940	2293	92,32	none	3	600	30	8x3l	96	6,25
20040150	20.09.2004	16:45:00	20.09.2004	23:26:45	81,3008	2,8373	81,2991	-0,8591	1598	65,21	none	0	600	30	8x3l	96	6,25
20040155	21.09.2004	00:04:00	21.09.2004	02:06:00	81,2780	-1,0098	81,1008	-1,0027	486	20,19	none	0	600	30	8x3l	96	6,25
20040160	21.09.2004	02:24:00	21.09.2004	20:00:00	81,0979	-0,9011	81,1001	9,2497	4194	178,36	sb0411	1	600	30	8x3l	96	6,25
20040200	22.09.2004	10:43:00	22.09.2004	12:46:00	79,8994	-0,8598	79,7000	-0,8467	490	22,41	sb0412	0	600	30	8x3l	96	6,25
20040201	22.09.2004	13:00:00	22.09.2004	15:10:00	79,7005	-0,8001	79,9054	-0,7998	517	22,98	sb0412 / sb0413	0	600	30	2xG+1xGI	96	6,25
20040202	22.09.2004	15:21:00	22.09.2004	17:31:00	79,8998	-0,7522	79,6955	-0,7522	517	22,95	sb0412 / sb0413 / sb0414	0	600	30	2xG+1xGI	96	6,25
20040203	22.09.2004	17:42:45	22.09.2004	19:48:00	79,6983	-0,7009	79,9051	-0,7022	499	23,20	none	0	600	30	2xG+1xGI	96	6,25
20040204	22.09.2004	20:02:00	22.09.2004	22:17:00	79,8991	-0,6549	79,6986	-0,6492	537	22,49	none	1	600	30	2xG+1xGI	96	6,25
20040210	23.09.2004	14:13:00	23.09.2004	15:13:00	79,0476	-10,9979	78,9496	-10,9976	715	11,00	none	0	600	30	1x2,4l (G)	96	6,25
20040211	23.09.2004	15:37:00	23.09.2004	16:45:00	78,9499	-11,1114	79,0556	-11,1074	808	11,90	none	0	600	30	1x2,4l (G)	96	6,25
20040212	23.09.2004	17:14:00	23.09.2004	18:22:00	79,0547	-11,3299	78,9467	-11,3216	808	12,16	none	0	600	30	1x2,4l (G)	96	6,25
20040213	23.09.2004	18:40:00	23.09.2004	19:51:00	78,9371	-11,4034	79,0555	-11,3895	841	13,14	none	0	600	30	1x2,4l (G)	96	6,25
20040214	23.09.2004	20:00:00	23.09.2004	21:10:00	79,0537	-11,4242	78,9466	-11,4238	832	12,04	none	0	600	30	1x2,4l (G)	96	6,25
20040215	23.09.2004	21:22:00	23.09.2004	22:28:00	78,9497	-11,4579	79,0538	-11,4615	787	11,70	none	0	600	30	1x2,4l (G)	96	6,25
20040216	23.09.2004	22:48:00	23.09.2004	23:54:00	79,0504	-11,5496	78,9453	-11,5483	750	11,80	none	0	600	30	1x2,4l (G)	96	6,25
20040217	24.09.2004	00:54:00	24.09.2004	02:48:00	78,9928	-11,8007	78,9933	-10,8556	1354	20,31	none	1	600	30	1x2,4l (G)	96	6,25
20040220	25.09.2004	04:45:00	25.09.2004	06:25:00	77,2987	1,4938	77,1397	1,5025	399	17,86	sb0415	0	600	30	2xG+1xGI	96	6,25
20040221	25.09.2004	07:27:00	25.09.2004	09:41:00	77,0984	1,2958	77,3053	1,2994	533	23,19	sb0415 / sb0416	0	600	30	2xG+1xGI	96	6,25
20040222	25.09.2004	10:13:00	25.09.2004	12:29:00	77,3058	1,0962	77,0963	1,1070	541	23,52	sb0415 / sb0416 / sb0417	1	600	30	2xG+1xGI	96	6,25
20040223	25.09.2004	13:08:00	25.09.2004	15:20:00	77,1005	0,9036	77,3051	0,9002	525	22,96	sb0417 / sb0418	1	600	30	2xG+1xGI	96	6,25
20040224	25.09.2004	15:54:30	25.09.2004	18:06:00	77,3000	0,6998	77,0947	0,6990	524	23,04	sb0417 / sb0418	1	600	30	2xG+1xGI	96	6,25
20040230	26.09.2004	09:38:00	26.09.2004	12:25:00	75,8033	-8,9457	75,6611	-9,9653	665	32,43	sb0419	0	600	30	8x3l	96	6,25
20040231	26.09.2004	12:25:00	26.09.2004	17:55:00	75,6611	-9,9653	75,2996	-11,3196	1178	55,70	sb0419 / sb0420	0	600	30	8x3l	96	6,25
20040232	26.09.2004	17:55:00	27.09.2004	00:05:15	75,2996	-11,3196	74,7475	-12,4960	1472	70,69	none	0	600	30	8x3l	96	6,25

Table 8.2: Statistics on the sonobuoy deployments

Profil	Beginn		End		Sonobuoy			Shots	Length of Profil	Start of R	egistration	Airguns	Streamer
	Date	Time	Date	Time	Тур	Buoy	Number		w hile registrating	Lon	Lat		
20040080	10.09.2004	12:34:48	10.09.2004	17:26:40	DOSH	sb1	sb0401	1033	44,97	80°52N	00°36E	VFL	600
20040095	11.09.2004	11:59:02	11.09.2004	15:33:00	DOSH	sb1	sb0402	802	37,43	80°36N	01°29E	VFL	600
20040100	11.09.2004	15:40:00	11.09.2004	18:49:20	DOSH	sb1	sb0402	752	33,66	81°36N	10°30E	VFL	600
	12.09.2004	08:36:43	12.09.2004	13:22:00	DOSH	sb1	sb0403	1126	45,80	81°36N	12°28E	VFL	600
20040111	14.09.2004	18:08:00	14.09.2004	20:59:00	DOSH	sb3	sb0404	681	30,73	81°17N	25°09E	VFL	600
20040112	14.09.2004	21:03:00	14.09.2004	23:07:00	DOSH	sb3	sb0404	494	21,83	81°19N	27°00E	VFL	600
20040113	15.09.2004	00:03:00	15.09.2004	02:07:00	DOSH	sb3	sb0404	494	22,47	81°30N	26°30E	VFL	600
20040114	15.09.2004	03:09:00	15.09.2004	05:23:00	DOSH	sb3	sb0404	533	23,29	81°18N	26°00E	VFL	600
	15.09.2004	03:10:00	15.09.2004	05:23:00	DOSH	sb1	sb0405	529	23,11	81°18N	26°00E	VFL	600
20040115	15.09.2004	06:14:00	15.09.2004	08:26:00	DOSH	sb3	sb0404	525	23,21	81°30N	25°30E	VFL	600
	15.09.2004	06:14:00	15.09.2004	08:21:00	DOSH	sb1	sb0405	505	22,26	81°30N	25°30E	VFL	600
20040120	17.09.2004	11:57:38	17.09.2004	20:09:45	DOSH	sb1	sb0406	1949	86,20	81°24N	11°01E	VFL	600
20040125	17.09.2004	20:32:00	17.09.2004	22:32:00	DOSH	sb1	sb0406	477	20,88	81°23N	05°59E	VFL	600
20040130	17.09.2004	22:45:00	18.09.2004	06:40:00	DOSH	sb1	sb0406	1865	83,79	81°12N	06°07E	VFL	600
	17.09.2004	01:03:00	18.09.2004	06:40:00	DOSH	sb3	sb0407	1317	60,45	81°12N	06°07E	VFL	600
20040135	18.09.2004	06:47:00	18.09.2004	07:43:00	DOSH	sb1	sb0406	221	10,21	81°13N	11°00E	VFL	600
	18.09.2004	06:47:00	18.09.2004	07:43:00	DOSH	sb3	sb0407	221	10,21	81°13N	11°00E	VFL	600
20040140	18.09.2004	07:49:00	18.09.2004	08:16:00	DOSH	sb1	sb0406	108	4,81	81°18N	10°57E	VFL	600
	18.09.2004	07:49:00	18.09.2004	17:22:00	DOSH	sb3	sb0407	2194	96,92	81°18N	10°57E	VFL	600
	18.09.2004	09:23:00	18.09.2004	17:00:00	DOSH	sb1	sb0408	1735	76,21	81°17N	9°58E	VFL	600
20040160	21.09.2004	13:35:00	21.09.2004	16:44:47	DOSH	sb2	sb0411	755	33,63	81°06N	5°16E	VFL	600
20040200	22.09:2004	10:43:00	22.09:2004	12:46:00	DOSH	sb1	sb0412	490	22,41	79°53N	0°52W	VFL	600
20040201	22.09:2004	13:00:00	22.09:2004	15:10:00	DOSH	sb1	sb0412	517	22,98	79°42N	0°48W	2xG+1xGI	600
	22.09:2004	13:00:00	22.09:2004	15:10:00	DOSH	sb2	sb0413	517	22,98	79°42N	0°48W	2xG+1xGI	600
20040202	22.09:2004	15:21:00	22.09:2004	16:28:00	DOSH	sb1	sb0412	267	12,15	79°54N	0°45W	2xG+1xGI	600
	22.09:2004	15:21:00	22.09:2004	16:43:00	DOSH	sb2	sb0413	326	14,76	79°54N	0°45W	2xG+1xGI	600
	22.09:2004	15:21:00	22.09:2004	17:30:00	DOSH	sb4	sb0414	513	22,78	79°54N	0°45W	2xG+1xGI	600
20040220	25.09.2004	04:47:00	25.09.2004	06:25:00	DOSH	sb1	sb0415	391	17,52	77°17N	1°29E	2xG+1xGI	600
20040221	25.09.2004	07:27:00	25.09.2004	09:41:00	DOSH	sb1	sb0415	533	23,19	77°06N	1°18E	2xG+1xGI	600
	25.09.2004	07:27:00	25.09.2004	09:41:00	DOSH	sb2	sb0416	533	23,19	77°06N	1°18E	2xG+1xGI	600
20040222	25.09.2004	10:13:00	25.09.2004	10:20:00	DOSH	sb1	sb0415	29	1,23	77°18N	1°06E	2xG+1xGI	600
	25.09.2004	10:13:00	25.09.2004	10:40:00	DOSH	sb2	sb0416	108	4,74	77°18N	1°06E	2xG+1xGI	600
	25.09.2004	10:16:00	25.09.2004	12:29:00	DOSH	sb4	sb0417	529	33,00	77°18N	1°05E	2xG+1xGI	600
20040223	25.09.2004	13:08:00	25.09.2004	15:20:00	DOSH	sb4	sb0417	525	22,96	77°06N	0°54E	2xG+1xGI	600
	25.09.2004	13:08:00	25.09.2004	15:20:00	DOSH	sb1	sb0418	525	22,96	77°06N	0°54E	2xG+1xGI	600
20040224	25.09.2004	15:54:30	25.09.2004	16:26:00	DOSH	sb4	sb0417	127	5,70	77°18N	0°42E	2xG+1xGI	600
	25.09.2004	15:54:30	25.09.2004	16:13:00	DOSH	sb1	sb0418	75	3,36	77°18N	0°42E	2xG+1xGI	600
20040230	26.09.2004	09:53:00	26.09.2004	12:25:00	DOSH	sb1	sb0419	605	29,39	75°47N	9°01W	VFL	600
20040231	26.09.2004	12:25:00	26.09.2004	13:58:00	DOSH	sb1	sb0419	237	11,27	75°40N	9°58W	VFL	600
	26.09.2004	13:39:00	26.09.2004	16:28:00	DOSH	sb2	sb0420	671	31,49	75°36N	10°15W	VFL	600

The Yermak Plateau east of 10°E was this season not accessible. Thick multiyear ice floes did not allow any single ship seismic investigations. This is also valid for the areas north of 82°N. Only on two profiles this part of the Arctic Ocean could be entered till the ship got stuck in thick ice floes.

Parallel to the seismic investigations gravity and magnetic data were continuously acquired by fixed installed instruments. Harbour values for the gravimeter were taken in Tromsoe and Bremerhaven. Both instruments worked without any problems or failures.

#### Results

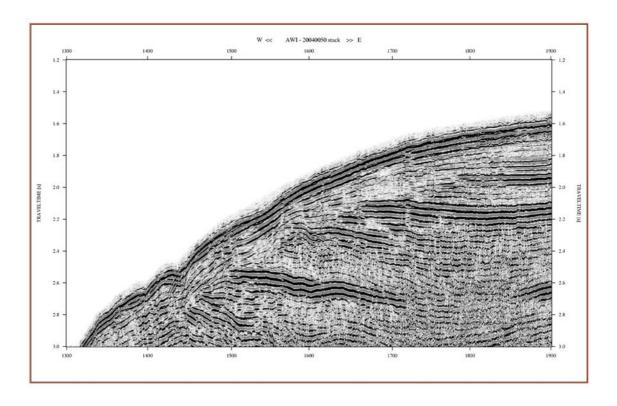
The ice conditions during this season allowed the first time to conduct a detailed seismic survey across the north-western part of the Yermak Plateau (Fig. 8.1). In addition two long almost N-S trending transect were acquired across the entire plateau, which provided a surprising image of the deeper structure of the Yermak Plateau at 10°E longitude. The seismic data show that the ancient topography is quite rough. Ridges and mountains with altitudes up to 1000 m are present. The present plateau-like shape is a consequence of continuous sedimentation mostly driven by currents, which filled the deep depressions between the topographic highs.

The seismic network west of 6°E indicates some kind of change in the style of drift deposition. Close to topographic highs the sedimentation pattern shows that currents were guided along these obstacles. The fast changes in the geometry of sediment drifts indicate quite dynamic processes, which are active till present. However, the drift sediment could build up only to a depth range of 600 to 800 m. In the shallowest parts the deposits were eroded by icebergs. More or less constant erosion by the icebergs prevented any formation of shallower sediment drifts. These processes caused on several locations that old drift sediments crop out at the surface or are only covered by a thin (200 m) sedimentary drape. Such locations are present on the Yermak Plateau, where old sedimentary units terminate close to the seafloor. However, their exact age is not known. From geodynamic models it can be assumed that the basement blocks were formed during the final split of the Yermak and Morris Jessup plateaus and/or during the earlier rifting period. Thus, the oldest sediments on the Yermak Plateau might be around 35 Ma old. The basement, if continental. might be as old as Devonian.

A different situation is found in the west towards the Lena Trough (Fig. 8.1). This mid-ocean ridge system actually formes the deep part of the Fram Strait and channels the deep-water exchange between the North Atlantic and the Arctic Ocean. Again the exact timing how this ridge evolved during the geological times is not known. The older oceanic crust is covered on both sides of the ridge by thick sediments. Magnetic spreading anomalies terminate north of 83°N. Thus, no detailed information on the development of the Lena Trough between 83°N and 80°N from the magnetic data is available. Especially this area is critical to all models, which deal with the initiation of the deep-water exchange of both oceans. The seismic profiles were planned to stretch from the central part of the Yermak Plateau to the present day rift shoulder of the Lena Trough.

Only one profile enters also the rift valley. Two observations from the seismic data are important:

- The geological history of both rift flanks seems not to be symmetric for the Greenland and the Svalbard margins. From the seismic data which were gathered in the last years, the North-Greenland margin is more rugged and is less covered by sediments than it counterpart on the Yermak Plateau. Models which suggest that the northward running Atlantic water carries more sediment load than the Arctic water masses are confirmed by the seismic data sets.
- Along the western margin of the Yermak Plateau large drift deposits were formed through the sediment load of the Atlantic water masses. They are around 2000 m thick and show a very constant depositional pattern. However, there are two exceptions. Close to the flank of the Lena Trough the sediments are slightly folded. More eastward an interesting area is indicated by a smooth change in the depth gradient of the sea floor. Below this stronger increase in water depth, the seismic units terminate or thin significantly. Old drift deposits, which might have formed during Miocene times, are close below the sea floor (Fig. 8.4, ca. 200 m drape). From the present stage of interpretation it is not clear what kind of process caused this image. A likely scenario is that the current speed increased due to a faster subsidence of the ridge or an increased amount of water exchange. Thus, the new drift sediments were deposited at greater depths and the older drift deposits were partly eroded.



**Fig.8.4**: Seismic profile on the eastern flank of the Lena Trough. Please note that deeper strata terminate and/or thin towards the seafloor.

These first ideas on the explanation of the seismic images have to be verified during future analyses. An important aspect to do so is that the new seismic network has several tie lines to the existing ODP drill holes. Here, the drill hole 910 is of special interest since the oldest sediments were drilled of Plio/Pleistocene age. This information can be extrapolated into the seismic lines of the network to provide at least time constraints for the last 5 Mio. years.

To provide constraints on the crustal fabric of the Yermak Plateau some seismic refraction work was planned before the cruise. The intention was to place RefTek recording stations on ice floes and to shoot with a large volume airgun. It turned out that this season the ice floes in the area of investigation were to small and/or were not covered by any snow, which would allow to hide the geophones from the wind. Thus, the only wide-angle data were gathered with a new type of sonobuoys. These instruments do not directly transmit the data to the ship but store them on an internal disc. They have to be retrieved by helicopter. The advantage is that longer offsets can be recorded than the line of sight, which is typical at 25-35 km. The instruments recorded data up to 70 km. Longer offsets were not tested due to the profile planning. The signal quality is highly variable due to different geological settings and the level of environmental noise.

## Helicopter-borne Magnetic Data

Parallel to the seismic profiling across the Yermak Plateau an aeromagnetic survey was performed by helicopter. The aim was to complete and fulfil planned lines of a past magnetic survey from 1999 (Fig. 8.5). The scientific goal was to densify the existing magnetic data sets in order to resolve more details on the magnetic structure of the Yermak Plateau, especially in the area of the strong positive anomalies (ca. 1000 nT strong anomalies). The complete area of interest has its boundaries along 81°N to 83°N and 5°W to 25°E. The block-corners were transformed to UTM coordinates and a central meridian of 21 was used. The line direction is on this longitude 0° which changes with higher distances to the central meridian. The line-spacing was 7.5 km. The flight altitude was mainly 100 m and the data were acquired at a speed of 80 ktns (~ 40 m/s). A new flying procedure was introduced. At the end of each line the pilots made a 270° turn and cross the lines in order to get more intersections. This will lead to an improvement for levelling the data.

Within nine days and 20 flights most of the survey area was successfully completed and more than 7250 line-kilometres were flown. An overview of flights per day and flying hours per day is given in Table8.3. Unfortunately the profile lines could not be finished to the northern end due to bad weather and long transit respectively. Crossing the lines of the past survey gives us now the opportunity to level the data and combine them into one database.

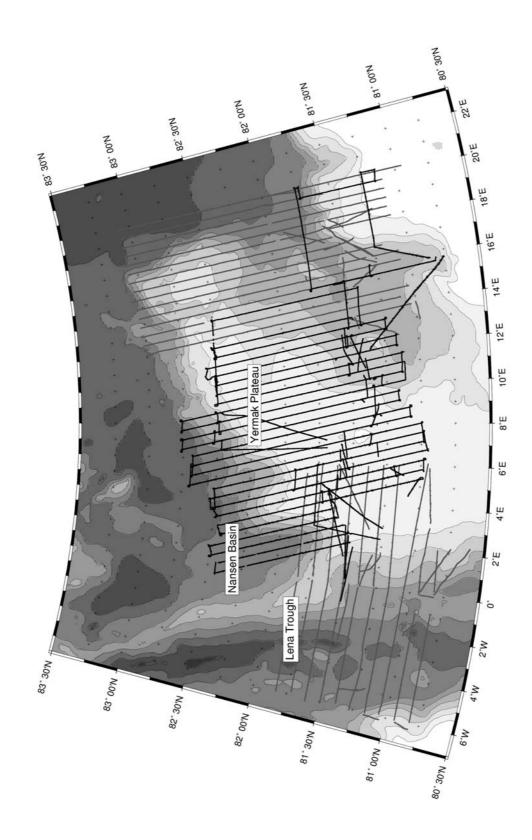


Fig. 8.5: Survey area for the magnetic measurements performed with the "Polarstern" helicopters

Table 8.3: Flight and data statistics for the magnetic survey

text filename	binary filename	line #	fiducial start	end	time start	end	flight #	ascii raw-filename
<b>03.09.2004</b> S4090310.T49	testflight	radartest						0409031.raw
<b>07.09.2004</b> S4090708.T17	S4090708.B17 S4090709.B13	110.0N 120.0S	1 3332	3331 5685	08:18:55 09:14:27	09:14:25 09:53:40	1	0409071.raw
S4090712.T04 S4090712.T27	S4090712.B04 S4090712.B27	??? 130.0N	1 1	1328 2461	12:05:37 12:29:19	12:27:44 13:10:19	2	0409072.raw 0409073.raw
S4090713.T40	S4090713.B08 S4090713.B40	140.0S 140.0S	2462 1	4331 1077	13:10:21 13:42:09	13:41:29 14:00:05		0409074.raw
S4090714.T44	S4090714.B45 S4090715.B47	150.0N 140.0N	1 3750	3749 4088	14:46:47 15:49:17	15:49:15 15:54:55	3	0409076.raw
	S4090715.B53 S4090715.B5D	150.0S 160.0S	4089 4090	4089 7548	15:54:57 15:54:59	15:54:57 16:52:37		024 km
08.09.2004							Н	924 km
S4090808.T48	S4090808.B48 S4090808.B49 S4090808.B50 S4090809.B07 S4090809.B0H S4090809.B0Y S4090810.B03	160.0S 150.0S 140.0S 130.0S 120.0S 110.0S 100.0N 90.0S	1 23 28 32 1101 1103 1107 4431	22 27 31 1100 1102 1106 4430 8845	08:50:40 08:51:03 08:51:09 08:51:51 09:09:41 09:09:44 09:09:49 10:05:14	08:51:01 08:51:07 08:51:12 09:09:39 09:09:42 09:09:47 10:05:12 11:18:48	1	0409081.raw
S4090811.T34	S4090811.B34 S4090811.B3E S4090811.B3V	100.0S 90.0S ???	1 26 28	25 27 683	11:36:11 11:36:37 11:36:40	11:36:35 11:36:38 11:47:35	2	0409083.raw
S4090811.T48	S4090811.B49 S4090811.B4J	90.0N ???	1 4	3 971	11:51:01 11:51:05	11:51:03 12:07:12		0409084.raw
S4090812.T06	S4090812.B06 S4090812.B0G	90.0S 80.0N	1 13	12 856	12:08:38 12:08:51	12:08:49 12:22:54		0409085.raw
S4090812.T22 S4090812.T35	S4090812.B22 S4090812.B35 S4090812.B3F	80.0N 80.0S 70.0S	1 1 6	720 5 934	12:24:18 12:37:38 12:37:44	12:36:17 12:37:42 12:53:12		0409086.raw 0409087.raw
S4090812.T52 S4090813.T11 S4090813.T30	S4090812.B52 S4090813.B11 S4090813.B30	??? 70.0S 70.0N	2 1 3	1048 1031 1221	12:54:54 13:13:49 13:32:58	13:12:20 13:30:59 13:53:17		0409088.raw 0409089.raw 0409080.raw
0.0000.00.00	0.0000.000	, 0.0.1			10.02.00	10.00.11		782 km
<b>09.09.2004</b> S4090908.T27	S4090908.B28 S4090908.B46 S4090909.B17	60.0N 160.0S 150.0N	1 1106 2938	1105 2937 2941	08:30:28 08:49:19 09:19:52	08:48:52 09:19:50 09:19:55	1	0409091.raw
	S4090909.B1H	140.0S	2942	5369	09:19:57	10:00:24		
10.09.2004	testflight	cable and ra	adintest				$\vdash\vdash\vdash$	260 km 040910*.raw
12.09.2004	tootingrit	Jabio dilu I					$\vdash \vdash$	5-7010 .1aw
S4091209.T12	S4091209.B12 S4091210.B26	290.0N 280.0S	1 4442	4441 7432	09:15:42 10:29:44	10:29:42 11:19:34	1	0409121.raw
S4091211.T46	S4091211.B47 S4091211.B4H S4091211.B4Y	280.0N 290.0S 300.0N	3 40 44	39 43 3294	11:50:43 11:51:21 11:51:26	11:51:19 11:51:24 12:45:36	2	0409122.raw
S4091212.T48	S4091212.B42 S4091212.B4C S4091212.B48 S4091212.B4I S4091212.B49 S4091212.B4J	290.0N 300.0S 300.0S 310.0S 320.0S 330.0S	3295 3311 1 8 27 34	3310 3646 7 26 33 37	12:45:38 12:45:55 12:52:08 12:52:16 12:52:36 12:52:44	12:45:53 12:51:30 12:52:14 12:52:34 12:52:42 12:52:47		0409123.raw
S4091212.T49	S4091212.B50	330.0S	1	15	12:53:33	12:53:47		0409124.raw

text filename	binary filename	line #	fiducial start	end	time start	end	flight #	ascii raw-filename
	S4091212.B5A	340.0S	16	2692	12:53:49	13:38:25		
	S4091213.B34	330.0S	2693	2695	13:38:26	13:38:29		
	S4091213.B35	320.0S	2696	2698	13:38:31	13:38:33		
	S4091213.B3F	310.0S	2699	2701	13:38:35	13:38:37		
	S4091213.B3W	300.0N	2702	4122	13:38:39	14:02:19		
S4091214.T30	S4091214.B30	300.0N	1	2	14:33:49	14:33:50	3	0409125.raw
04001214.100	S4091214.B3A	290.0N	3	4	14:33:52	14:33:53	ΙĭΙ	0400120.141
	S4091214.B3R	280.0N	5	5	14:33:55	14:33:55		
	S4091214.B3c	270.0N	6	6	14:33:57	14:33:57		
	S4091214.B3t	260.0N	7	3962	14:33:59	15:39:54		
	S4091215.B36	270.0S	3963	6657	15:39:56	16:24:51		
S4091216.T22	S4091216.B23	270.0S	1	741	16:26:33	16:38:54		0409126.raw
14.09.2004	testflight	radiotest					$\vdash$	1043 km
16.09.2004	testilight	Tadiotest					Н	
S4091608.T51	S4091608.B51	350.0N	1	1129	08:55:06	09:13:54	1	0409162.raw
S4091609.T11	S4091609.B11	350.0N	1	910	09:15:53	09:31:02		0409163.raw
	S4091609.B27	360.0N	911	912	09:31:04	09:31:05		
	S4091609.B2H	370.0N	913	914	09:31:07	09:31:08		
	S4091609.B2Y	380.0N	915	916	09:31:10	09:31:11		
	S4091609.B2j	390.0N	917	917	09:31:13	09:31:13		
	S4091609.B2a	400.0N	918	918	09:31:15	09:31:15		
	S4091609.B2b S4091609.B2c	410.0N 420.0N	919 921	920 924	09:31:17 09:31:20	09:31:18 09:31:23		
	S4091609.B2d	430.0N	925	928	09:31:25	09:31:28		
	S4091609.B2e	440.0N	929	931	09:31:30	09:31:32		
	S4091609.B2g	460.0S	936	3074	09:31:39	10:07:17		
	S4091610.B03	450.0S	3075	6399	10:07:19	11:02:43		
	S4091610.B58	440.0S	6400	6402	11:02:45	11:02:47		
	S4091610.B5I	430.0S	6403	6404	11:02:49	11:02:50		
	S4091610.B5Z	420.0S	6405	6407	11:02:52	11:02:54		
	S4091610.B59	410.0N	6408	6410	11:02:56	11:02:58		
	S4091610.B5J	400.0N	6411	6414	11:03:00	11:03:03		
	S4091611.B00	360.0S	6415	8443	11:03:58	11:37:46		
S4091611.T55	S4091611.B55	360.0N	1	4	11:59:57	12:00:00	2	0409164.raw
01001011.100	S4091611.B56	350.0N	5	1887	12:00:01	12:31:25	~	0400104.114.11
	S4091612.B27	340.0S	1888	1890	12:31:27	12:31:29		
	S4091612.B2H	330.0S	1891	1893	12:31:31	12:31:33		
	S4091612.B2Y	320.0S	1894	1895	12:31:35	12:31:36		
	S4091612.B2j	310.0S	1896	1897	12:31:38	12:31:39		
	S4091612.B2a	300.0N	1898	3616	12:31:40	13:00:19		
	S4091612.B56	290.0N	3617	4617	13:00:21	13:00:21		
	S4091613.B13 S4091613.B34	280.0S 270.0S	4618 5908	5907 8850	13:00:22 13:38:34	13:38:32 14:27:36		
	3409 IO I3.D34	210.03	3900	0000	13.30.34	14.27.30		859 km
17.09.2004							$\Box$	
S4091708.T05	S4091708.B05	310.0N	1	107	08:10:00	08:11:46	1	0409171.raw
	S4091708.B07	300.0N	108	136	08:11:48	08:12:16		
	S4091708.B08	310.0N	137	2672	08:12:18	08:54:33		
	S4091708.B50	300.0S	2673	2673	08:54:34	08:54:34		
	S4091708.B5A	290.0S	2674	2674	08:54:35	08:54:35		
	S4091708.B5R S4091708.B5c	280.0S 270.0S	2675 2676	2675 2676	08:54:36 08:54:37	08:54:36 08:54:37		
	S4091708.B5t	260.0S	2677	2677	08:54:38	08:54:38		
	S4091708.B5U	250.0S	2678	7565	08:54:45	10:16:12		
S4091710.T30		250.0S 250.0N	1	28	10:35:11	10:35:38	2	0409172.raw
	S4091710.B3B							
	S4091710.B3B S4091710.B3S	240.0N	29	2816	10:35:40	11:22:07		
S4091711.T18								0409173.raw

flight ascii # raw-filename	end	time start	end	fiducial start	line #	binary filename	text filename
	12:48:16	11:36:14	4350	28	230.0S	S4091711.B32	
	12:58:53	12:48:18	4986	4351	240.0N	S4091712.B44	
0400475	40.44.50	40.44.50	0	,	040.001	04004740 D07	S4091713.T07
3 0409175.raw	13:11:58 13:12:02	13:11:56 13:12:00	3 6	1 4	240.0N 230.0N	S4091713.B07 S4091713.B0H	\$4091713.107
	14:04:39	13:12:04	3162	7	220.0N	S4091713.B0H S4091713.B0Y	
	15:13:31	14:04:41	7293	3163	210.0S	S4091714.B00	
	15:31:10	15:13:33	8351	7294	220.0N	S4091715.B09	
1149 km							
							18.09.2004
1 0409182.raw	09:02:54	08:47:35	920	2	240.0S	S4091808.B43	S4091808.T43
	09:02:55	09:02:55	921	921	250.0S	S4091808.B58	
	09:26:40	09:02:58	2344	922	260.0N	S4091808.B5I	
	09:54:34	09:26:42	4017	2345	250.0S	S4091809.B22	
	09:54:40 10:10:50	09:54:36 09:54:42	4022 4991	4018 4023	240.0S 230.0N	S4091809.B50 S4091809.B5A	
	10:10:50	10:10:52	4997	4992	220.0N	S4091810.B06	
	10:11:03	10:10:52	5002	4998	210.0S	S4091810.B0G	
	10:38:30	10:11:05	6648	5003	200.0S	S4091810.B0X	
	11:04:51	10:38:32	8228	6649	190.0N	S4091810.B34	
2 0409183.raw	12:21:04	11:25:50	3315	1	200.0N	S4091811.B21	S4091811.T21
	13:33:24	12:21:06	7654	3316	190.0S	S4091812.B16	
3 0409184.raw	13:47:49	13:47:45	5	1	190.0S	S4091813.B43	S4091813.T43
	14:42:50	13:47:51	3305	6	180.0N	S4091813.B4D	
4400 km	16:19:40	14:42:52	9114	3306	170.0S	S4091814.B38	
1160 km							20.09.2004
1 0409203.raw	08:48:35	08:14:09	2069	3	200.0N	S4092008.B09	S4092008.T09
	09:04:01	08:48:37	2994	2070	210.0N	S4092008.B44	01002000.100
	09:24:45	09:04:03	4237	2995	220.0S	S4092008.B59	
	09:43:21	09:24:47	5352	4238	230.0N	S4092009.B20	
	10:37:29	09:43:22	8599	5353	240.0S	S4092009.B38	
2 0409204.raw	10:55:44	10:55:36	9	1	240.0S	S4092010.B50	S4092010.T50
	10:55:46	10:55:46	10	10	230.0S	S4092010.B51	
	10:55:49	10:55:48	12	11	220.0S	S4092010.B5B	
	10:55:52	10:55:51	14	13	210.0S	S4092010.B5S	
	10:55:00	10:55:00	16 17	15 17	200.0S	S4092010.B5d	
	10:55:56 10:55:59	10:55:56 10:55:59	17 18	17 18	190.0S 180.0S	S4092010.B5u S4092010.B5x	
	10:55:59	10:55:59	21	19	170.0S	S4092010.B5x S4092010.B5y	
	11:29:22	10:56:04	2019	22	160.0S	S4092010.B5y	
	11:29:25	11:29:24	2021	2020	170.0N	S4092011.B24	
	11:54:16	11:29:26	3511	2022	180.0N	S4092011.B2E	
	12:17:44	11:54:18	4918	3512	170.0S	S4092011.B49	
	12:17:47	12:17:46	4920	4919	160.0N	S4092012.B13	
	12:49:40	12:13:11	7108	4921	150.0N	S4092012.B1D	
	13:15:13	12:17:49	8364	7109	140.0S	S4092012.B49	
274	13:31:59	13:15:15	9369	8365	130.0N	S4092013.B10	
851 km							21.09.2004
1 0409211.raw	12:14:36	12:14:27	10	1	130.0S	S4092112.B09	21.09.2004 S4092112.T09
0-700211.10W	12:55:35	12:14:27	2468	11	120.0N	S4092112.B03	5-1002112.109
	13:32:02	12:55:37	4654	2469	130.0S	S4092112.B50	
	13:53:17	13:32:04	5928	4655	140.0N	S4092113.B27	
301 km							

# 9. Marine Geology

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#### 9.1 Introduction

One of the key marine-geology objectives was the study of sediment dynamics of megaslides along the Svalbard continental margin (as part of the ESF Program "EUROMARGINS"), using Hydrosweep (bathymetry) and PARASOUND profiling data, multi-sensor-core-logging data, and long sediment cores. In this context, the characterization of sediment facies within debris flow and turbidite sequences as well as undisturbed pelagic sequences, the dating of sediment mass flows and estimates of sedimentray budgets are of interest. Furthermore, detailed multidisciplinary (i.e., sedimentological, mineralogical, micropaleontological, and geochemical) studies will be performed on the sediment cores for high-resolution reconstructions of paleoclimate, paleooceanic circulation patterns, paleosea-ice cover, and paleoproductivity and their variability during late Quaternary times. In order to reach these goals, an intensive sampling program was performed using large-volume box corer, multicorer, gravity corer, and kastenlot corer. In addition, gravity cores were taken from the central part of the Yermak Plateau as well as the East Greenland continental margin as part of the IODP Site survey work. All geological stations are shown in Fig. 9.1 (for details see station list, Chapter 12.1, Annex).

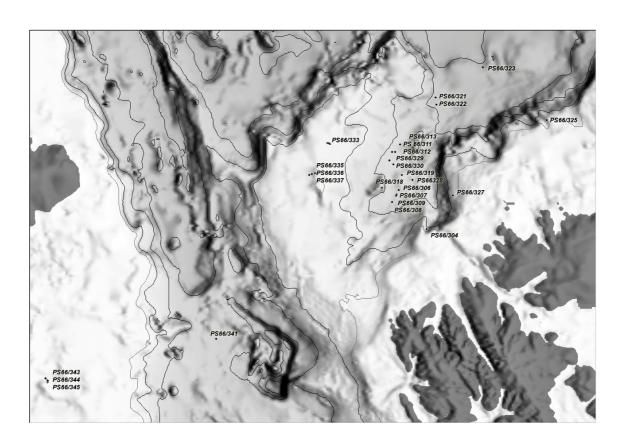


Fig. 9.1: Location map of geological stations of "Polarstern" Cruise ARK-XX/3

## 9.2. Marine Sediment Echosounding using PARASOUND

F. Niessen, J. Rogenhagen and F. Schoster

Scientific objectives, technical settings and operation conditions

Bottom and sub-bottom reflection patterns obtained by PARASOUND characterize the uppermost sediments in terms of their acoustic behavior. This can be used to interpret the sedimentary environments and their changes in space and time. The use and basic principles of the system are described in Spiess (1992). During ARK XX/3 the aims of PARASOUND profiling were (i) to select coring locations for gravity and box cores, (ii) to identify lateral differences in sedimentary facies with particular emphasis on submarine slides and debris flows north of Svalbard. (iii) to characterize the topmost sedimentary cover at locations proposed for IODP drilling as part of seismic site-survey work carried out during the cruise in the framework of various IODP-preproposals.

The hull-mounted PARASOUND sediment echosounder was in 24-hour operation along all cruise tracks starting from 2 September 2004, 15:35 UTC (77° 34.79'N, 05° 55.260'E) until after the completion of the seismic survey on 27 September 2004, 10:00 UTC (73° 35.336'N, 10° 35.663'W). The PARASOUND system (ATLAS Hydrographic, GmbH, Bremen, Germany) generates two primary frequencies between 18 and 23.5 kHz transmitting in a narrow beam of 4° at high power. As a result of the interaction of the primary frequencies within the water column, a secondary frequency is created based on the parametric effect. The parametric frequency is the difference frequency of the two primary waves transmitted. As a result of the longer wave length, the parametric frequency allowed sub bottom penetration up to 70m with a vertical resolution of ca. 30 cm. The parametric pulse length was set to 2 under normal operating conditions. Under extreme conditions, such as above steep slopes and while operating in heavy sea ice, the pulse length was increased up to 6. In water depth shallower than 2000m the mode of operation was PAR. At water depth larger than 2000m operation modes was PAR, NBS-PAR or PAR-pilot. In general, the operating conditions during the cruise were excellent due to mostly ice-free areas and relatively calm sea. Only in the most northerly areas the data acquisition was affected by moderate sea ice conditions causing noisy records with some traces missing.

## The new ATLAS PARASOUND-DS2 system onboard "Polarstern"

Before the start of the first cruise leg of ARK-XX, the "Polarstern" PARASOUND system received a major upgrade by installation of ATLAS PARASOUND-DS2. This equipment consists of the transmitter/receiver electronics and its control/data acquisition system. The former remained mostly unchanged as compared to the previous version, whereas the latter was completely replaced by an operator personal computer (PC) working as control, processing, display and recording unit. The function of the new system was checked during the Sea Acceptance Test (SAT) carried out during ARK-XX/1. The results are published in Budeus (in prep.). The basic elements of the entire system are summarized in a flow diagram (Fig. 9.2) and described in full details in the different reference manuals (ATLAS Hydrographic, 2004a, 2004b, 2004c). The first ATLAS PARA-

SOUND-DS2-system was installed onboard RV "Academic Boris Petrov" in 2003. The basic principles and first results using the system during cruise "Karas Sea 2002" are described in Dittmers and Schoster (2003).

The ATLAS PARASOUND-DS2 system comprises of the following essential components (Fig. 9.2):

- (1) Transducer arrangement SW 6051 G 003
- (2) Transceiver Cabinets GE 6019 A 003 and 004
- (3) Control Unit GE 6030
- (4) User interface in form of an Operator PC containing the control software ATLAS HYDROMAP SERVER, ATLAS HYDROMAP CONTROL and the data acquisition software ATLAS PARASTORE-3

The RV "Polarstern" PARASOUND-system is equipped with additional components:

- (5) Two colour printers HP Business Inkjet 1100 for printing of echograms and auxiliary data (navigation, depth and PARASOUND settings), respectively
- (6) Data Storage PC for data management, recording and data replay in offline mode
- (7) Spare PC of the Operator PC as hardware backup but fully installed in the winch control room echogram in order to provide track plot and single trace information for station work
- (8) Flat-Screen Monitor for duplication of echogram information of the Operator PC on the bridge
- (9) black and white chart recorder ATLAS Deso-25 as slave

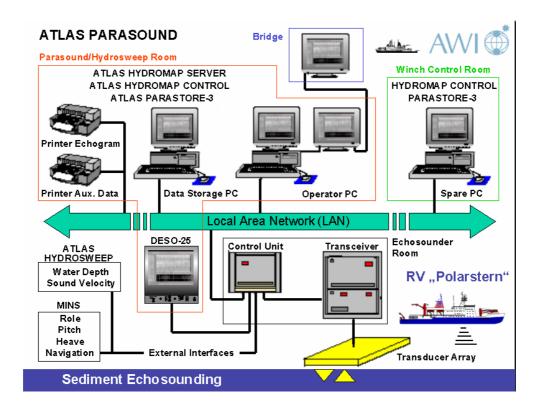


Fig. 9.2 Schematic overview of different hardware components of the new system PARA-SOUND DS-2 onboard RV "Polarstern"

The Data Storage PC and the Spare PC are equipped with the same software components (except for the ATLAS HYDROMAP SERVER installed on the Operator PC only) and linked to printers and Control Unit via a Local Area Network (Fig. 9.2). One major advantage of the new system is that several units of the former system were replaced by modern standard equipment for the benefit of easier handling, maintenance and repair. The former control unit (COLOR-SCOPE) has been replaced by a standard Windows-XP-based PC together with the new Control Unit GE 6030, which mostly contains of-the-shelf products. The equipment is now controlled by means of the ATLAS HYDROMAP CONTROL software via Operator PC (Fig. 9.2). The operating elements affect the system according to their functions (e.g. parametric frequency, pulse length, depth range, modes PAR/NBS-PAR/PAR-pilot, Narrow Beam Sounder (NBS) operation modes, beam steering, depth control and others).

The hardware and software configuration onboard "Polarstern" using two Windows-based PC's for system control and data management (Fig. 9.2), respectively, offers a lot of additional functions by means of the ATLAS PARASTORE-3 software via both Operator and Data Storage PC. ATLAS PARASTORE-3 provides a user-friendly graphical interface and has been designed to acquire, visualize, process, store, convert, quality control, replay and print data from AT-LAS PARASOUND-DS2 profiles. This includes not only the sub-bottom parametric data but also the primary narrow single-beam signal (NBS, 18 kHz). The recorded data of both parametric and NBS signals are stored in a hybrid raw data format (ASDF format) in either a ring buffer or into user selected folders on hard disc. The important improvement compared to the previous system is the acquisition of full soundings of both signals and not only data selected in specific depth windows by the operator. Optionally, the parametric data can be converted on-line into standard SEG-D format for further processing using industrial post-processing software. In addition, the parametric data can be stored in the formerly used PS-3 format (Paradigma). This function, however, requires careful watch keeping because only data kept in a selected depth window of commonly 200m length is stored. In addition to data acquisition, ATLAS PARASTORE-3 provides the following functions:

- (10) simultaneous operation of parametric signal and narrow beam signal
- (11) data storage reduction by user selecTable input of wait times (e.g. during station work)
- (12) colour-coded visualization of parametric and NBS data in arbitrary number of windows with different scales including single-trace visualization
- (13) appropriate processing algorithms such as bandpass filtering, clipping, negative flank subtraction and threshold filtering
- (14) colour-coded online print of echograms on user selecTable scales
- (15) Table print of navigation data and PARASOUND system settings in user-selecTable time intervals (auxiliary data)
- (16) information window for all system parameters combined with navigation information
- (17) visualization of ships track information on user selecTable map scales
- (18) storage of auxiliary data in user selecTable time intervals on hard disc

- (19) storage of meta information for each echogram window for monitoring operator activity during the survey on hard disc
- (20) network capability
- (21) replay of recorded data in off-line mode including visualization, processing, printing and storage in SED-D and/or PS3-formats

We have used the Operator PC for system control, data acquisition and storage as well as visualization of data in on-line mode. Both parametric and NBS signals were acquired. The stronger backscatter of the NBS-signal turned out to be very useful for keeping the echo in the selected time window when the sediment character or slope steepness did not allow sufficient reception of parametric reflections. The data storage PC was used for data management and replay of signals in off-line mode. Data transfer was carried out via Local Area Network (Fig. 9.2). This has allowed rapid visualization of previously stored PARA-SOUND data from ARK-XX/3 as well as PARASOUND data (PS-3 format) from previous cruises. In replay modes we have used software PARASTORE-3 and SENT (H. van Loom, University of Bremen). Unfortunately, the PARASOUND SEG-D data format was unreadable by shipboard seismic software "Focus".

### Data Management

The capabilities of the ATLAS PARASTORE-3 software in acquiring the PARA-SOUND data are two-fold: (i) simultaneously as raw data from the parametric channel and the NBS channel for full trace length including the water column, and, (ii) selected in specific depth windows, in PS3 and SEG-D formats. Because all the data can be produced and stored simultaneously, a well-organized data management is necessary. As outlined above the full soundings of the parametric and NBS signals are transferred from the Control Unit to the Operator PC and may be acquired there with the ATLAS PARASTORE-3 software and converted to PS3 and SEG-D format. So, it is possible to store parametric, NBS, PS3 and SEG-D data of the same profile. In addition, the newly implemented replay option of PARASTORE-3 gives the opportunity to replay the soundings of parametric data and/or NBS data and record them again as PS3 and/or SEG-D data, for example, for specifically defined profiles. In order to avoid confusion with the different formats and the status of the data, the data management was carried out by one designated person.

Taking into account the amount of hard disk space and/or storage media, the user has to decide what kind of data set is recorded and stored. The minimum data set would be the PS3 data of a specific depth window, mostly 200 m, and leads to a data amount of 30 to 60 MB per hour (dependent of the sounding rate) or 0.7 to 1.5 Gbyte per day. This could be further reduced with a functional resample option during storage, which is not yet implemented in PARASTORE-3. Keeping only PS3 data requires that the depth window is always manually adjusted to the topographic variations of the seafloor.

The other extreme is keeping all raw data and subsets of PS3 and SEG-D data, selected by a depth window of again 200 m. This data amount is ranging between 500 and 700 MB per hour (12 to 16 GByte per day), again dependent of the sounding rate. Hard disk capacity of the Data Storage PC is around 138

GByte, and would last for nine to ten days of PARASOUND measurement. Including the hard disk capacity of the Operator PC this time span can be doubled.

Both PC's are equipped with a DVD Writer and an LTO High Density Tape Drive, which allows the use of mass storage media. During cruise ARK XX/3, we decided to store the full data set of parametric and NBS raw data. NBS data has not been stored previously but the experience of this cruise showed that it also contains valuable information on the sediment structure especially in regions with steep slopes and irregular sea-floor topography such as measured at the surface of the Mega Slide. In addition, the PARASOUND system was monitored online all the time and the depth window adjusted, so that the out coming PS3 data were of good quality and stored as well. Because, according to testing, the converted SEG-D data from PARASTORE-3 could not be handled by other standard seismic software, we refrained from acquiring this data during the cruise.

The data storage was organized in such a way that the parametric raw data (.asd files) and the subsequent data set of PS3 data (.ps3) were recorded on DVD and LTO tapes, while the NBS data (.asd files) were recorded on LTO Tapes only. For backup reasons the LTO tapes were copied.

First Results with examples from the different working areas

There are six working areas, where PARASOUND profiling was of interest (from east to west, Fig. 9.3):

- (1) The IODP preproposal NSB-01A site survey grid north of North-Eastland (Svalbard)
- (2) The mega-slide area covering most of the Sophia Basin north of the Hinlopen Strait
- (3) A relatively small subbasin characterized by pelagic sediments between the Sophia Basin and the Yermak Plateau which also includes the sediment covered Mosby Seamount
- (4) The Yermak Plateau including its eastern and western slopes together with IODP preproposal site survey grids YP-03A and LE-01A
- (5) The central Fram Strait locations of the IODP preproposal FR-01A and FR-02A site survey grids
- (6) The East Greenland Shelf area of the IODP preproposal GR-01A-GR-05A site survey grid

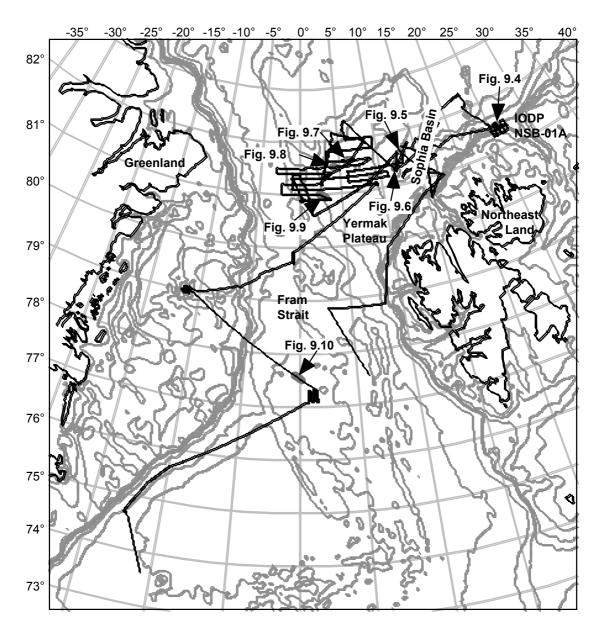
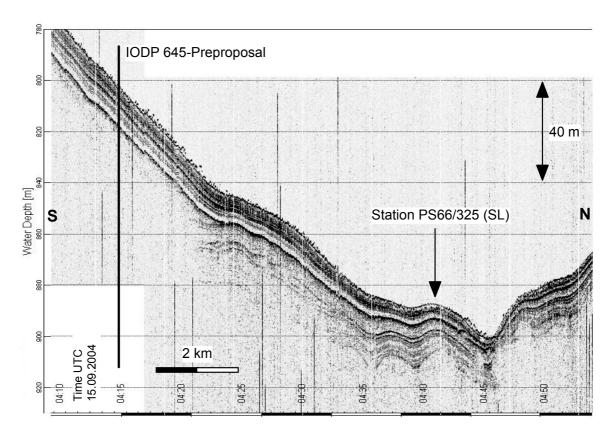


Fig. 9.3: Track map of RV "Polarstern" Cruise ARK-XX/3 during which PARASOUND was in operation and location of selected sections presented as examples in this report

In this report, examples from the different areas above are presented and described in brief including some preliminary interpretation. In particular, for areas (2), (3) and (4), the reader is referred to previous reports (Stein and Fahl, 1997; Jokat, 2000), in which PARASOUND profiles are presented from the same areas recorded during earlier "Polarstern" cruises. In order to avoid repetition, the examples described here give new information about the sedimentary environment and do not always represent the overall character of the different regions above. The locations of the PARASOUND profile examples presented are highlighted in the track lines along which PARASOUND was in operation during the entire cruise (Fig. 9.3).

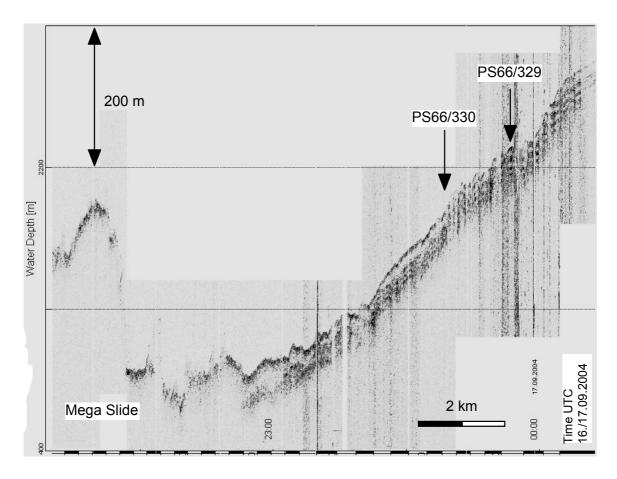
The example from area (1) including the location of coring station PS66/325, exhibits a typical situation adjacent to IODP coring location outlined in preproposal NSB-01A (Fig. 9.4). An upper undisturbed and continuous sediment unit

is defined by a strong basal reflector of maximum amplitude. In the profile presented (Fig. 9.4), there is a general S-to-N-decrease in thickness of this unit from the Svalbard side (about 10m) towards the Nansen Basin (about 2m). The underlying strata is well-stratified in the deeper area of the profile (>860m water depth), but may be invisible along the southern slope as a result of limited sound penetration. In places, lenticular-shaped debris flows, are intercalated at the boundary between the upper and lower unit (Fig. 9.4), and suggest a south to north flow direction.



**Fig. 9.4**: PARASOUND example from the area of IODP preproposal NSB-01 including coring location of PS66/325. Note that the dots plotted just above the sediment surface are an artefact caused by malfunction of the PARASOUND DS2-System (Bottom TVC)

The transition from area (2), the western end of the mega-Slide, to pelagic sediments west of the Sophia basin (3) exhibits hummocky relief with no or limited penetration. The slide is grading into a distal debris-flow tongue of transparent acoustic character overlying pelagic sediments before lensing out at the western end of the profile (Fig. 9.5). Although of complete appearance, well-stratified pelagic sediments built up a hummocky relief too, possibly as a result of post-sedimentary overprint by the adjacent slide. The latter may have caused overloading of pelagic sediments buried by the slide and subsequent humocky gravitational deformation and uplift west of the slide in areas largely uncovered by debris.



**Fig. 9.5**: PARASOUND example from the western end of the mega slide grading into pelagic sediments including coring locations PS66/329 and /330.

A chain of small seamounts are located in S-N-direction along the boundary between the Sophia Basin (2) and the small sub basin (3) towards the west. According to the PARASOUND data, one of them, the Mosby Seamount (Fig. 9.6) is covered by well-stratified sediments. Along the profile across the mount, the uppermost sediments decrease in thickness over a relatively small lateral distance from 30m (1350m water depth) to <5m (1250m water depth) on the top of the mount. The observed condensation of the sedimentary record is associated with increase in reflector amplitudes. This may suggest that currents have a stronger influence on sedimentation over the top of the mount compared to the flanks thereby decreasing sedimentation rates and possibly leading to the formation of lag-deposits with high acoustic reflectivity towards the top. The steeper N-E slope of the mount is marked by slump scars (Fig. 9.6 left).

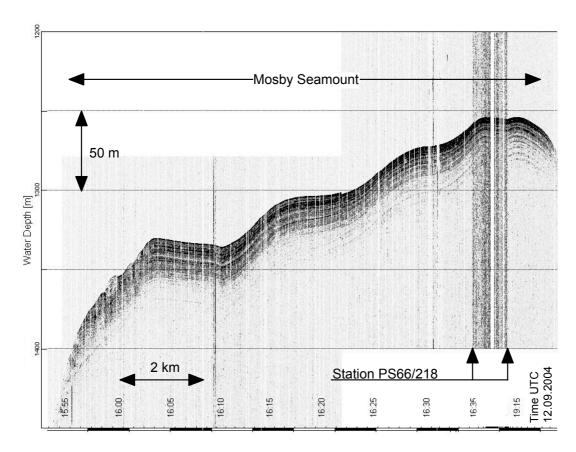
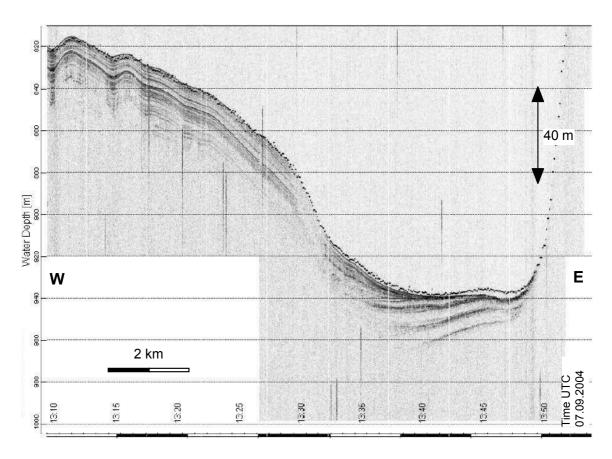
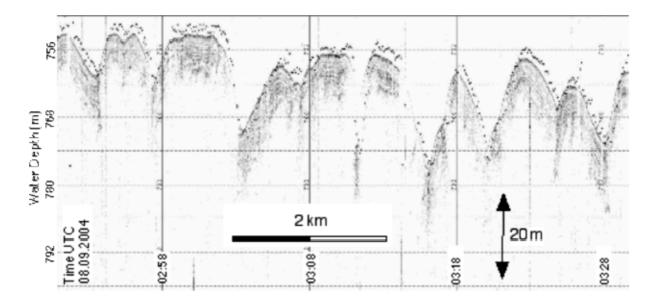


Fig. 9.6: PARASOUND example from the top area of the Mosby Seamount including coring location PS66/318

The topography of the Yermak Plateau (4) is marked by a number of small subplateaus which appear to be uplifted areas or relicts of basement reaching into relatively shallow water. In places, along the steep flanks of these sub-plateaus, PARASOUND indicates well-stratified sediments, which decrease in thickness significantly over short lateral distances down-slope from 60m in 820m water depth to less than 1m in 940m water depth directly adjacent to steep slopes (Fig. 9.7). In places, these deeper areas may also be marked by erosional discontinuities at the sediment surface. The overall character of the pattern is typical for marine drift deposits, which are controlled by the influence of bottom currents. On the shallowest areas of the Yermak Plateau (Fig. 9.8) the relief is characterized by deep furrows. The fact that PARASOUND exhibits a pelagic cover of well-stratified sediments overlying furrows and hummocks suggests a paleo-relief possibly formed by massive ice-berg gouging during earlier glaciations with intensive ice-berg drift (e.g. Vogt et al., 1994).



**Fig. 9.7**: PARASOUND example from the eastern part of the Yermak Plateau indicating drift deposits adjacent to steep slopes of elevated basement. Note that the dots plotted just above the sediment surface are an artefact caused by malfunction of the PARASOUND DS2-System (Bottom TVC).



**Fig. 9.8**:PARASOUND example from the top of the Yermak Plateau showing sediment-covered furrows related to ice erosion. Note that the dots plotted just above the sediment surface are an artefact caused by malfunction of the PARASOUND DS2-System (Bottom TVC).

The western slope of the Yermak Plateau (4) towards the Fram Strait (5) is often characterized by well-stratified pelagic sediments intercalated with up to 20m thick transparent debris-flow deposits (Fig. 9.9). These debris flows are rarely found directly at the sediment surface but are mostly covered by packages of undisturbed sediments of 5 to 10m in thickness including several distinct sub-parallel reflectors (Fig. 9.9).

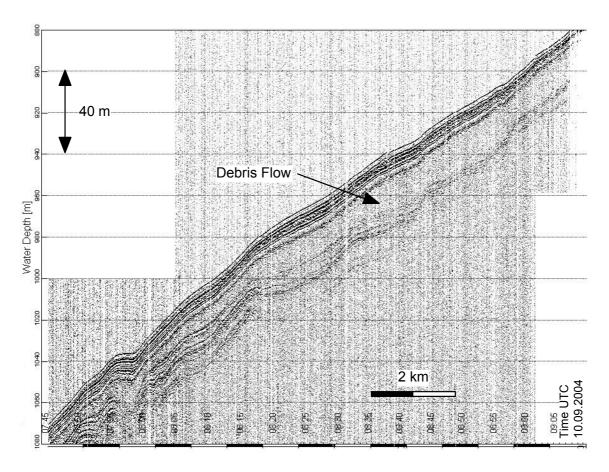
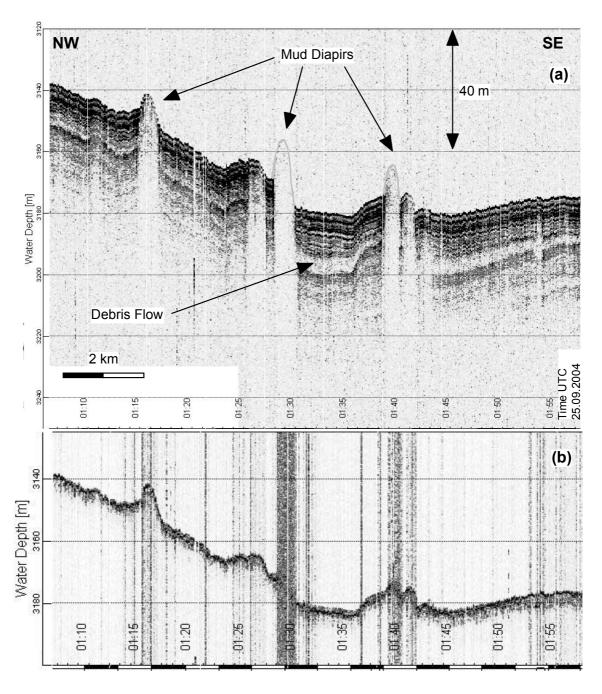


Fig. 9.9: PARASOUND example from the western slope of the Yermak Plateau showing typical intercalation of pelagic sediments with debris flows

The presence of debris-flow deposits is also common off the slope of the East Greenland Shelf and in the area of the Southern Fram Strait (6). The sequence of pelagic sediments intercalated with debris flows is in places laterally interrupted by diapirs visible as acoustic voids but forming distinct hummocks on the sediment surface up to 20m high (Fig. 9.10). Possibly the source of the diapirs are muds of high porosity/low density derived from debris-flow deposits in greater depth (not visible in Fig. 9.10 as a result of limited sound penetration). In places, these muds were uplifted in diapirs due to overburden of denser pelagic sediments. The general pattern is typical for marine slopes and basins off formerly glaciated continental margins. In the presented example the primary 18 kHz NBS-record is compared to the secondary parametric 4 kHz PARASOUND signal. Note the differences in resolution and penetration (Fig. 9.10).



**Fig. 9.10**: PARASOUND example from the southern Fram Strait indicative of mud diapirs cutting through pelagic sediments and debris flows

## 9.3 Physical Properties and Core Logging

F. Niessen, J. Rogenhagen, C. Gebhardt, D. Penshorn, J. Schneider, D. Winkelmann

### 9.3.1 Multi-sensor Core Logging

F. Niessen, J. Rogenhagen, C. Gebhardt, J. Schneider, D. Penshorn

The physical properties of whole sediment cores (gravity cores and box cores) were measured in one-meter sections using a GEOTEK ,Multi-Sensor-Core-Logger' (MSCL). P-wave travel times, magnetic susceptibility, and  $\gamma$ -ray-absorption were measured simultaneously including control measurements of core diameter and sediment temperature. From these data, the physical properties density, magnetic susceptibility, fractional porosity, P-wave velocity, and impedance can be calculated using the MSCL software package. The technical description of the system is given in Table 9.1.

Table 9.1: Technical specifications of the GEOTEK MSCL14

## P-wave velocity and core diameter

Plate-transducer diameter: 4 cm Transmitter pulse frequency: 500 kHz

Pulse repetition rate: 1 kHz Received pulse resolution: 50 ns

Gate: 5000 Delay: 0 □s

#### **Density**

Gamma ray source: Cs-137 (1983)

Activity: 356 MBq Energy: 0.662 MeV Collimator diameter: 5.0 mm

Gamma detector: Gammasearch2, Model SD302D, Ser. Nr. 3043, John Count Scientific

Ltd.,

10 s counting time

#### Fractional porosity

Mineral grain density = 2.75, water density = 1.026

### **Magnetic susceptibility**

Loop sensor: BARTINGTON MS-2C, Ser. Nr. 208

Loop sensor diameter: 14 cm

Point sensor: BARTINGTON MS-2F, Ser. Nr. 139

Alternating field frequency: 565 Hz, counting time 10 s, precision 0.1 \* 10<sup>-5</sup> (SI)

Magnetic field intensity: ca. 80 A/m RMS

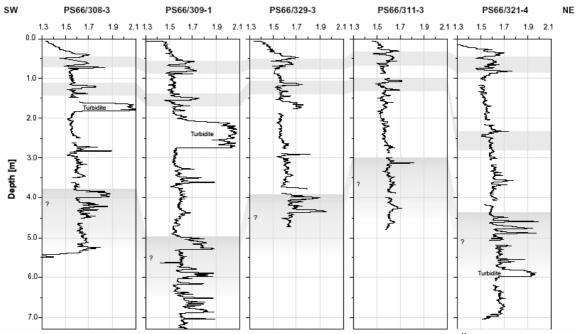
Krel: 1.56 (SL, 12 cm core- $\emptyset$ ), 0.60 (KAL, cross section 50.27 cm<sup>2</sup>, = 8 cm core- $\emptyset$ ) Loop sensor correction coefficient: 6.391 (SL) for 10^-6 (SI), 16.689 (KAL) for 10^-6 (SI)

Point sensor: BARTINGTON MS-2F, Ser. No. 139

counting time 10 s

In addition, measurements of magnetic susceptibility using a Bartington point sensor were performed on split cores. For practical reasons, the point sensor was hooked up on the GEOTEK Colour-Line-Scan-Logger. Data acquisition was carried out in combination with colour scans of split-cores surfaces. The latter is described in more detail under 9.3.2. of this report. The step intervals of the measurements were 0.5 or 1.0 cm (see appendix for more details). Gravity cores were measured in coring liners, whereas Kastenlot cores were measured in sub-cores retrieved from the original core using length-wise open transparent plastic boxes of 1000mm length and 75 x 75 mm cross section (outside includ-

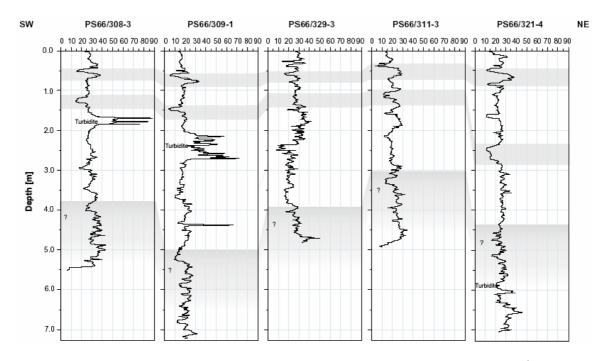
ing 2 x 2.5 mm plastic wall thickness). In order to allow for magnetic-susceptibility sensor correction (according to the Bartington correction requirements), the rectangular cross section of the box was equalised to an size-equivalent circular section, of which a fictive core diameter was calculated as input parameter for loop-sensor correction coefficient (Table 9.1). For both gravity and Kastenlot cores, the density calibration was carried out using a set of defined mixtures of aluminium and distilled water in a gravity liner and Kastenlot sub-sampling box, respectively (Best and Gunn, 1999).



**Fig. 9.11**: Lateral core-correlation using wet-bulk density (g/cm<sup>3</sup>).

Physical properties are commonly used for lateral core correlation. For the area of the Fram Strait and Yermak Plateau, wet bulk density proofed to function as useful correlation parameter for gravity and Kastenlot cores obtained during a previous "Polarstern" expedition (ARK-XIII/2; Stein and Fahl, 1997). For the area under investigation during this cruise wet-bulk density again exhibits distinct down-core pattern of variability, which is found in several cores obtained from pelagic facies. For the pelagic sediments west of the Sophia Basin, largely covered by the mega-slide, a wet-bulk density-correlation for a cross section from core PS66/308-3 (south-west) to core PS66/321-4 (north-east) is presented as example (Fig. 9.11). Some distinct density wiggles can be correlated along the entire profile, whereas peaks with maximum densities are observed in turbidites of cores PS66/308-3 and PS66/309-1 only (Fig. 9.11). This local occurrence of turbidites is related to distal deposition caused by the mega-slide. Thus, the physical property records reveal an ideal tool to put the slide event into a first stratigraphic framework. Along the same profile, an overlay of the density correlation onto magnetic susceptibility data indicates that some density wiggles coincide with depth intervals, in which susceptibility exhibits distinct fluctuations too (Fig. 9.12). The latter, however, do not occur in exactly the same core-depth levels suggesting that the fluctuation of magnetic susceptibility is not purely a function of variation in density/porosity but has its own signature.

Also, it appears that the correlation using magnetic susceptibility is less clear than that of density, and, for some depth intervals weak or even invisible (Fig. 9.12). Again, very high magnetic susceptibility is mostly measured in turbidites (Fig. 9.12). The down-core pattern of all physical properties measured during the cruise is presented in the Appendix of this report.



**Fig. 9.12**: Lateral core-correlation of magnetic susceptibility (point-sensor data in 10<sup>-6</sup> SI; sensor susceptibility not converted to sediment volume). Note that correlation of grey-shaded areas is an overlay obtained by density data (Fig. 9.11) in order to test the core correlation for consistency.

## 9.3.2 Colour Scan Logging

A. C. Gebhardt, F. Niessen, D. Penshorn, J. Schneider

In combination with the magnetic susceptibility point sensor measurements (see Chapter 9.3.1), all split-core surfaces were colour-scanned with a resolution of 10 to 20 scan lines per millimetre in vertical and 1000 pixels in horizontal direction to have a digital downcore colour signal of the fresh sediment. The GEO-SCAN camera was installed on an MSCL array and calibrated as described in the GEOTEK Multi-Sensor Core Logger Manual (2000). The camera uses a diachronic colour interference filter and three line-scan device arrays (1024 pixels each) to continuously record the three red-green-blue (RGB) colour channels. Its aperture was set to maximize contrast for the lightest-coloured sediment of each core. Synchronisation between the camera and the track is achieved by using the stepper motor pulses to trigger the line acquisition of the camera. Camera data is demultiplexed in the interface box, and 24-bit data is transmitted differentially to the PC interface card. In the PC data are corrected for convergence, gain and offset with the previously set calibration parameters. Profiles for each RGB channel were produced by averaging pixels in 6 cm x 1 cm and 8 cm x 1 cm rectangles for Kastenlot and gravity cores, respectively. along the central axis of the core using GEOTEK MSCL imaging software.

As an example, results from colour-scan logging of Kastenlot Core PS66/309-1 are shown in Fig. 9.13. Further records are presented in Chapter 12.5 (Annex).

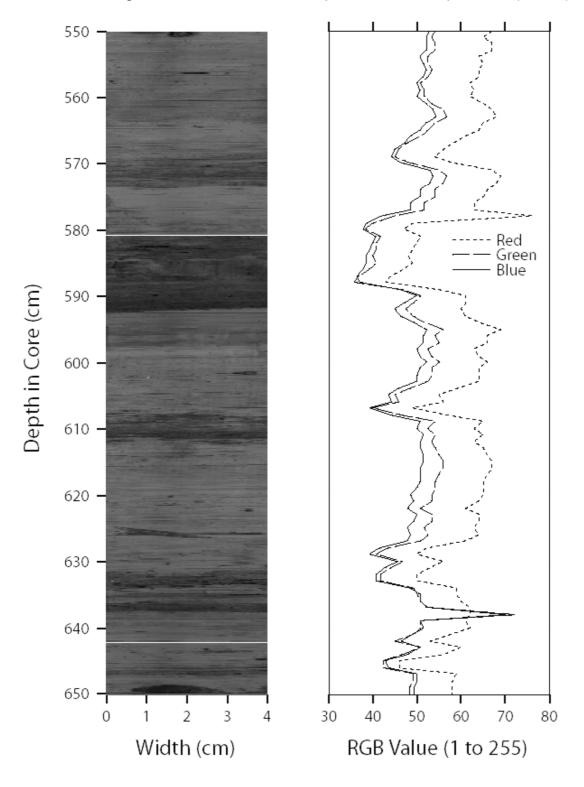


Fig. 9.13: Gray-value image of Core PS66/309-1, depth 550 to 650 cm, and RGB values of this section.

## 9.3.3 Shear Strength

D. Winkelmann, D. Penshorn

Shear strength measurements of undrained marine sediment (commonly silty clays) have been carried out by using both the fall-cone device Geotechnik g-200 and for comparison the Haake Rotovisco RV-12/500M on material from selected gravity kastenlot cores.

The fall cone method originally developed by John Olsson, member of the technical commission of the Swedish Royal Railway, in 1915, was used to determine the consistency of clays. The penetration of the fall cone into the clay or soil is considered to be directly proportional to its weight. According to Hansbo (1957) the relation between undrained shear strength s and the penetration h of a cone of the weight m can be written as:

$$s = K * m * h^{-2}$$

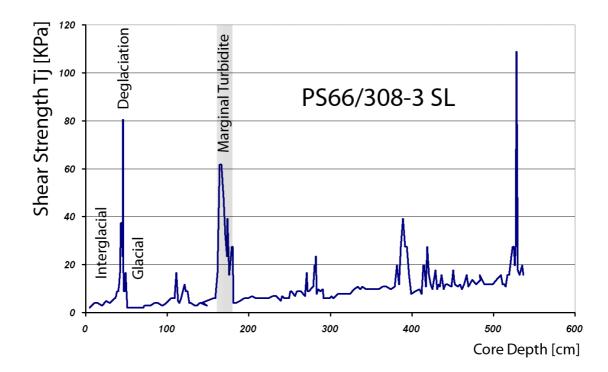
The constant K represents the angle of the cone and the sensity of the clay. For comparison to the SI system gravity has to be considered as:

$$Tj = s * g$$

The rotovisco method uses a winged wheel pressed into the sediment followed by rotational momentum applied to it until the wheel turns. The relation of the rotational momentum and shear strength Tj can be written as:

$$Tj = f * a$$

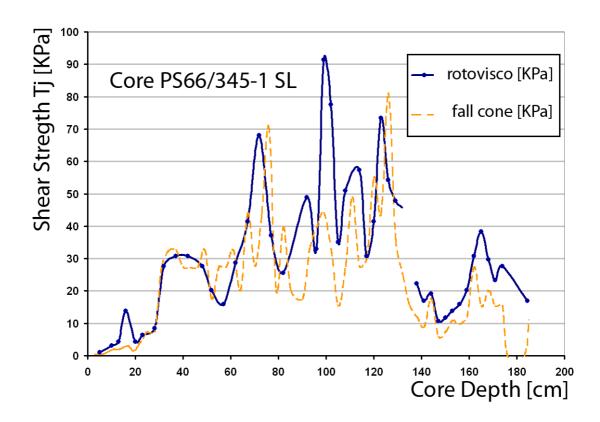
The shape factor f represents a proportionality constant [f =  $1/(2*\pi*h*R^2)$ ] while a is the rotational momentum per scale unit.



**Fig. 9.14**: Shear strength plot of core PS66/308-3 SL from the northern Svalbard continental margin close to the mega-slide margin.

Variability of measured shear strength can be attributed to major lithological changes within the sediment core mirroring glacial and interglacial conditions or turbidites from the adjacent submarine slide (Fig. 9.14). In some sediment cores from the Greenland shelf extraordinary shear strength characterises overconsolidated diamictons from an extensive glaciation (Fig. 9.15).

Problems of the method arise from sandy layers within the sediment cores. Due to dewatering effects the fall cone method overestimates the shear strength of such layers. Secondly, the application of both methods to the same core material leaves the second method with a slightly altered sediment resulting in lower or higher shear strength values.



**Fig. 9.15**: Shear strength record from Core PS66/345-1, East Greenland continental margin. Comparison of parameters from the two methods applied for determination of the shear strength.

## 9.4 Geological Sampling, Description, and Methods Applied

F. Schoster, B. Bahr, C. Gebhardt, N. Kukina, N. Lensch, S. Nam, H. Noffke, D. Penshorn, A. Puhr, R. Saraswat, Chr. Schäfer, J. Schneider, R. Stein, J. Thiele, D. Winkelmann, Y. Yanina)

Surface and near-surface sediment sampling

Surface and near-surface sediment sampling was carried out by using a giant box corer and a multicorer.

The Giant Box Corer (weight of ca. 500 kg; volume of sample 50\*50\*60 cm; manufactured by Fa. Wuttke, Henstedt-Ulzburg, Germany) was successfully used 14 times at 16 stations. Two times there was no recovery due to technical problems. From the box corer surface sediments and three sediment cores (diameter 12 cm) for (1) sedimentology, mineralogy, geochemistry, (2) heavy minerals, and (3) archiving were taken. The second core was logged (MSCL, colour, magnetic susceptibility; see Chapter 9.3), opened and sampled for heavy mineral determination (MMBI) and bulk parameters like total organic carbon, total carbon, carbonate, and Rock-Eval analysis (AWI). The following samples were obtained from the surface sediments:

```
10x10 cm<sup>2</sup>
               (100cc)
                              Foraminifera (IfM-GEOMAR)
10x10 cm<sup>2</sup>
                              Org. Geochemistry (AWI)
               (100cc)
10x10 cm<sup>2</sup>
               (100cc)
                              Sedimentology (AWI)
10x10 cm<sup>2</sup>
               (100cc)
                              Micropaleontology (KIGAM)
3x3 cm<sup>2</sup>
               (ca. 10 cc; freeze-drying onboard) TOC (AWI)
10x10 cm<sup>2</sup>
                              Sedimentology (IORAS)
               (100cc)
```

The standard 8-tubes-version multicorer (manufactured by Fa. Wuttke, Henstedt-Ulzburg, Germany) with an inner tube diameter of 10 cm was used. The penetration weight was always 250 kg. The multicorer was successfully used 6 times at 7 stations, and usually recovered undisturbed surface sediments and overlying bottom water. One multicorer was empty due to a technical problem. In general, the multicorer cores were sampled for the following investigations in slices of 1cm throughout the whole core:

```
    3 cores Org. Geochemistry (IfBM Hamburg)
    2 cores Micropaleontology (AWI)
    2 cores Org. Geochemistry-Sedimentology (AWI)
    1 Core (Only surface 0-1 cm) TOC, freeze-drying onboard (AWI)
```

### Sampling of long sediment cores

Long sediment cores were taken by a gravity corer and a kastenlot. The gravity corer (GC or "Schwerelot", SL) has a penetration weight of 1.5 t. It was successfully used in variable barrel lengths of 3, 5, or 10 m at 26 stations (33 cores; see Table 9.1 for details). The recovery of the gravity corer varied between 0.1 and 7.11 m, the penetration between 0.1 and 11 m (Table 9.2).

The kastenlot (Kögler, 1963), a gravity corer with a rectangular cross section of 30x30 cm, has a penetration weight of 3.5 t and a corebox segment sized 30x30x575 cm (manufactured by Hydrowerkstätten Kiel). The length of the kastenlot boxes used was 11.75 m plus about 30 cm for the core catcher. The great advantage of this kastenlot is the wall-thickness of only 0.2 cm. Because of the great cross-sectional area (900 cm2) and the small thickness of the walls, the quality of the cores was generally excellent. The kastenlot was successfully used at 3 stations. The recovery of the kastenlot cores varied between 2.80 and 7.27 m (Table 9.2, Fig. 9.16).

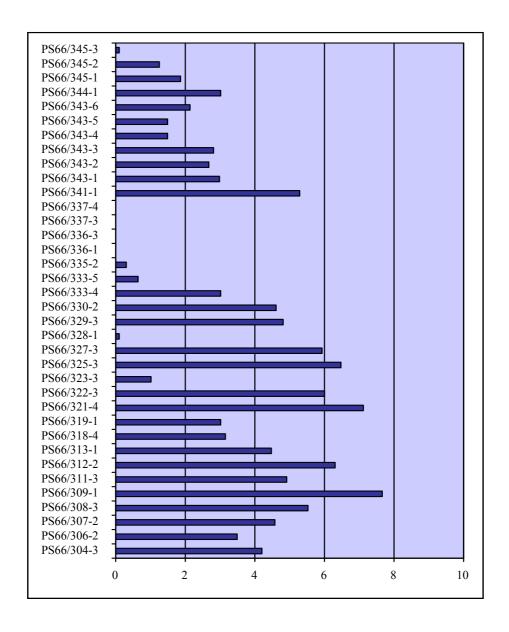


Fig. 9.16: Recovery of the gravity and kastenlot cores (in m). For location of cores see Fig. 9.1.

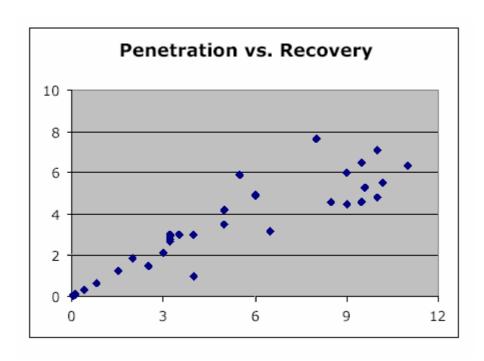


Fig. 9.17: Penetration (m) vs. recovery (m) of the gravity and kastenlot cores.

Fig. 9.17 shows the relation between recovery and penetration of the gravity and kastenlot cores.

All gravity cores and kastenlot cores were logged (MSCL, colour, magnetic susceptibility; Chapter 9.3). Most of them were opened, described (see Chapter 9.6.4. and Annex), and sampled for heavy minerals (MMBI) and bulk parameters like total organic carbon, total carbon, carbonate, and Rock-Eval parameters (AWI). Core PS66/309-1 KAL was also sampled for coarse fraction and foraminifera determination (see Chapter 9.6.3.). From all opened cores slices were taken in order to do X-ray photographs (see Chapter 9.6.1.).

**Table 9.2:** Gravity and kastenlot cores from ARK XX/3 with penetration and recovery values. Cores opened during the expedition, are marked with an "x".

					Depth			_	
Station	Date	Time	Latitude	Longitude	[m]	Gear	Penetration	Recovery	Opened
PS66/304-3	03.09.04	17:40	80°43,20'N	14°38,02' E	1095,0	GC-5	5	4,19	Х
PS66/306-2	04.09.04	03:34	81°14,56' N	13°18,62' E	2268,0	GC-5	5	3,5	Х
PS66/307-2	04.09.04	06:21	81°11,94' N	13°2,68' E	2275,0	GC-10	8,5	4,57	Х
PS66/308-3	04.09.04	10:47	81°7,30' N	12°35,97' E	2218,0	GC-10	10,2	5,54	Х
PS66/309-1	04.09.04	13:44	81°11,22' N	12°59,08' E	2269,0	KAL	8	7,65	Х
PS66/311-3	05.09.04	02:53	81°41,78' N	13°28,20' E	2192,0	GC-5	6	4,93	Х
PS66/312-2	05.09.04	05:30	81°41,24' N	13°42,65' E	2275,0	GC-10	11	6,32	Х
PS66/313-1	05.09.04	07:35	81°45,56' N	14°16,31' E	2298,0	GC-10	9	4,49	Х
PS66/318-4	12.09.04	18:46	81°18,60' N	12°2,71' E	1262,0	KAL	6,5	3,15	Х
PS66/319-1	12.09.04	22:26	81°24,49' N	13°48,79' E	2178,0	GC-3	4	3	Х
PS66/321-4	13.09.04	11:51	82°10,27' N	18°13,47' E	2360,0	GC-10	10	7,11	Х
PS66/322-3	13.09.04	16:16	82°5,40' N	18°5,25' E	3028,0	GC-10	9	6	х
PS66/323-3	14.09.04	06:06	82°18,44' N	22°59,67' E	3913,0	GC-3	4	1	Х
PS66/325-3	15.09.04	16:32	81°26,18' N	25°59,86' E	895,5	GC-10	9,5	6,48	х
PS66/327-3	16.09.04	17:18	81°0,92' N	17°9,39' E	701,0	KAL	5,5	5,92	х
PS66/328-1	16.09.04	21:03	81°19,25' N	14°30,88' E	1979,0	GC-5	0,1	0,1	х
PS66/329-3	17.09.04	03:44	81°36,31' N	13°6,13' E	2211,0	GC-10	10	4,82	х
PS66/330-2	17.09.04	06:44	81°32,88' N	13°21,71' E	2305,0	GC-10	9,5	4,6	х
PS66/333-4	19.09.04	22:30	81°56,42' N	8°30,17' E	725,5	GC-3	3,5	3	х
PS66/333-5	19.09.04	23:26	81°56,10' N	8°31,50' E	719,1	GC-3	0,8	0,63	х
PS66/335-2	20.09.04	06:47	81°36,05' N	6°34,06' E	654,8	GC-3	0,4	0,32	х
PS66/336-1	20.09.04	07:40	81°36,48' N	6°45,14' E	505,2	GC-3	0,05	0,05	х
PS66/336-3	20.09.04	08:22	81°36,41' N	6°47,40' E	505,9	GC-3	0,05	0,05	х
PS66/337-3	20.09.04	10:29	81°36,30' N	7°1,44' E	503,0	GC-3	0,05	0,05	х
PS66/337-4	20.09.04	11:12	81°36,61' N	7°3,18' E	587,8	GC-3	0,05	0,05	х
PS66/341-1	22.09.04	23:43	79°44,07' N	0°44,97' W	2822,0	GC-10	9,6	5,28	х
PS66/343-1	24.09.04	03:50	79°0,83' N	11°6,91' W	216,2	GC-3	3,2	2,97	
PS66/343-2	24.09.04	04:20	79°0,98' N	11°6,82' W	226,0	GC-3	3,2	2,67	
PS66/343-3	24.09.04	04:49	79°1,05' N	11°6,88' W	227,2	GC-3	3,2	2,8	
PS66/343-4	24.09.04	05:18	79°1,17' N	11°7,12' W	240,9	GC-3	2,5	1,49	
PS66/343-5	24.09.04	05:44	79°1,31' N	11°6,80' W	252,2	GC-3	2,5	1,49	
PS66/343-6	24.09.04	06:09	79°1,44' N	11°6,84' W	264,2	GC-3	3	2,14	
PS66/344-1	24.09.04	06:48	79°2,55' N	11°6,65' W	338,7	GC-3	3,2	3,02	
PS66/345-1	24.09.04	07:30	79°3,58' N	11°15,11'W	317,9	GC-3	2	1,86	х
PS66/345-2	24.09.04	08:08	79°3,58' N	11°16,03' W	313,3	GC-3	1,5	1,27	x
PS66/345-3	24.09.04	08:40	79°3,58' N	11°17,37' W	308,6	GC-3	0,1	0,1	x

## Sampling of crystalline basement

In order to collect samples from outcropping crystalline basement a dredge (chain bag) was successfully used two times on three stations (Table 9.3). At one station the overlying sediment layer was too thick. The chain bag was  $0.95 \times 0.35 \times 0.40$  m large and had a weight of 100 kg. The rope length during dredging was three times the water depth.

At station PS66/333-2 and PS66/335-1 stones were collected. The stones varied in size and composition, some of them were covered by a black layer (Fe-Mn-crust?). Identification and quantification will take place in the home laboratories.

Table 9.3. Positions of the dredges during ARK XX/3.

					Waterdepth		
Station	Date	Time	Latitude	Longitude	[m]	Gear	Action
PS66/333-2	19.09.04	17:54	81° 55.54' N	8° 34.29' E	751.5	DRG_C	Dredge Beginn
PS66/333-2	19.09.04	19:05	81° 56.32' N	8° 31.62' E	720.8	DRG_C	Dredge Ende
PS66/333-3	19.09.04	20:45	81° 56.53' N	8° 22.74' E	718.5	DRG_C	Dredge Beginn
PS66/333-3	19.09.04	21:34	81° 56.49' N	8° 22.77' E	716.8	DRG_C	Dredge Ende
PS66/335-1	20.09.04	05:08	81° 35.99' N	6° 32.00' E	863.7	DRG_C	Dredge Beginn
PS66/335-1	20.09.04	05:29	81° 36.02' N	6° 35.42' E	558.5	DRG_C	Dredge Ende

#### 9.5 Characteristics of Surface Sediments

S. Nam, N. Kukina, R. Saraswat, J. Thiele, Y. Yanina

During ARK-XX/3 expedition, at a total of 14 geological stations the giant box corer (GKG) was deployed to take undisturbed surface sediments from water depths varying between 503 and 3910 m. Recovery of GKG cores ranges between 16 and 55 cm in length. A total of 11 box cores were mostly undisturbed, while the surface sediments of two cores (PS66/306-1GKG and PS66/323-1GKG) were strongly and/or probably disturbed, respectively (Table 9.4). In general, the mostly undisturbed surface sediments mainly consist of brown to olive brown silty clay to mud. Core PS66/336-4 comprised many black-coloured (black coating?) dropstones overlying nearly undisturbed surface sediments. Core PS66/337-2GKG, retrieved from the shelf of the Yermak Plateau, mainly contained high amounts of dropstones with strongly varying sizes between gravel and boulder, and a little amount of coarse sand. The majority of dropstones of gravel to boulder sizes seem to be also black-coloured (basalts? and/or black-coated sedimentary rocks?). Table 9.4 shows general information about all box core sediments collected during this expedition.

**Table 9.4**: General overview of box cores collected during ARK-XX/3.

Core (GKG)	Core length (cm)	Water depth (m)	S	Sampling No.		Color & lithology	Remarks
		,	100cm <sup>3</sup>	biomarker	tube		
PS66/306-1	51	2271	1	Х	2	brown silty clay	abundant benthic forams ( <i>Pyro</i> )
PS66/307-1	48	2276	6	1	3	brown silty clay	abundant benthic forams ( <i>Pyro</i> )
PS66/308-1	53	2217	6	1	3	brown silty clay	moderate benthic forams ( <i>Pyro</i> ), worms
PS66/311-1	-	2195	6	1	3	brown silty clay	traces of benthic organisms, not described
PS66/312-1	51	2273	6	1	3	brown silty clay	worms, abundant benthic forams ( <i>Pyro</i> ), traces of benthic organisms
PS66/321-2	25	2359	6	1	3	brown silty clay	traces of benthic organisms, sponge
PS66/322-2	23	2992	6	1	3	brown silty clay	traces of benthic organisms, shell fragments, benthic forams
PS66/323-1	24	3910	Х	X	2	brown silty clay	probably surface disturbed
PS66/325-2	53	896	6	1	3	olive gray to brown silty clay	abundant benthic organisms and its traces, sea-stars of <~1 cm, worms, bivalve shell (~2 cm) with flat form
PS66/327-2	55	701	6	1	3	olive gray to brown silty clay	abundant benthic organisms, i.e. sea-stars, nema- todes and/or an- nelids, bird feather
PS66/329-1	52	2209	6	1	3	brown mud	various benthic organisms includ- ing sea-stars, nematodes and/or annelids, benthic forams
PS66/330-1	50	2306	6	1	3	brown mud	dropstone of ca. 25 cm long
PS66/336-4	16	505	X	X	3	gravelly sand mud	abundant drop- stones of gravel to boulder, abundant benthic organisms (forams, brittle stars, bivalves, worms, etc.)
PS66/337-2	51	2271	Х	Х	Х	gravel	abundant drop- stones of gravel to boulder

Prior to opening the box cores onboard, wet bulk density and P-wave velocity and magnetic susceptibility (10-5 SI) were determined in 1 cm intervals using Multi Sensor Core Logger (MSCL-14) on all box cores. The box cores were then splited in two halves one each for archive and working. The working half was used for preliminary description. Both half cores were photographed. The colors and magnetic susceptibility (10-5 SI) of the archive cores were additionally logged in 0.5 cm intervals using Color Scanner. Sediment colors were visually noted using Munsell Soil Color Charts (1954). Lithological changes and primary structures of sediment were also described with an exception of core PS66/311-1GKG. X-radiographs were made for the determination of sedimentary structures and to estimate the contents of IRD. Generally, the sediments of box cores are composed mainly of brown to dark grayish brown homogenous mud and silty clay. However, slightly bioturbated and/or laiminated structures are also observed in several cores (see Chapter 12.3, Annex).

Table 9.5: Bulk minerals of surface (0-1cm) sediments

Stations	Quartz	Feld- spar	Mica	Terr. carbo nate	Vol- canic glass	Clay mine- rals	Org. re- mains	Biog. Carbo- nate	Fe- hydro- xide	Other Heavy
PS66/304- 2MUC	30,2	18,1	7,3	4,1	1,6	5,3	8,8	2,1	10,2	12,3
PS66/307- 1GKG	28,2	20,6	7,8	0	2,1	7,3	8,9	1,1	4,1	19,9
PS66/308- 1GKG	30,2	10,1	6,2	1,1	2,1	15,4	20,1	4,9	2,1	7,8
PS66/311- 1GKG	20,8	11,2	5,3	1,1	1,5	11,1	12,2	2,2	25,5	9,1
PS66/312- 1GKG	26,3	10,1	5,3	1,2	0	21,2	11,3	4,6	3,2	16,8
PS66/321- 2GKG	33,5	14,2	8,6	0	0	10,2	8,5	3,6	6,6	14,8
PS66/322- 2GKG	22,1	8,2	7,2	0	0	25	11,2	1,1	12	13,2
PS66/325- 2GKG	22,6	11,1	8,8	0	0	26,6	9,1	0	14,1	7,7
PS66/327- 2GKG	19	6,4	4,3	6,4	2,2	17,2	18,3	2,2	7,5	16,5
PS66/329- 1GKG	28,1	14,2	6,3	1	0	20,3	18	0,5	4,6	7
PS66/330- 1GKG	20,3	8,9	9,3	0	0	5,6	25	8	6	16,9
PS66/336- 4GKG	30,1	10,8	7,1	0	0	18,3	5,9	2,5	6,8	18,5
PS66/337- 2GKG	25,3	11,2	8,9	1,2	0	6,9	9,9	12,2	8,1	16,3

Table 9.6: Heavy minerals of surface (0-1cm) sediments

Stations	Pyroxene	Amphibole	<b>Epidote</b>	Clorite	Garnet	Black ores	Opaque
PS66/304-2MUC	38,1	31	9,2	2,6	5,8	6,2	7,1
PS66/307-1GKG	38,7	25,1	12,6	4,4	6,9	7,2	5,1
PS66/308-1GKG	22,6	41,1	5,6	0	14,2	11,3	5,2
PS66/311-1GKG	21,5	39,9	2,8	0	8,9	16,3	10,6
PS66/312-1GKG	40	39	6,3	0	5,2	3,2	6,3
PS66/321-2GKG	38,4	29,6	7,9	0	3,3	10,6	10,2
PS66/322-2GKG	32,2	30	8,9	2,2	8,2	15,3	3,2
PS66/325-2GKG	40,1	30,1	11,2	2,6	5,5	7,8	2,7
PS66/327-2GKG	23,8	28,6	9,5	0	4,8	19	14,3
PS66/329-1GKG	30	35,2	4,8	0	12,2	14,6	3,2
PS66/330-1GKG	39	28,3	4,6	2,1	2,5	12,3	11,2
PS66/336-4GKG	40	30	5	1,8	2,1	14,2	6,9
PS66/337-2GKG	32,7	21,2	5,8	3,8	9,6	9,6	17,3

Surface (0-1cm) sediments from 13 box cores collected during the ARK-XX/3 cruise were investigated with smear slides. Results of smear slide analysis are given in Table 9.5 and 9.6. Smear slide analysis shows that major surface sediments are comprised of quartz and clay minerals with minor amounts of feld-spars, heavy minerals and Fe-oxides/hydroxides. The abundance of heavy minerals varies from 11.0% (PS66/329-1) to 34.6% (PS66/311-1). The increase in the amount of heavy minerals is on account of Fe-oxides/hydroxides. The biogenic components and diatoms are almost absent in all the samples.

Quartz, feldspar and mica constitute the majority of the light minerals. In all surface sediments quartz content is higher than feldspar content. The Q/Fsp ratio is more than 2.0 in seven samples (PS66/308-1, PS66/312-1, PS66/322-2, PS66/325-2, PS66/330-1, PS66/336-4, PS66/337-2) and between 1.0 and 2.0 at the other stations. Terrigenous carbonates were found to be significant at two stations (PS66/304-2 and PS66/327-2, 4.0% and 6.4%, respectively).

In most of the studied samples the heavy minerals include pyroxenes, amphiboles and black ore minerals (almost everywhere more than 15%). The garnet and epidote content also exceeds 10% in some of the samples investigated onboard. The concentration of other minerals does not reach 10%.

#### 9.6. Characteristics of ARK-XX/3 Sediment Cores

R. Stein, N. Kukina, H. Noffke, A. Puhr, Chr. Schäfer, D. Winkelmann, Y. Yanina

# 9.6.1 X-ray Photographs: Sediment Structure and IRD Content

D. Winkelmann, H. Noffke, A. Puhr, Chr. Schäfer

X-ray imagery of sediment sections has become a common tool in accessing the internal structure of soft marine sediments. Standard slices (10x25x1cm) of sediment continuously taken downcore from all long gravity and kastenlot cores as well as from short cores, have been processed onboard. The structures identified on the photographs reflect marine sedimentary environments ranging from

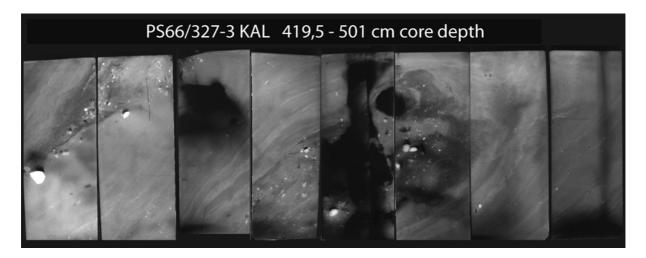
PS66/308-3 SL 029 - 067 cm PS66/308-3 SL 154 - 182 cm

**Fig. 9.18**: Examples of x-ray images from gravity core PS66/308-3 SL situated near the western margin of the mega-slide in the Sophia Basin north of Svalbard (For core location see Figure 9.1).

partly bioturbated pelagic glacio-marine deposits, tubidites and debris flows to diamicton.

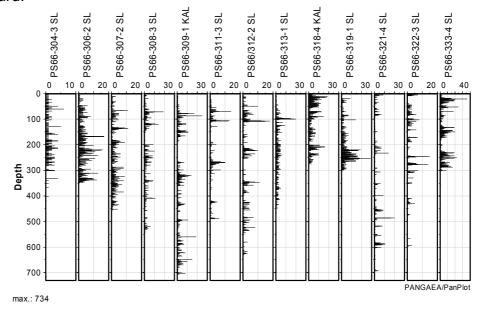
Examples of x-ray images from gravity core PS66/308-3 SL situated near the western margin of the mega-slide in the Sophia Basin north of Svalbard.are shown Figure 9.18. The marginal debris flow / turbidite is visible as unit of sandv silty material (brighter in the image due to lower penetration of the x-rays) with cross bedding and fining upwards sequences indicating higher velocities of current speed (right column of figure). The unit displays individual pulses of sediment accumulation presumably from a single event. The "normal" pelagic sediments partly bioturbated with IRD mirror glacio-marine conditions while layers of higher silt/sand content often laminated may point todeglacial conditions further up in the core (left column of figure).

From the Kastenlot Core PS66/327-3 sediment slices have been taken every ten centimetre to cover large sediment structures of the full core width (Fig. 9.19). The core displays tilted sediment layers as well as sediment folding. Hence standard slices have been taken in 10 centimetres interval to cover almost the full width of the kastenlot. Note the scale of deformational features which would make their identification in a standard gravity core (e.g. Gravity corer with a core barrel of 10 cm in diameter) very unlikely.



**Fig. 9.19**: Examples of x-ray images from Kastenlot Core PS66/327-3 KAL, northern Svalbard continental slope (For core location see Figure 9.1).

The x-ray photographs have also been used for quantification of ice rafted debris (IRD). According to the method of Grobe (1987) the IRD content can be evaluated by counting lithogenic particles with more than one millimetre diameter per ten cubic centimetre (per cm sediment slice). In the IRD records from selected sediment cores, distinct changes in IRD abundances occur (Fig. 9.20), which are probably related to glacial/interglacial variability in extension of glaciation on Svalbard.



**Fig. 9.20**: IRD record of selected gravity and kastenlot cores recovered during Cruise ARK-XX/3 (For core locations see Figure 9.1).

### 9.6.2 Results of Smear-slide Analysis

N. Kukina

Description of the smear slides allow to indicate mineral associations and contents of terrigenous and biogenous components. The results of smear-slide estimates in cores PS66/304, PS66/308, PS66/309, PS66/312, PS66/318, PS66/319, PS66/325, and PS66/333 are shown in Figures 9.21 to 9.28 (for location of cores see Fig. 9.1). In these figures, on one hand the abundances of quartz, feldspars, mica, terrigenous carbonate, volcanic glass, clay minerals, organic remains, Fe-hydroxides, and heavy minerals (summed up to 100%) are shown. On the other hand, the composition of heavy minerals (summed-up to 100%) is shown in separate graphs of the figures. Tables of the smear-slide data are listed in Chapter 12.4 (Annex).

Based on smear-slide analysis, the main minerals are quartz, feldspar, terrigenous carbonates, mica and clay minerals. The heavy minerals are presented by pyroxenes, amphiboles, epidote, garnets, Fe-oxides/hydroxides and black ores. Pyroxenes and amphiboles dominate, but in some intervals the Fe-oxides and hydroxides are enriched.

The study of the mineral composition and lithological description allowed to distinguish the following types of sediments (units), ordered from core top to bottom:

The type 1 sediments are composed of dark brown and dark grayish brown soft silty clay. Clay minerals and quartz dominate in the mineralogical composition. Amphiboles, Fe oxides/hydroxides and black ores increase, terrigenous carbonates and organic remains decrease. The heavy minerals concentration changes from 10% (PS66/308) to 23% (PS66/309).

The type 2 sediments have very dark brown color, and oxidations, Fe-Mn nodules, sandy silt with gravel grain-size composition ("hard ground") were found. The mineral composition is characterized by much higher abundances of Fe oxides/hydroxides in these sediments. Terrigenous and biogenic carbonates are absent. Contents of pyroxenes and epidote increase from 20% to 35%.

The type 3 sediments are characterized by strong variability of lithological parameters. The sediments are dark olive gray sandy silty clay. In the upper part of the unit elevated content of sand were observed. With increase of sand content the heavy minerals content increase also. The terrigenous and biogenic carbonates were completely absent. Only rare amount of organic remains was found. The content of amphiboles and garnets increases.

Type 4 sediments, composed of very viscous gray or dark gray silty clay, are observed in cores PS66/308, -329, and -330. Quartz and feldspar are the main minerals, pyroxenes, epidotes and chlorites dominate in the heavy fraction.

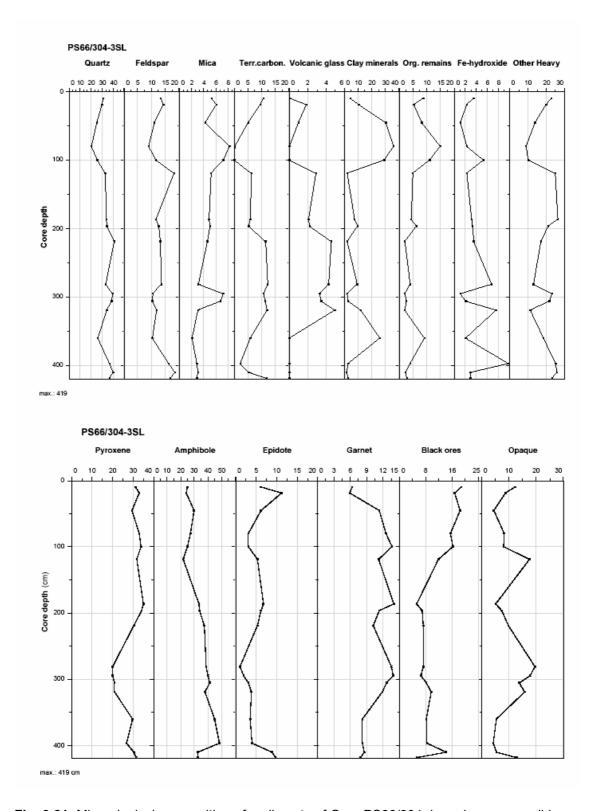
The type 5 sediments (PS66/306, -309, -311, -313) are different from the other units due to high grad of compaction and fine to coarse grain-size. Color changes from dark grayish brown to olive brown and to olive gray are typical.

The sediment is laminated. In the mineralogical composition, a relatively high content of volcanic glass and organic remains and a decrease of content of the heavy minerals were determined.

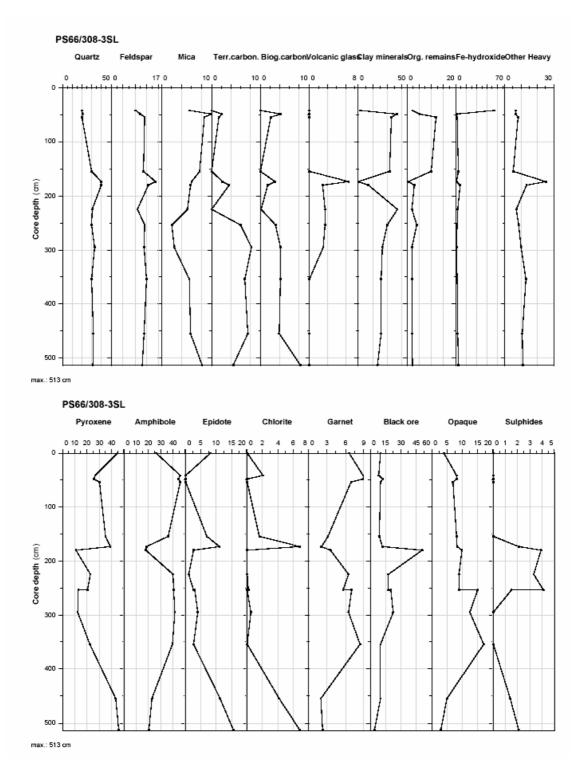
The type 6 sediments (Cores PS66/308 and 309) are of brown, olive gray and dark grayish brown colours and bioturbated, contain some lenses of olive gray sand and dropstones. The main minerals are quartz, feldspar and mica. Volcanic glass and biogenic carbonates are absent.

Type 7 sediments are generally fine-grained and have gray and dark gray colours with black points. Description of smear slide showed increased abundances of sulphides and opaques. There are more amphiboles than pyroxenes. The content of terrigenous and biogenic carbonates is increased (to 16-18% in cores PS66/318 and 319).

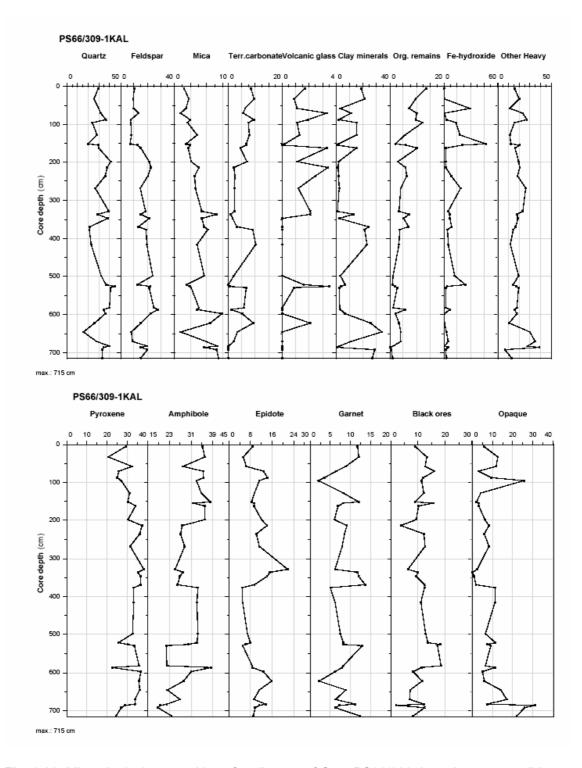
Future studies of mineral assemblages and grain-size composition have to be performed to describe the sediment type and to identify lithological units in more detail.



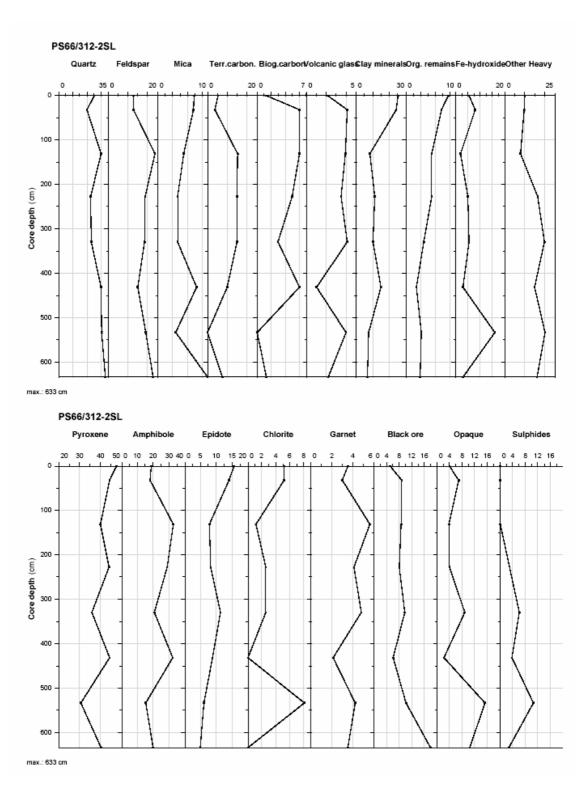
**Fig. 9.21**: Mineralogical composition of sediments of Core PS66/304, based on smear-slide estimates.



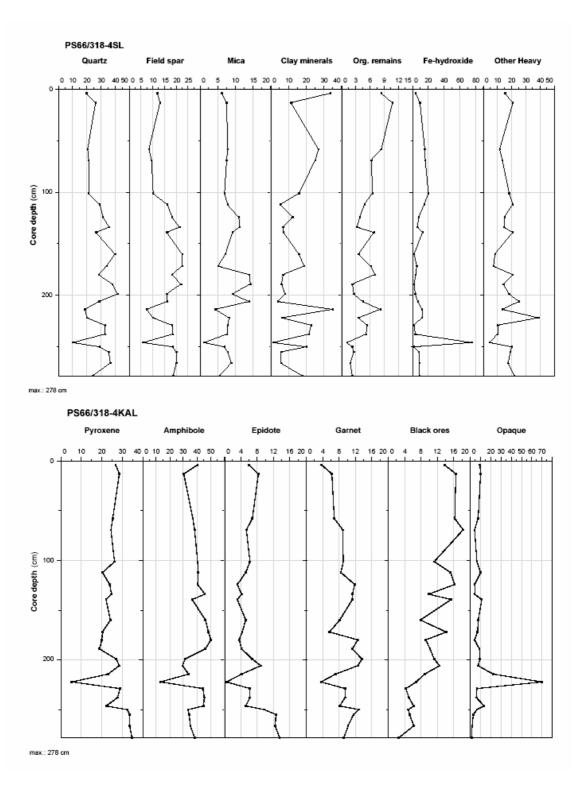
**Fig. 9.22**: Mineralogical composition of sediments of Core PS66/308, based on smear-slide estimates.



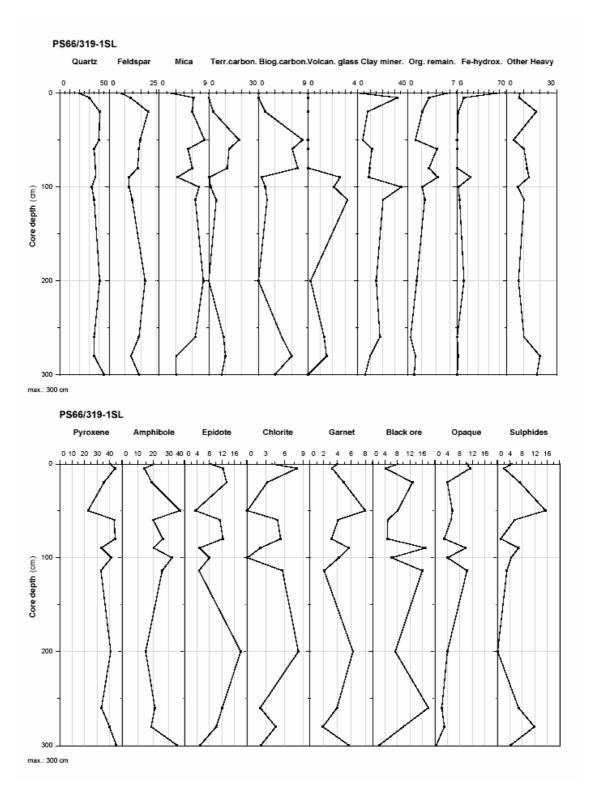
**Fig. 9.23**: Mineralogical composition of sediments of Core PS66/309, based on smear-slide estimates.



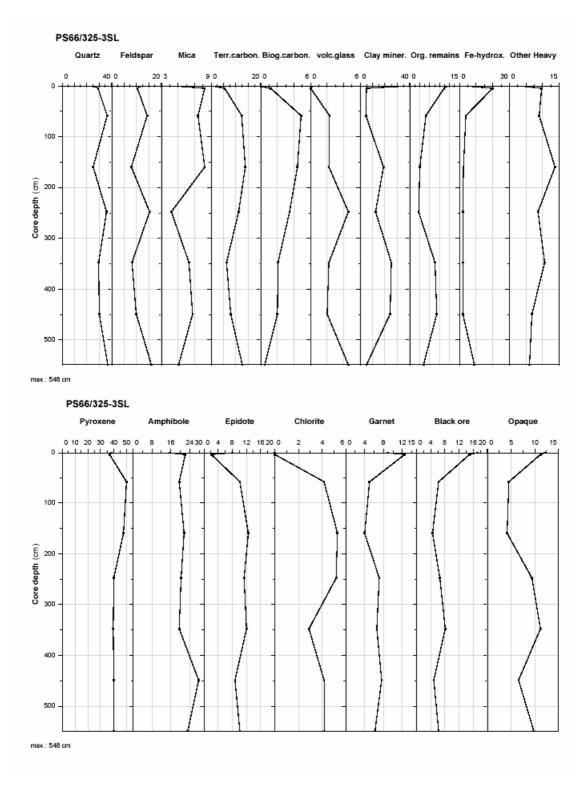
**Fig. 9.24**: Mineralogical composition of sediments of Core PS66/312, based on smear-slide estimates.



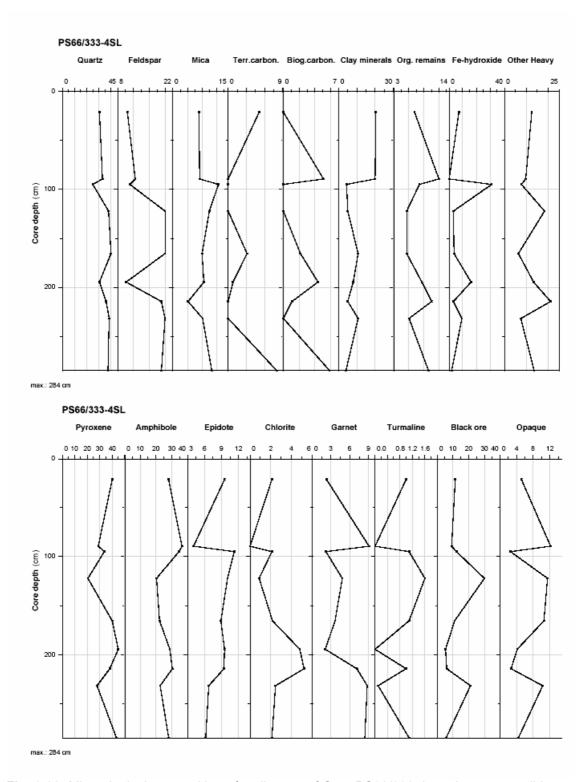
**Fig. 9.25**: Mineralogical composition of sediments of Core PS66/318, based on smear-slide estimates.



**Fig. 9.26**: Mineralogical composition of sediments of Core PS66/319, based on smear-slide estimates.



**Fig. 9.27**: Mineralogical composition of sediments of Core PS66/325, based on smear-slide estimates.



**Fig. 9.28**: Mineralogical composition of sediments of Core PS66/333, based on smear-slide estimates.

### 9.6.3 Results of Coarse Fraction Analysis

J. Thiele, R. Saraswat

Foraminifers (unicellular, almost exclusively marine protists), have been used extensively for paleoclimatic reconstructions and are extremely helpful in providing the chronology based on their oxygen and carbon isotopic composition. Keeping the objective of reconstructing the past climatic variations and especially the dating of megaslide events in view, a preliminary analysis of the coarse fraction was carried out onboard.

The longest Kastenlot core PS66/309-1 recovered during "Polarstern" Cruise ARK XX/3 was investigated in order to assess the coarse-fraction composition of different lithologies. This core was recovered from the north of Spitsbergen (see Figure 9.1 and annex for location details). Based on the differences in colour and also to get an overview of the complete core, 24 samples were processed for coarse fraction analysis as per prescribed procedures (freeze-drying followed by wet sieving over 63µm and 125µm sieves). A small representative fraction from the >125µm fraction was taken with the help of splitter, in order to count a minimum of 300 particles. A total of 15806 particles were counted. The particles were grouped into two categories, i.e., biogenic and abiogenic origin. The abiogenic fraction constituted the larger part of the majority of samples (Fig. 9.29). In the abiogenic fraction, quartz grains with angular to subangular edges were most abundant. In few samples, considerable numbers of pyritized worm tubes were also found. The biogenic fraction mainly consists of planktic foraminiferal tests with few benthic foraminifers and rare sponge spiculae. As reported earlier the planktonic foraminiferal assemblage was dominated by Neogloboquadrina pachyderma, the polar bear of planktonic foraminifers. The planktonic foraminiferal assemblage is almost monospecific. The benthic foraminiferal component is represented by few agglutinated specimens along with considerably large specimens of *Pyrgo* and few specimens of *Cibicidoides*.

From the intervals where sufficient planktonic foraminifers were available, specimens of *N. pachyderma* were picked for AMS dating (five in number) as well as stable isotopic analysis (24 in number).

The correlation of the major constituents at different intervals with the colour of the sediments shows that some of the brown coloured sediments are devoid of biogenic components while some of the gray coloured intervals are high in biogenic components, mainly composed of planktonic foraminiferal tests. Out of the total of 24 samples analysed onboard, the maximum amount of biogenic components (approx. 92%), was reported in the very dark brown coloured silty clay sediments at 179 cm interval. One of the samples (249-252cm) taken from the sandy sediment layer that probably is the result of debris flow, is low in biogenic components. Further analysis of sandy sediments showed the presence of few shallow water benthic foraminifers, probably indicating the transport of sediments from the shelf to the deeper regions from where the core has been recovered. This needs to be confirmed by detailed species identification and subsequent grouping into foraminiferal assemblages.

This preliminary onboard analysis provides an overview of the composition (biogenic and abiogenic) of the different lithologies. For detailed reconstruction of the past climatic events, closer spaced sample analysis has to be carried out.

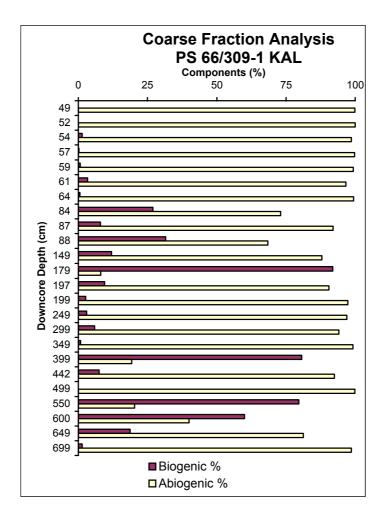


Fig 9.29: Down-core plot of abundance of biogenic and abiogenic components in 24 samples taken for preliminary onboard coarse fraction (>125 $\mu$ m) analysis. The figure shows the broad overview of the composition of different lithologies.

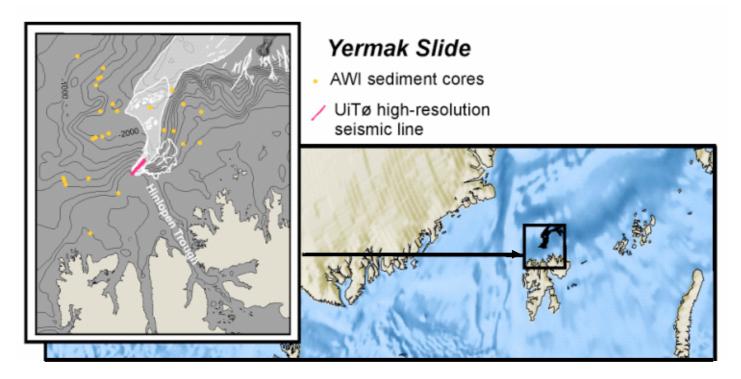
# 9.6.4 Main Lithologies and Preliminary Interpretation of ARK-XX/3 Sediment Cores

R. Stein

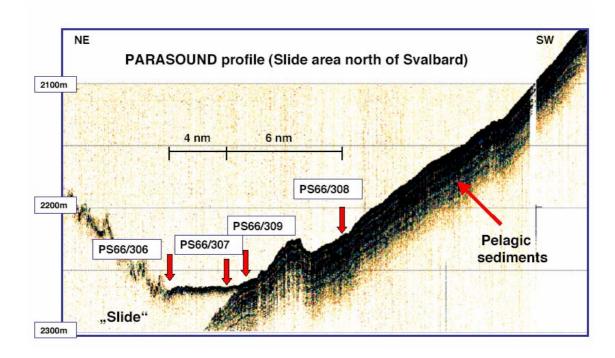
During the "Polarstern" Cruise ARK-XX/3, sediment cores were taken at 26 stations in total (see Chapter 9.4 for details). In the following, the main lithologies of the recovered sediments are described and preliminary interpretations are given. The detailed lithological core descriptions are presented in Chapter 12.3 of the Annex.

Lithostratigraphy of sediment cores from the Svalbard continental margin and eastern slope of the Yermak Plateau

One of the main targets of the ARK-XX/3 geology program is the study of sediment dynamics of mega-slides along the Svalbard continental margin (as part of the ESF Program "EUROMARGINS") (Fig. 9.30). In this context, the characterization of sediment facies within debris flow and turbidite sequences as well as undisturbed (hemi-) pelagic sequences, the dating of sediment mass flows, and estimates of sedimentary budgets are of interest. In order to sample sedimentary sequences within the mega-slide area as well as the adjacent undisturbed (hemi-) pelagic facies, detailed PARASOUND and Hydrosweep surveys have been carried out (see Chapters 7 and 9.2) to select optimum coring stations. At the southwestern flank of the mega-slide area, several couples of sediment cores representing slide-influenced facies and (hemi-) pelagic facies were taken (Fig. 9.31).



**Fig. 9.30**: Assumed extension of the mega-slide ("Yermak Slide") area north of Svalbard (light gray area in the left figure) (Figure taken from ESF-EUROMARGINS proposal on "Slope Stability on Europe's Passive Continental Margins", coordinated by J. Mienert, Tromsø University, Norway).



**Fig. 9.31**: PARASOUND profile and locations of cores PS66/306 to PS66/309 as an example for the sampling strategy. Cores PS66/306 and PS66/307 were taken inside the slide area, PS66/309 at the marginal part, and PS66/308 in the (hemi-) pelagic facies outside the slide area.

A key core within our research will be Kastenlot Core PS66/309-1 obtained close to the edge of the mega slide (Fig. 9.31). This core is composed of very different lithologies as shown in Figure 9.32 (for detailed lithological core description see Chapter 12.3, Annex). Based on sediment colour and texture, the sequence can be divided into nine units. Unit I (0 to 61 cm) is composed of brown and very dark brown silty clay with some layers of mud clasts. Underneath an about 20 cm thick interval of dark gray to very dark gray silty clay is obvious (Unit II). Unit III (89 – 138 cm) is an olive brown, partly bioturbated silty clay. Unit IV (138 - 204 cm) and Unit VI (272 - 299 cm) consist of olive gray silty clay. Between these two units, an interval characterized by dark gray sandy silty clay and sand alternation and fining-upwards sequences (Unit V). This turbidite unit has a sharp (erosional?) contact to the underlying Unit VI. Unit VII (299 - 510 cm) is mainly composed of dark grayish brown, partly bioturbated silty clay with occasional occurrence of very dark gray layers and dropstones. Unit VIII (510 – 641 cm) is characterized by alternation of olive/dark olive gray and dark gray/very dark gray (sandy) silty clay. The uppermost part of this unit is more coarse-grained, and at the bottom of this unit (636 - 641 cm) gray and dark olive brown beds occur. The lowermost unit (Unit IX; 641 - 765 cm) consists of dark gray to very dark gray silts clay.

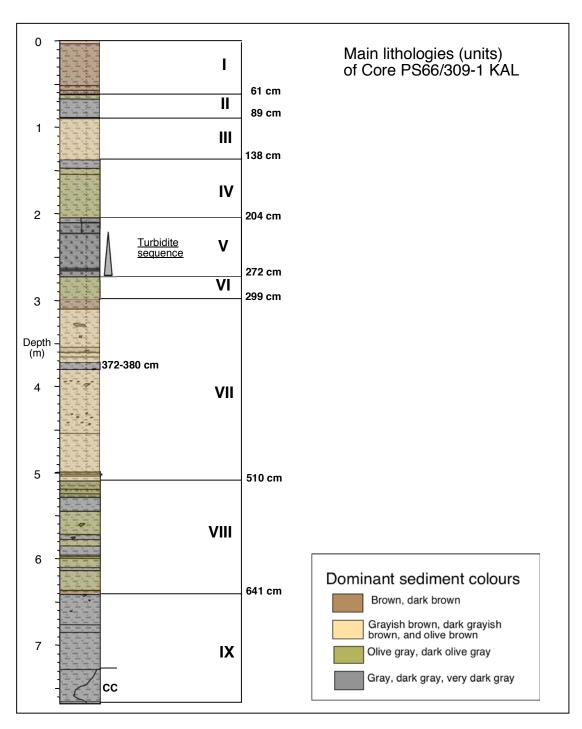
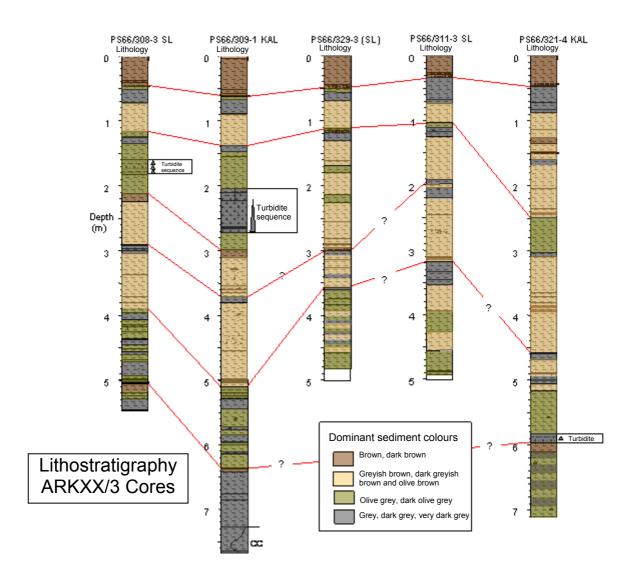


Fig. 9.32: Major lithologies of Kastenlot Core PS66/309-1

The distinct changes in sediment colour (and grain size) in the record of Core PS66/309-1 may represent glacial/interglacial variability (with the dark gray/dark olive gray intervals such as the lower part of the sedimentary sequence probably representing glacial and the brownish intervals probably representing interglacial intervals; e.g., Jakobsson et al. 2000). The turdidite sequence (Unit V) is related to the mega-slide event. The turbidites occur in reduced thickness also in the more distal Core PS66/308 (Fig. 9.33). If a precise dating of the sedi-

ments under- and overlying the turbidite unit can be done, this gives the unique possibility to reconstruct the history of the mega-slide event at the northern Svalbard continental margin.

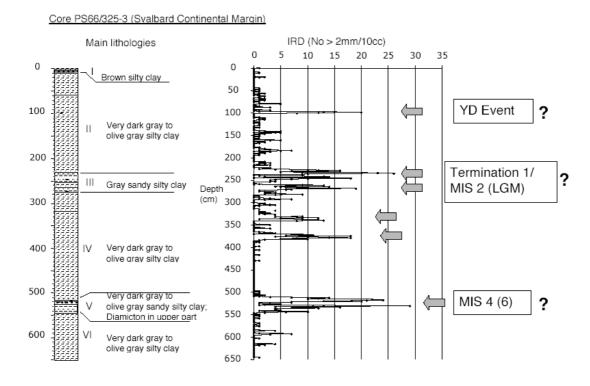
Similar lithologies as described for Core PS66/309-1 were also recorded in the other (pelagic) cores obtained in the study area during this cruise. Based on the lithological core description, a correlation between these cores is possible (Fig. 9.33). In general, this correlation is also supported by the wet bulk density and magnetic susceptibility reords (see Chapter 9.3.1, Figs. 9.11 and 9.12).



**Fig. 9.33**: Major lithologies of cores PS66/308-3, PS66/309-1, PS66/311-3, PS66/321, and PS66/329 (For core locations see Figure 9.1).

The lithology of Core PS66/325-3 taken from the continental slope north of North East Land at a water depth of 895 m, differs from those described above. Except the uppermost dark brown-coloured silty clay, very dark gray and dark olive gray silty clay and sandy silty clay are the dominent type of sediment (Fig.

9.34; for details see Annex 12.3). The sedimentary sequence can be divided into six main lithological units. Unit I (0 to 10 cm) is composed of very dark brown silty clay with mud clasts in the lower part. Units II (10 - 230 cm), IV (274 - 516 cm) and VI (545 - 648 cm) consist of very dark gray to olive gray silty clay. In between these units, two more coarse-grained (sandy silty clay) units are intercalated: Unit III (230 - 274 cm) and Unit V (516 - 545 cm). The upper part of Unit V (516 - 524 cm) is a diamicton-type sediment. These two more coarse-grained units probably represent glacial stages. Furthermore intervals with significant amount of ice-rafted debris (IRD), especially in Units III and V, are obvious (Fig. 9.34). These pulses of maximum IRD input may be related to advances of the near-by Svalbard-Northern Barents Sea Icesheet, as described for Marine Isotope Stages (MIS) 6, 5d, 4, and 2 (e.g., Mangerud et al. 1996).



**Fig. 9.34**: Major lithologies (Units I to VI) and amount of ice-rafted debris (IRD) of Core PS66/325-3. Arrows indicate maxima in IRD input, probably related to fluctuations (advances) of the Svalbard-Northern Barents Sea Icesheet during glacial stages.

Future sedimentological, micropaleontological, and geochemical studies and, especially, the dating of the sediment cores are needed for more detailed interpretation in terms of paleoenvironmental history.

Sediments and rocks from the top of the central Yermak Plateau: Implications for basement origin?

As outlined in Chapter 8, it is assumed that the northeastern Yermak Plateau is oceanic in origin, while its southern part might consist of stretched continental crust. Thus, on top of the central Yermak Plateau where old strata (basement?) probably terminate close to the seafloor, the sediment dredge as well as the giant box corer and the gravity corer was used for sampling at several stations (Fig. 9.35). The sampling stations were selected based on detailed PARA-SOUND and Hydrosweep profiling. Fur this purpose, steep slopes and peak elevations with lacking young soft sediments on top were of interest (Fig. 9.36).

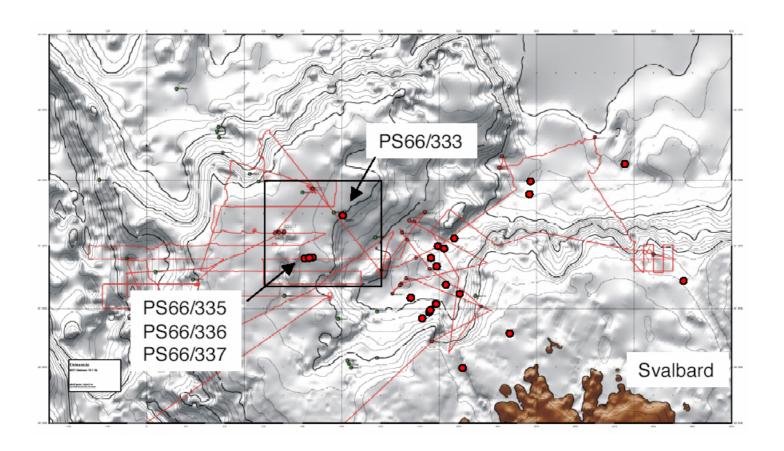
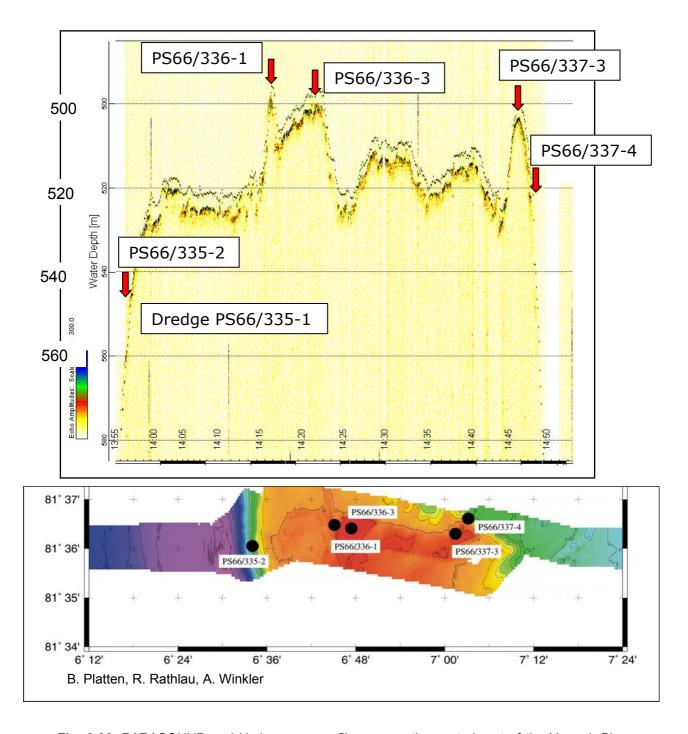


Fig. 9.35: Stations of "basement sampling" on top of the central Yermak Plateau



**Fig. 9.36**: PARASOUND and Hydrosweep profiles across the central part of the Yermak Plateau with sampling stations.

At the western steep slope (location PS66/335), huge amounts of stones have been dredged (Fig. 9.37; for exact dredge location see Table 9.3). The majority of stones is black-coloured/-coated. After cutting, however, a large number of stones were identified as red sandstones. At the same location, also a gravity core (PS66/335-2) was taken. In this core, only 32 cm of sediments were recovered. Part of the recovered sequence is composed of – for this area – non-typical red-coloured sediments (Fig. 9.38).



Fig. 9.37: Dredged rocks at Station PS66/335, central Yermak Plateau

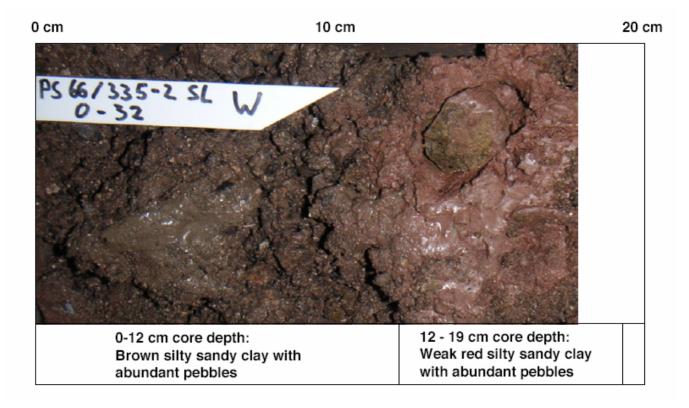


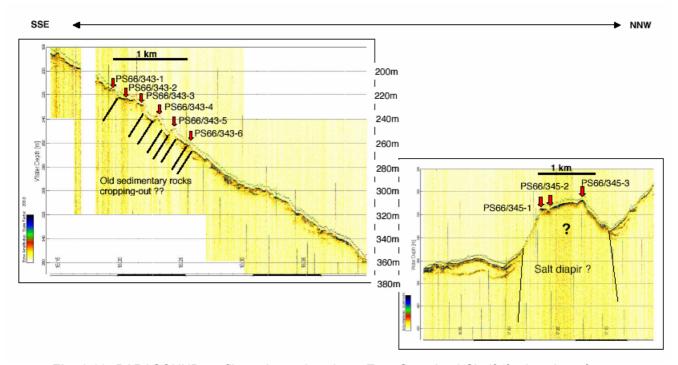
Fig. 9.38: Sediments recovered in gravity core PS66/335-2, central part of the Yermak Plateau

At sites PS66/337-3 and PS66/337-4 located in the eastern part of the profile (Fig. 9.36), no penetration with the gravity corer was possible. At Site PS66-337-3 the core catcher of the gravity corer even became damaged when hitting the hard "basement" rocks at the seafloor. In the core catcher of both cores, mainly black (basalt-like looking crystalline) stones or fragments of them were recovered.

Based on these very preliminary results and still speculative interpretation, the basement may have a continental origin (Devonian Old-Red sandstones?) at the western slope of the profile whereas in the eastern part basaltic material may be cropping out, indicating an oceanic origin. Of course, more detailed work on the petrology of the dredged material should be performed before a more precise interpretation about the basement origin of the Yermak Plateau in this area can be done.

Sediments from the East Greenland continental shelf: Pre-Quaternary basement and Quaternary paleoenvironment

As part of the IODP site survey, seismic profiling was carried out on the east Greenland shelf during this expedition (see Chapter 8). In this area, seismic data reveal the presence of a prominent salt province and suggest that Mesozoic sediment are cropping out at the surface (Schmitz and Jokat, 2005). In order to sample these old strata, 10 gravity cores (using the short 3m core barrel) were taken. The location of the cores was selected based on PARASOUND profiling (Fig. 9.39). Based on comparison with the seismic profile, faulted old sediment strata and a salt diapir can also be assumed in the PARASOUND profile.



**Fig. 9.39**: PARASOUND profile and core locations, East Greenland Shelf (for location of area see Fig. 9.1).

In the southern part of the profile (PS66/343-1 to PS66/343-6), where pre-Quaternary (Mesozoic?) sediments should be sampled, unfortunately only soft (young) sediments were recovered (Fig. 9.40). Some of the gravity corers even over-penetrated into the seafloor. At Site PS66/345 located on top of the assumed salt diapir structure (Fig. 9.39), however, coring of old strata was successful. At Station PS66/345-2, a 1.5 m long gravity core was recovered (Fig. 9.41). In the lowermost part of the sequence(core catcher), small light gray "crystals" and one large white-orange "crystal" of about 4 cm in diameter were found (Fig. 9.42-A). X-Ray analysis of the large "crystal" indicates that it is pure gypsum (Fig. 9.42-B). This may suggest that the gravity corer may have hit the (Permian?) salt diapir surface (or came at least very close to it). This important finding may support the interpretation of the seismic data of Schmitz and Jokat (2005).

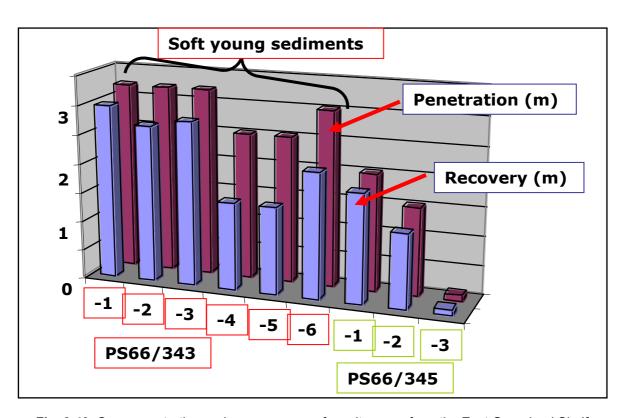


Fig. 9.40: Corer penetration and core recovery of gravity cores from the East Greenland Shelf.

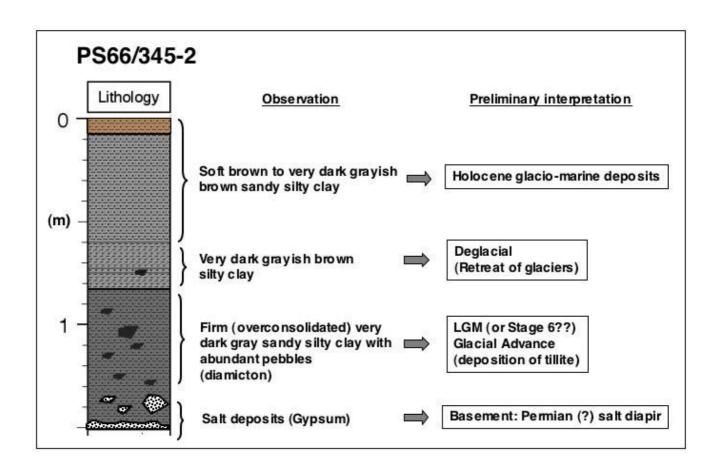
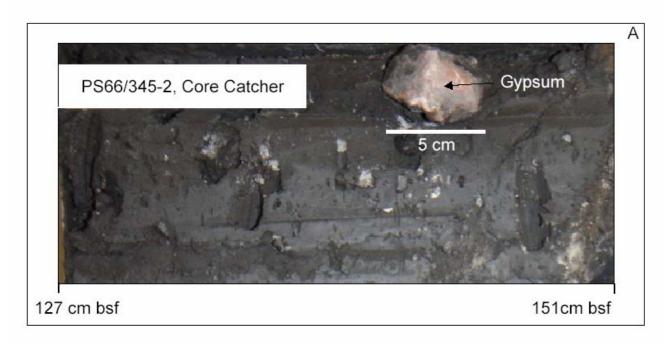
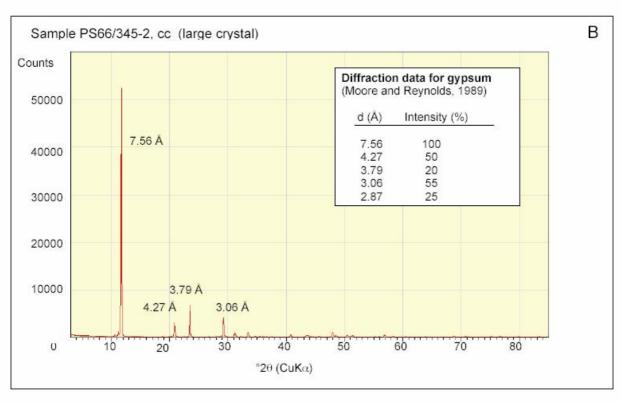


Fig. 9.41: Lithology and preliminary interpretation (Core PS66/345-2, East Greenland Shelf)





**Fig. 9.42**: (A) Photograph of core catcher sediments recovered in gravity core PS66/345-2 (East Greenland shelf). Common small white rock fragments ("crystals") and one large "crystal" of about 4 cm in diameter are obvious. (B) X-Ray diffractogram of the large crystal found in the core catcher, identified as gypsum (X-Ray analysis has been performed by Chr. Vogt, Bremen University).

The main part of the sedimentary record of Core PS345-2 that is certainly of Quaternary age can be divided into three main units (Fig. 9.41).

- Unit I (0 60 cm): soft brown to very dark grayish brown sandy silty clay
- Unit II (60 83 cm): very dark grayish brown silty clay
- Unit III (83 base): firm (overconsolidated) very dark gray sandy clay with abundant pebbles (diamicton)

From correlation with similar sequences from the East Greenland continental margin (e.g., Stein et al., 1993; Hubberten et al., 1995; Funder et al., 1999), these lithologies can be interpreted in terms of East Greenland glacial history. The sediments of Unit III (a tillite) probably represent the Last Glacial Maximum (LGM) glacial advance. The site was overridden by glaciers as shown by overconsolidation. Overconsolidated sediments are also reflected in the shear-strength record (see Chapter 9.3.3, Fig. 9.15). The deglacial retreat of glaciers is documented in the sediments of Unit II. The uppermost Unit may represent the glacial-marine post-glacial (Holocene) time interval. Finally, it has to be mentioned that it cannot be excluded at this stage, that the diamicton of Unit III may be related to older (Marine Isotope Stage 6?) glaciations. AMS<sup>14</sup>C datings are needed to give a more precise interpretation of the sedimentary record of Core PS66/345.

## 10. The Lost Deutsche Arktische Expedition 1912: A Pilot Study in the North East Land of Spitsbergen N.Fricke, I.Mende

In 1912, the world's attention focussed around events in the Antarctic: Scott and Amundsen's race to the South Pole. Almost simultaneously similar heroic, pathetic and tragic events unfolded on Spitsbergen – the lost Deutsche Arktische Expedition (DAE) 1912, also known as Schroeder Stranz expedition: a trial expedition to the North East Land of Spitsbergen (Fig. 10.1) thought to be a forerunner of the ambitious plan to cross the North East passage. The illprepared expedition ended in chaos and disintegration. Of the 10 German members, 7 including lieutenant Schroeder Stranz, died. It became Germany's most tragic polar expedition.

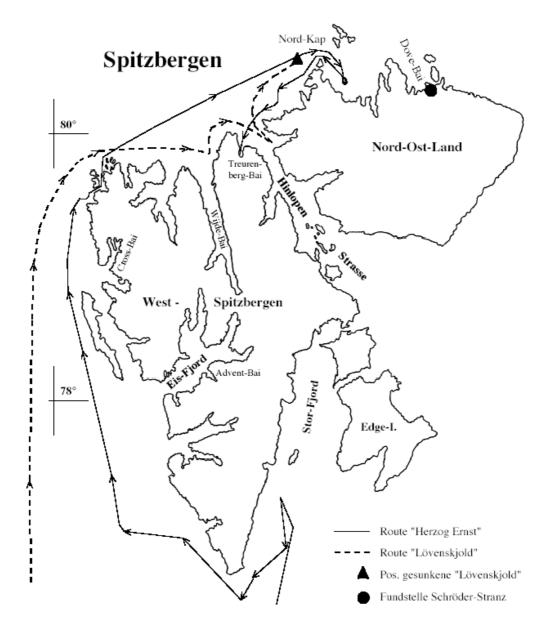


Fig. 10.1: Locations and tracks of lost "Deutsche Arktis Expedition 1912 (DAE)"

One of the objectives of the present pilot study during the "Polarstern" Cruise ARK-XX/3 was to get preliminary information and to confirm some of the positions of the Schroeder Stranz Expedition of 1912 and the Lerner Expedition of 1913, as available in the literature.

Our first search point was Cape Rubin situated at the western side of the North Cape of North East Land (Fig. 10.1). We were interested in this place, since in 1913 the German expedition lead by Theodor Lerner and organized to search the lost German Artic expedition, the "Schroeder Stranz Expedition" of 1912, started from the North East Land of Svalbard. The participants of Lerner's expedition built a base at Cape Rubin and spent two month in this camp and searched around 600 km of coastline, unfortunately without any success.

The expedition vessel "Loevenskiold" of Lerner got entrapped in the pack ice and sunk off the Beverlysund near the North Cape. This wreck is the most northern wreck of the world. It is important to know the precise position of the place, where the ship sank. The information available about the Learner's expedition included the latitude and the water depth of the place of accident from the ships' logbook as well as from historical pictures shot by the participants of the Lerner expedition. The pictures shot from the land, show the "Loevenskiold" in the ice, whereas in other pictures taken from the ship, the camp with rock formations in the background is shown. These pictures helped in locating the place of accident.

On September 03, when "Polarstern" was near to the first search point, we flew to North Cape/Cape Rubin (Fig. 10.1). Supported by a helicopter and the pictures we had, we could locate the place of the camp. We found some remains of the camp, which included old clothes, wooden snow-shoes (Fig. 10.2a), shekels (Fig. 10.2b) and cartridge cases (Fig. 10.2c). Though taking a bearing is not an appropriate way to find this wrack, the coordinates we got from the present pilot study will definitely help to find the shipwreck later with the submarine "Jago" and other appropriate techniques.

a)





c)



Fig. 10.2: Remains of the camp at Cape Rubin (a) wooden snow-shoes, (b) shekels, and (c) cartridge cases

Our second search location was the assumed position where remains of the Schroeder Stranz Expedition were reported in 1937. The fisherman Amandus Wilhelmsen found equipment of the 1912 expedition in the Duvefjord (Dove Bay) area (Fig. 10.1). He gave coordinates as well as the description of the place. But the coordinates and the description of the place as given by Amandus Wilhelmsen do not fit together. In the planning of our project we selected three possible locations that we wanted to confirm.

In search for this place, first we flew to the Schroeder Stranz Eldet (Fig. 10.3) on September 14. Fortunately, this seems to be a place which is concurrent with the description of Wilhelmsen. But, unfortunately, we could not find any fur-

ther remains of the expedition because of bad weather. Fresh snow covered the ground, and continuous heavy snowfall made a long stay with a substantial search almost impossible. We shot some films and photographs at this place. One hour later we flew to the two other possible locations. From the helicopter we observed that these two places did not at all fit to the description of the place where Amandus Wilhelmsen found the remains.



Fig. 10.3: Schroeder Stranz Eldet, Duvefjord area, North East Land

Our next point of interest was from where Wilhelm Dege, a soldier of the secret German weather station "Haudegen", made an interesting discovery in 1945, few dishes that were later on recognised by Prof. Dr. Hans Fricke (MPI Seewiesen) as the remains of Schroeder Stranz Expedition. The place of recovery of these remains is known. Unfortunately, we could stay there only for five minutes because of the bad weather conditions. After taking some pictures and making video sequences we had to leave.

In summary, the major goals of this pilot study were reached. We could (1) precisely document the place of recovery of the remains from the Lerner expedition of 1913, (2) ascertain the coordinates where the "Loevenskiold" sunk, (3) locate the place where Amandus Wilhelmsen found remains of the Schroeder Stranz Expedition and (4) collect video clippings and photographs of the region around Cape Rubin and Dove-Bay. The results will help in the more exact planning of the multidisciplinary submersible expedition of the shallow shelf and fjords of North Spitsbergen in 2006.

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Table 12.1: Station List

Station	Date	Time	Position Lat.	Position Long.	Depth [m]	Gear	Remarks
PS66/300-1	01.09.04	00:25	70° 31,46' N	19° 16,14' E	190,4	M	start- magnetic turncircels
PS66/300-1	01.09.04	01:07	70° 31,33' N	19° 16,33' E	187,2	M	end-magnetic turncircels
PS66/301-1	02.09.04	22:22	78° 50,25' N	2° 48,20' E	2496,0	PIES 71	
PS66/302-1	03.09.04	00:45	78° 49,87' N	5° 0,93' E	2707,0	PIES 141	
PS66/303-1	03.09.04	04:15	78° 50,03' N	8° 19,91' E	793,0	PIES 61	
PS66/304-1	03.09.04	16:07	80° 43,36' N	14° 39,37' E	1098,0	GKG	
PS66/304-2	03.09.04	16:57	80° 43,21' N	14° 38,47' E	1093,0	MUC	
PS66/304-3	03.09.04	17:40	80° 43,20' N	14° 38,02' E	1095,0	GC	
PS66/305-1	03.09.04	18:03	80° 42,95' N	14° 38,91' E	1058,0	HS_PS	start track
PS66/305-1	04.09.04	01:38	81° 14,62' N	13° 18,90' E	2269,0	HS_PS	profile end
PS66/306-1	04.09.04	02:17	81° 14,58' N	13° 18,96' E	2271,0	GKG	
PS66/306-2	04.09.04	03:34	81° 14,56' N	13° 18,62' E	2268,0	GC	
PS66/307-1	04.09.04	05:11	81° 11,95' N	13° 2,74' E	2277,0	GKG	
PS66/307-2	04.09.04	06:21	81° 11,94' N	13° 2,68' E	2275,0	GC	
PS66/308-1	04.09.04	08:22	81° 7,28' N	12° 35,96' E	2217,0	GKG	
PS66/308-2	04.09.04	09:38	81° 7,31' N	12° 36,16' E	2217,0	MUC	
PS66/308-3	04.09.04	10:47	81° 7,30' N	12° 35,97' E	2218,0	GC	
PS66/309-1	04.09.04	13:44	81° 11,22' N	12° 59,08' E	2269,0	KAL	
PS66/310-1	04.09.04	14:40	81° 12,20' N	13° 7,40' E	2271,0	HS_PS	start track
PS66/310-1	04.09.04	23:38	81° 41,50' N	13° 28,00' E	2198,0	HS_PS	profile end
PS66/311-1	05.09.04	00:20	81° 41,76' N	13° 28,25' E	2192,0	GKG	
PS66/311-2	05.09.04	01:33	81° 41,79' N	13° 28,45' E	2193,0	MUC	
PS66/311-3	05.09.04	02:53	81° 41,78' N	13° 28,20' E	2192,0	GC	
PS66/312-1	05.09.04	04:18	81° 41,26' N	13° 42,48' E	2273,0	GKG	
PS66/312-2	05.09.04	05:30	81° 41,24' N	13° 42,65' E	2275,0	GC	
PS66/313-1	05.09.04	07:35	81° 45,56' N	14° 16,31' E	2298,0	GC	
PS66/314-1	05.09.04	11:35	81° 17,87' N	11° 59,47' E	1337,0	M	start- magnetic turncircels
PS66/314-1	05.09.04	12:21	81° 17,83' N	11° 59,10' E	1347,0	M	end-magnetic turncircels
PS66/315-1	05.09.04	13:11	81° 16,46' N	12° 6,91' E	1708,0	SEISREFL	Streamer into water
PS66/315-1	05.09.04	13:37	81° 17,36' N	12° 3,00' E	1477,0		airguns in the water
PS66/315-1	05.09.04	13:40	81° 17,50' N	12° 2,39' E	1440,0		profile start
PS66/317-1	10.09.04	10:00	80° 53,75' N	3° 34,35' E	791,4	M	start- magnetic turncircels
PS66/316-1	10.09.04	10:01	80° 53,73' N	3° 34,76' E	787,1	SEISREFL	array on deck
PS66/317-1	10.09.04	11:24	80° 54,38' N	3° 41,51' E	761,5	M	end-magnetic turncircels
PS66/316-1	10.09.04	11:39	80° 55,05' N	3° 39,79' E	763,0		airguns in the water
PS66/316-1	10.09.04	13:24	80° 54,01' N	2° 50,47' E	-	SEISREFL	cont. Profile
PS66/316-1	12.09.04	13:22	81° 35,98' N	12° 29,68' E	2073,0		end of profile
PS66/316-1	12.09.04	13:39	81° 36,50' N	12° 34,26' E	2085,0		streamer on deck
PS66/316-1	12.09.04	13:48	81° 36,77' N	12° 36,70' E	2091,0	SEISREFL	array on deck
PS66/318-1		16:40	81° 18,52' N	12° 2,55' E	1261,0	SD	
PS66/318-2	12.09.04	17:09	81° 18,60' N	12° 2,39' E	1261,0	GKG	
PS66/318-3	12.09.04	18:00	81° 18,58' N	12° 2,87' E	1260,0	MUC	
PS66/318-4	12.09.04	18:46	81° 18,60' N	12° 2,71' E	1262,0	KAL	
PS66/319-1	12.09.04		81° 24,49' N	13° 48,79' E	2178,0	GC	
PS66/320-1	12.09.04	23:00	81° 24,60' N	13° 48,96' E	2175,0	HS_PS	start track
PS66/320-1	13.09.04 13.09.04	08:38	82° 10,34' N	18° 15,35' E	2361,0	HS_PS	profile end
PS66/321-1		08:45	82° 10,38' N	18° 15,02' E	2359,0	SD	
PS66/321-2	13.09.04	09:21 10:40	82° 10,38' N	18° 14,89' E	2359,0	GKG MUC	
PS66/321-3	13.09.04	11:51	82° 10,32' N 82° 10,37' N	18° 14,36' E	2360,0 2360,0	MUC GC	
PS66/321-4 PS66/322-1	13.09.04 13.09.04	14:04	82° 10,27' N 82° 5,36' N	18° 13,47' E 18° 5,27' E	3041,0	SD	
PS66/322-1	13.09.04	14:47	82° 5,47' N	18° 4,13' E	2992,0	GKG	
PS66/322-3	13.09.04	16:16	82° 5,40' N	18° 5,25' E	3028,0	GC	
PS66/323-1	14.09.04	02:12	82° 18,48' N	22° 58,77' E	3910,0	GKG	
PS66/323-2	14.09.04	04:14	82° 18,69' N	22° 59,40' E	3910,0	MUC	
PS66/323-3	14.09.04	06:06	82° 18,44' N	22° 59,40 E	3914,0	GC	
PS66/324-1	14.09.04	14:52	81° 32,62' N	24° 54,50' E	1513,0		Streamer into water
PS66/324-1	14.09.04	15:15	81° 31,81' N	25° 2,30' E		SEISREFL	airguns in the water
PS66/324-1	14.09.04	15:45	81° 30,07' N	25° 0,00' E	846,1		profile start
PS66/324-1	14.09.04	18:08		25° 9,70' E		SEISREFL	Sono-buoy
1 000/024-1	17.00.04	10.00	0. 10,00 N	20 3,10 L	505,6		30110-buoy

Station	Date	Time	Position Lat.	Position Long.	Depth [m]	Gear	Remarks
PS66/324-1	15.09.04	03:10	81° 18,07' N	26° 0,17' E	537,8	SEISREFL	Sono-buoy
PS66/324-1	15.09.04	13:31	81° 24,01' N	26° 59,71' E	871,2		end of profile
PS66/324-1	15.09.04	13:52	81° 23,96' N	27° 8,43' E		SEISREFL	streamer on deck
PS66/324-1	15.09.04	13:56	81° 23,95' N	27° 9,44' E		SEISREFL	array on deck
PS66/325-1	15.09.04	15:35	81° 26,19' N	25° 59,85' E	896,7	SD	,
PS66/325-2	15.09.04	15:57	81° 26,17' N	25° 59,86' E	896,2	GKG	
PS66/325-3	15.09.04	16:32	81° 26,18' N	25° 59,86' E	895,5	GC	
PS66/326-1	15.09.04	17:00	81° 26,31' N	25° 56,93' E	864,7	HS_PS	start track
PS66/326-1	16.09.04	13:28	80° 43,05' N	15° 51,42' E	642,4	HS_PS	profile break
PS66/326-1	16.09.04	14:22	80° 43,30' N	15° 52,45' E	655,1		continue the profile
				,	-	HS_PS	-
PS66/326-1	16.09.04	16:30	81° 1,04' N	17° 9,11' E	703,5	HS_PS	profile end
PS66/327-1	16.09.04	16:32	81° 0,99' N	17° 9,54' E	702,0	SD	
PS66/327-2	16.09.04	16:47	81° 0,93' N	17° 9,60' E	701,0	GKG	
PS66/327-3	16.09.04	17:18	81° 0,92' N	17° 9,39' E	701,0	KAL	
PS66/328-1	16.09.04	21:03	81° 19,25' N	14° 30,88' E	1979,0	GC	
PS66/329-1	17.09.04	01:08	81° 36,25' N	13° 6,19' E	2209,0	GKG	
PS66/329-2	17.09.04	02:25	81° 36,27' N	13° 6,14' E	2207,0	MUC	
PS66/329-3	17.09.04	03:44	81° 36,31' N	13° 6,13' E	2211,0	GC	
PS66/330-1	17.09.04	05:37	81° 32,87' N	13° 21,77' E	2306,0	GKG	
PS66/330-2	17.09.04	06:44	81° 32,88' N	13° 21,71' E	2305,0	GC	
PS66/331-1	17.09.04	08:16	81° 26,53' N	13° 5,68' E	2302,0	SEISREFL	Streamer into water
PS66/331-1	17.09.04	08:33	81° 25,53' N	13° 3,20' E	2306,0	SEISREFL	airguns in the water
PS66/331-1	17.09.04	11:57	81° 24,01' N	11° 2,03' E	1940,0	SEISREFL	Sono-buoy
PS66/331-1	18.09.04	01:10	81° 12,02' N	7° 31,45' E	787,2	SEISREFL	Sono-buoy
PS66/331-1	18.09.04	09:24	81° 17,97' N	9° 57,31' E	1784,0	SEISREFL	Sono-buoy
PS66/331-1	18.09.04	11:27	81° 17,97' N	8° 42,59' E	1205.0	SEISREFL	end of profile
PS66/331-1	18.09.04	11:37	81° 17,98' N	8° 36,95' E	-	SEISREFL	profile start
PS66/331-1	19.09.04	01:08	81° 35,23' N	3° 21,56' E	843,4		Sono-buoy
PS66/331-1	19.09.04	08:26	82° 8,01' N	4° 19,07' E	-	SEISREFL	end of profile
PS66/331-1	19.09.04	08:37	82° 8,38' N	4° 18,88' E		SEISREFL	streamer on deck
PS66/331-1	19.09.04	08:41	82° 8,40' N	4° 18,81' E		SEISREFL	array on deck
PS66/332-1	19.09.04	15:10	81° 58,10' N	8° 19,62' E	791,0	HS_PS	start track
PS66/332-1	19.09.04	17:14	81° 54,55' N	8° 29,17' E	752,0	HS_PS	profile end
PS66/333-1	19.09.04	17:14	81° 55,19' N	8° 35,81' E	747,0	SD	profile end
				8° 34,29' E	-		start dradging
PS66/333-2	19.09.04	17:54	81° 55,54' N	,	751,5	DRG_C	start dredging
PS66/333-2	19.09.04	19:05	81° 56,32' N	8° 31,62' E	720,8	DRG_C	stop dredging
PS66/333-3	19.09.04	20:45	81° 56,53' N	8° 22,74' E	718,5	DRG_C	start dredging
PS66/333-3	19.09.04	21:34	81° 56,49' N	8° 22,77' E	716,8	DRG_C	stop dredging
PS66/333-4	19.09.04		81° 56,42' N	8° 30,17' E	725,5	GC	
PS66/333-5	19.09.04		81° 56,10' N	8° 31,50' E	719,1	GC	
PS66/334-1	20.09.04		81° 36,99' N	6° 39,26' E	536,7	HS_PS	start track
PS66/334-1	20.09.04	04:39	81° 36,32' N	6° 32,22' E	819,5	HS_PS	profile end
PS66/335-1		05:08	81° 35,99' N	6° 32,00' E	863,7	DRG_C	start dredging
PS66/335-1	20.09.04	05:29	81° 36,02' N	6° 35,42' E	558,5	DRG_C	stop dredging
PS66/335-2	20.09.04	06:47	81° 36,05' N	6° 34,06' E	654,8	GC	
PS66/336-1	20.09.04	07:40	81° 36,48' N	6° 45,14' E	505,2	GC	
PS66/336-2	20.09.04	08:09	81° 36,35' N	6° 47,74' E	502,6	SD	
PS66/336-3	20.09.04	08:22	81° 36,41' N	6° 47,40' E	505,9	GC	
PS66/336-4	20.09.04	08:56	81° 36,39' N	6° 47,69' E	504,5	GKG	
PS66/337-1	20.09.04	09:30	81° 36,33' N	7° 1,32' E	505,1	SD	
PS66/337-2	20.09.04	09:43	81° 36,32' N	7° 1,38' E	503,3	GKG	
PS66/337-3	20.09.04	10:29	81° 36,30' N	7° 1,44' E	503,0	GC	
PS66/337-4		11:12	81° 36,61' N	7° 3,18' E	587,8	GC	
PS66/338-1	20.09.04	16:11	81° 19,05' N	3° 2,82' E		SEISREFL	airguns in the water
PS66/338-1	20.09.04	16:15	81° 18,96' N	3° 1,54' E		SEISREFL	Streamer into water
PS66/338-1	20.09.04	16:30	81° 18,51' N	2° 55,19' E		SEISREFL	airguns in the water
PS66/338-1	21.09.04	13:36	81° 6,01' N	5° 16,66' E		SEISREFL	Sono-buoy
PS66/338-1	21.09.04	20:00	81° 6,00' N	9° 14,96' E			- 1
			=			SEISREFL	end of profile
PS66/338-1	21.09.04		81° 5,99' N	9° 19,67' E		SEISREFL	streamer on deck
PS66/338-1	21.09.04		81° 5,95' N	9° 20,57' E	1141,0		array on deck
PS66/339-1		20:21	81° 5,94' N	9° 22,12' E	1157,0	M	start- magnetic turncircels
PS66/339-1	21.09.04	22:09	81° 5,96' N	9° 24,00' E	1178,0	M	end-magnetic turncircels
PS66/340-1	22.09.04	10:15	79° 55,10' N	0° 41,87' W	2747,0	SEISREFL	Streamer into water

Station	Date	Time	Position Lat.	Position Long.	Depth [m]	Gear	Remarks
PS66/340-1	22.09.04	10:32	79° 54,58' N	0° 48,57' W	2766,0	SEISREFL	airguns in the water
PS66/340-1	22.09.04	10:43	79° 53,97' N	0° 51,64' W	2775,0	SEISREFL	profile start
PS66/340-1	22.09.04	13:03	79° 42,26' N	0° 47,94' W	2821,0	SEISREFL	Sono-buoy
PS66/340-1	22.09.04	22:17	79° 41,92' N	0° 38,95' W	2826,0	SEISREFL	end of profile
PS66/340-1	22.09.04	22:33	79° 41,26' N	0° 39,08' W	2823,0	SEISREFL	streamer on deck
PS66/340-1	22.09.04	22:36	79° 41,18' N	0° 39,13' W	2826,0	SEISREFL	array on deck
PS66/341-1	22.09.04	23:43	79° 44,07' N	0° 44,97' W	2822,0	GC	
PS66/342-1	23.09.04	13:17	79° 3,51' N	10° 57,87' W	357,9	SEISREFL	airguns in the water
PS66/342-1	23.09.04	13:27	79° 4,04' N	10° 57,85' W	354,8	SEISREFL	array on deck
PS66/342-2	23.09.04	13:41	79° 5,22' N	10° 59,11' W	345,6	SEISREFL	Streamer into water
PS66/342-2	23.09.04	13:59	79° 3,85' N	10° 59,04' W	360,0	SEISREFL	airguns in the water
PS66/342-2	23.09.04	14:11	79° 3,03' N	10° 59,50' W	358,9	SEISREFL	profile start
PS66/342-2	24.09.04	02:45	78° 59,59' N	10° 52,84' W	281,9	SEISREFL	end of profile
PS66/342-2	24.09.04	03:02	78° 59,58' N	10° 48,25' W	313,7	SEISREFL	streamer on deck
PS66/342-2	24.09.04	03:07	78° 59,77' N	10° 47,95' W	319,7	SEISREFL	array on deck
PS66/343-1	24.09.04	03:50	79° 0,83' N	11° 6,91' W	216,2	GC	
PS66/343-2	24.09.04	04:20	79° 0,98' N	11° 6,82' W	226,0	GC	
PS66/343-3	24.09.04	04:49	79° 1,05' N	11° 6,88' W	227,2	GC	
PS66/343-4	24.09.04	05:18	79° 1,17' N	11° 7,12' W	240,9	GC	
PS66/343-5	24.09.04	05:44	79° 1,31' N	11° 6,80' W	252,2	GC	
PS66/343-6	24.09.04	06:09	79° 1,44' N	11° 6,84' W	264,2	GC	
PS66/344-1	24.09.04	06:48	79° 2,55' N	11° 6,65' W	338,7	GC	
PS66/345-1	24.09.04	07:30	79° 3,58' N	11° 15,11' W	317,9	GC	
PS66/345-2	24.09.04	08:08	79° 3,58' N	11° 16,03' W	313,3	GC	
PS66/345-3	24.09.04	08:40	79° 3,58' N	11° 17,37' W	308,6	GC	
PS66/346-1	25.09.04	04:08	77° 19.27' N	1° 21.36' E	3261.0	SEISREFL	Streamer into water
PS66/346-1	25.09.04	04:24	77° 18.75' N	1° 24.87' E	3261.0	SEISREFL	airguns in the water
PS66/346-1	25.09.04	10:16	77° 18.09' N	1° 5.88' E	3261.0	SEISREFL	Sono-buoy
PS66/346-1	25.09.04	13:06	77° 5.81' N	0° 54.19' E	3255.0	SEISREFL	Sono-buoy
PS66/347-1	26.09.04	09:51	75° 47.56' N	9° 1.76' W	1699.0	SEISREFL	Sono-buoy
PS66/347-1	26.09.04	13:35	75° 36.94' N	10° 8.72' W	1576.0	SEISREFL	Sono-buoy
PS66/347-1	27.09.04	00:07	74° 44.70' N	12° 30.10' W	1983.0	SEISREFL	end of profile
PS66/347-1	27.09.04	00:48	74° 42.27' N	12° 30.61' W	2125.0	SEISREFL	streamer on deck
PS66/347-1	27.09.04	00:55	74° 41.70' N	12° 30.62' W	2167.0	SEISREFL	array on deck
PS66/348-1	27.09.04	01:03	74° 40.77' N	12° 30.64' W	2234.0	M	start- magnetic turncircels
PS66/348-1	27.09.04	02:45	74° 40.60' N	12° 30.56' W	2246.0	М	end-magnetic turncircels

## 12.2 Sea-ice conditions

## V. Shevchenko

Table 12.2: Sea ice conditions in the Yermak Plateau in September 2004

Date of September	Time UTC	Latitude, N	Longitude	True wind speed (m/s)	True wind direction (°)	Air temperature (°C)	Ship speed (kn)	Number of engines	Operation mode: chanel=0,	Total ice concentral	Conc. of thin ice <30 cm (%)	Typical sea ice thickness (m)	Snow thickness (cm)	Typical floe diameter (m)	Max. floe diameter (m)	Melt pond coverage (%)	Typical pond diameter (m)	Maximum pond diameter (m)	Dirty ice concentration (%)	Lead width (m)	Lead floes, diameter (m)	Typical ridge height (m)	Max. ridge height (m)	Typical ridge spacing (m)	Ridges: new=0, old=1, both=2	Rubble fields, coverage (%)	Icebergs, number of
7	9	81°36.45′	4°13.5'E	11.6	315	-6.0	5.8	2	3	10		0.5	5	20	30												
	12	81°36.7'	5°29.3'E	10.8	320	-5.7	3.3	2	3	5		0.5	10	10	20												
	15	81°36.22'	7°12.11′E	12.4	320	-5.0	5.3	2	3	1		0.5	10	10	15												1
	20	81°30.97'	8°04.5'E	6.2	330	-3.7	5.5	2	3	1		0.5	10	10	15												$\Box$
8	8	81°29.82'	1°32.88' E	4.9	240	-6.3	5.3	4	3	95	1	1.5	10	300	500							1.5	2	100	1		$\Box$
	10	81°30.05'	0°38.13' E	6.9	260	-4.2	5.0	4	3	90	5	1	10	15	50							1.0	1.5		1		П
	12	81°29.47'	0°26.60' W	7.1	245	-3.5	5.3	3	3	40		1	15	15	50												П
	16	81°29.93'	2°44.37' W	9.6	270	-3.4	4.3	3	3	50		1	10	15	50							1	1.5		1		П
	18	81°24.00'	2°40.51' W	8.5	280	-3.1	5.8	3	3	10		0.5	5	10	30												П
	20	81°23.94'	1°26.9' W	8.6	275	-2.3	5.7	3	3	50		1	10	20	100							1	1.5		1		П
	22	81°24.09'	0°18.63'W	8.7	280	-3.0	4.8	3	3	5		1	10	15	50										$\vdash$	$\Box$	П

Date of September	Time UTC	Latitude, N	Longitude	True wind speed (m/s)	True wind direction (')	Air temperature (°C)	Ship speed (kn)	Number of engines	Operation mode: chanel=0, lead=1, floe ice=3, ramming=4	Total ice concentration (%)	Conc. of thin ice <30 cm (%)	Typical sea ice thickness (m)	Snow thickness (cm)	Typical floe diameter (m)	Max. floe diameter (m)	Melt pond coverage (%)	Typical pond diameter (m)	Maximum pond diameter (m)	Dirty ice concentration (%)	Lead width (m)	Lead floes, diameter (m)	Typical ridge height (m)	Max. ridge height (m)	Typical ridge spacing (m)	Ridges: new=0, old=1, both=2	Rubble fields, coverage (%)	lcebergs, number of
9	8	81°13.93'	2°39.82' E	7.1	330	-5.6	4.5	3	3	40		1	10	15	30												
	10	81°12.05'	2°41.72' E	9.3	315	-6.6	4.0	3	3	50																	Ш
	13	81°11.87"	2°32.80' E	7.1	300	-6.5	5.9	3	3	30																	Ш
	14	81°11.86'	2°00.05' E	5.5	320	-6.5	6.0	3	3	70	10	1	5	15	50							1	1.5		1		
	16	81°12.00'	0°51.30' E	9.2	350	-6.0	3.3	3	3	40		1.5	10	15	30							1	1.5		1		Ш
	19	81°11.70'	0°48.10' W	8.4	5	-6.6	6.4	3	3	5		1	5	15	30							1	1.5		1		
	20	81°11.95'	1°22.80' W	6.0	340	-6.7	5.2	3	3	5		0.5	5	10	20												
	22	81°09.02"	2°16.42' W	4.7	10	-6.0	5.0	3	3	5		0.5	5	10	20												Ш
10	8	80°58.63'	2°34.80' E	4.0	45	-6.9	5.7	3	3	30		1	5	15	30							1	2		1		Ш
	10	80°53.70'	3°36.00' E	1.3	50	-6.0	5.3	3	3	30		1	10	10	20							1	1.5		1		Ш
	12	80°53.90'	3°22.00' E	6.6	35	-4.9	5.0	3	3	40		1	5	15	30							1	2		1		1
	16	80°54.00'	1°18.80' E	7.9	350	-4.5	6.0	3	3	10		8.0	5	10	25							1	2		1		
	19	80°51.54'	0°00.75' W	15.1	45	-0.7	5.1	3	3	5		0.5	5	10	20							1	2		1		
	21	80°56.00'	0°20.10' E	13.7	40	-1.9	6.0	3	3	5		0.5	10	10	20							1	2		1		

Date of September	Time UTC	Latitude, N	Longitude	True wind speed (m/s)	True wind direction (*)	Air temperature (°C)	Ship speed (kn)	Number of engines	Operation mode: chane=0, lead=1, floe ice=3, ramming=4	Total ice concentration (%)	Conc. of thin ice <30 cm (%)	Typical sea ice thickness (m)	Snow thickness (cm)	Typical floe diameter (m)	Max, floe diameter (m)	Mett pond coverage (%)	Typical pond diameter (m)	Maximum pond diameter (m)	Dirty ice concentration (%)	Lead width (m)	Lead floes, diameter (m)	Typical ridge height (m)	Max. ridge height (m)	Typical ridge spacing (m)	Ridges: new=0, old=1, both=2	Rubble fields, coverage (%)	Icebergs, number of
11	8	81°14.10'	0°13.17' E	10.0	350	-5.8	4.8	3	3	90		1	10	10	25							1	2		1		
	12	80°55.72'	0°59.90' E	10.4	360	-4.7	6.6	3	3	30		0.8	5	10	20							1	1.5		1		
	14	80°45.10'	1°16.13' E	8.7	10	-4.2	6.1	3		1		0.5	5	2	5												П
	16	80°37.00'	1°42.41' E	8.2	5	-4.0	5.6	2		0																	П
	18	80°42.20'	2°46.52' E	7.4	2.0	-4.6	5.9	2		0																	
	20	80°47.14'	3°48.83' E	6.6	35	-3.1	5.7	2		0																	П
	22	80°52.21'	4°53.10' E	6.5	5	-3.9	5.8	2		0																	П
12	8	81°19.35'	10°00.78' E	8.2	330	-3.3	5.8	2		2		0.4	5	5	10												$\square$
	10	81°25.97'	10°59.10' E	8.9	340	-3.3	4.9	2	3	10		0.5	5														$\square$
	12	81°31.62'	11°50.25' E	9.5	350	-3.4	5.9	2	3	50		1	10	10	30							1	1.5	1			$\square$
	14	81°35.82'	12°38.47' E	9.3	355	-3.1	3.5	2	3	10		0.8	5	7	15												$\square$
	16	81°22.92'	12°09.79' E	8.7	10	-2.7	8.2	3	3	2		0.5	5	5	10												$\square$
	20	81°20.17'	12°45.18' E	9.9	350	-2.8	8.0	3	3	5		0.5	2	2	5												

Date of September	Time UTC	Latitude, N	Longitude	True wind speed (m/s)	True wind direction (")	Air temperature (°C)	Ship speed (kn)	Number of engines	Operation mode: chanel=0, lead=1, floe ice=3, ramming=4	Total ice concentration (%)	Conc. of thin ice <30 cm (%)	Typical sea ice thickness (m)	Snow thickness (cm)	Typical floe diameter (m)	Max, floe diameter (m)	Melt pond coverage (%)	Typical pond diameter (m)	Maximum pond diameter (m)	Dirty ice concentration (%)	Lead width (m)	Lead floes, diameter (m)	Typical ridge height (m)	Max. ridge height (m)	Typical ridge spacing (m)	Ridges: new=0, old=1, both=2	Rubble fields, coverage (%)	loebergs, number of
13	8	82°05.55'	18°28.60' E	11.2	30	-1.7	6.8	3	3	85	30	1	5	15	30							1	2.5		1		П
	12	82°10.28'	18°12.28' E	11.9	50	-1.0	0.2	2	3	50	20	8.0	5	15	50							1	1.5		1		П
	14	82°05.36'	18°05.29' E	11.6	55	-2.2	0.2	2	3	50	20	1	5	15	50							1	1.5		1		
	17	82°05.26'	18°03.00' E	8.1	60	-1.5	3.3	2	3	70	30	1	5	20	50							1	2		1		1
	18	82°06.26'	18°55.53' E	8.7	60	-1.6	8.7	3	3	90	30	1	5	15	50							1	1.5		1		
	20	82°11.04'	20°24.74' E	10.6	60	-1.0	6.6	3	3	90	30	1	5	20	50							1	1.5		1		
14	8	82°14.36'	23°01.15' E	8.6	35	-2.1	5.4	3	3	95	30	1.2	5	50	300	5	5	10				1	2		1		
	12	81°52.53'	23°23.00' E	8.8	20	-2.3	9.7	3	3	30		8.0	5	15	60	5	5	10									
	14	81°38.92'	24°29.98' E	8.5	20	-1.7	10.4	3	3	2		0.5	0.5	1	5												
	16	81°27.85'	25°00.08' E	8.5	25	-1.2	5.7	2		0																	
15	8	81°20.20'	25°30.12' E	7.3	5	-3.0	5.4	2		0																	
	12	81°24.01'	26°02.20' E	8.7	5	-3.9	6.3	2		0																	
	16	81°26.17'	25°59.93' E	6.1	350	-4.9	0.1	2		0																	
	22	81°31.95'	22°43.40' E	5.5	320	-5.5	10.2	3		0																	

Date of September	Time UTC	Latitude, N	Longitude	True wind speed (m/s)	True wind direction (°)	Air temperature (°C)	Ship speed (kn)	Number of engines	Operation mode: chanel=0, lead=1, floe ice=3, ramming=4	Total ice concentr	Conc. of thin ice <30 cm (%)	Typical sea ice thickness (m)	Snow thickness (cm)	Typical floe diameter (m)	Max. floe diameter (m)	Melt pond coverage (%)	Typical pond diameter (m)	Maximum pond diameter (m)	Dirty ice concentration (%)	Lead width (m)	Lead floes, diameter (m)	Typical ridge height (m)	Max. ridge height (m)	Typical ridge spacing (m)	Ridges: new=0, old=1, both=2	Rubble fields, coverage (%)	Icebergs, number of
16	8	81°25.88'	15°18.67' E	3.7	335	-4.3	10.5	3		0																	
	16	80°43.23'	15°52.85' E	2.2	210	-3.4	0.6	3		0																	Ш
17	8	81°27.06'	13°07.48' E	10.0	180	-0.8	3.3	3		0																	
	12	81°24.00'	10°58.93' E	7.9	230	-0.2	5.9	3		0																	
	13	81°24.38'	10°22.16' E	5.8	225	-0.7	5.6	ფ	3	20		0.5	5	5	30							1	1.5		1		
	14	81°24.02'	9°47.00′ E	8.7	230	-0.3	4.5	3	3	40		0.5	10	10	50												
	16	81°24.00'	8°37.15′ E	6.3	245	-1.4	5.4	3	3	40		1.2	10	20	50							1	2		1		
	18	81°24.00'	7°20.60' E	7.0	185	-1.6	6.1	3		0																	
	20	81°23.97'	6°02.95′ E	9.7	195	-0.1	6.1	3		0																	
18	8	81°17.96'	10°48.55' E	7.1	220	+0.2	5.8	3		0																	
	10	81°17.94'	9°29.20' E	5.9	170	-0.6	5.3	ფ	3	90		1	10	20	50							1	3		1		
	12	81°18.0'	8°22.80' E	6.5	190	-0.7	5.6	ფ	3	50	20	1	10	20	50							1.5	3		1		
	14	81°17.92'	7°15.75' E	6.4	175	-0.6	4.3	3	3	20		8.0	10	10	25							1	1.5		1		
	16	81°17.98'	6°01.00′ E	8.4	170	-0.2	5.3	3	3	0																	

Date of September	Time UTC	Latitude, N	Longitude	True wind speed (m/s)	True wind direction (°)	Air temperature (°C)	Ship speed (kn)	Number of engines	Operation mode: chanel=0, lead=1, floe ice=3, ramming=4	Total ice concentration (%)	Conc. of thin ice <30 cm (%)	Typical sea ice thickness (m)	Snow thickness (cm)	Typical floe diameter (m)	Max. floe diameter (m)	Melt pond coverage (%)	Typical pond diameter (m)	Maximum pond diameter (m)	Dirty ice concentration (%)	Lead width (m)	Lead floes, diameter (m)	Typical ridge height (m)	Max. ridge height (m)	Typical ridge spacing (m)	Ridges: new=0, old=1, both=2	Rubble fields, coverage (%)	Icebergs, number of
18	18	81°18.00'	4°50.20' E	9.6	160	-0.1	5.7	3		0																	Ш
	20	81°17.89'	3°40.75' E	5.3	165	-0.4	6.3	3	3	50		1	15	15	50							1	1.5		1		
19	8	82°06.00'	4°16.83' E	7.2	160	-1.2	4.8	4	3	90		1.2	15	20	50							1	1.5		1		
	14	82°00.94'	7°16.00' E	8.3	145	-0.6	8.8	3	3	40		1	10	15	30				1			1.5	3		1		
	16	81°54.85'	8°34.00′ E	5.9	175	+0.1	5.1	3	3	80	30	0.6	10	15	30							1	2		1		П
	18	81°55.70'	8°32.74' E	4.5	160	-0.5	1.7	3	3	50	10	0.5	10	15	40							1.5	3		1		П
	20	81°57.08'	8°19.16′ E	7.2	175	-0.6	0.2	3	3	75	20	8.0	10	15	30												$\Box$
20	8	81°36.38'	6°47.64' E	6.6	60	-1.6	0.6	3		0																	$\Box$
	12	81°34.10'	6°28.10′ E	7.7	65	-0.7	10.0	3		<1		0.5	5	5	10												$\Box$
	14	81°25.91'	4°43.98′ E	7.1	65	-1.0	10.0	3	3	5		0.5	10	2	10												$\Box$
	16	81°19.54'	3°09.10' E	6.4	35	-1.6	8.0	3	3	25		1	15	5	20				1			1	3		1		一
	18	81°18.07'	2°06.80' E	6.0	50	-1.9	5.8	3	3	50		1	15	15	40							1	2		1		$\square$
	20	81°18.37'	0°58.63' E	5.5	15	-5.0	4.2	3	3	95		1.5	15	20	70							2	5		2		

Date of September	Time UTC	Latřiude, N	Longitude	True wind speed (m/s)	True wind direction (°)	Air temperature (°C)	Ship speed (kn)	Number of engines	Operation mode: chanel=0, lead=1, floe ice=3, ramming=4	Total ice concentration (%)	Conc. of thin ice <30 cm (%)	Typical sea ice thickness (m)	Snow thickness (cm)	Typical floe diameter (m)	Max. floe diameter (m)	Melt pond coverage (%)	Typical pond diameter (m)	Maximum pond diameter (m)	Dirty ice concentration (%)	Lead width (m)	Lead floes, diameter (m)	Typical ridge height (m)	Max. ridge height (m)	Typical ridge spacing (m)	Ridges: new=0, old=1, both=2	Rubble fields, coverage (%)	Icebergs, number of
21	8	81°05.70'	2°00.60' E	6.7	355	-4.9	5.3	3	3	95	20	1	10	10	25							1	2		1		
	12	81°06.04'	4°25.56' E	8.1	15	-2.4	5.4	3		0																	$\Box$
22	13	79°42.18'	0°47.94' W	5.9	10	-1.7	4.6	3		0																	3
23	7.5	79°12.10'	6°11.69' W	3.1	120	-4.4	10.0	2	3	50	40	0.5	5	2	5												3
	8	79°11.19'	6°44.43' W	5.0	125	-4.1	10.4	2		0																	1
	11	79°03.00'	9°14.80' W	5.0	230	-6.9	9.9	2		0																	2
	12	79°00.32'	10°12.22' W	5.7	280	-6.5	10.1	2		0																	2
	13	79°02.67'	10°55.56' W	5.7	280	-6.5	10.1	2		0																	16
	19	78°58.00'	11°23.46' W	3.6	215	-8.5	6.0	2		0																	9
24	8	79°03.50'	11°16.10' W	5.7	180	-3.4	0.2	2		0																	13

#### **Comments to Table 12.2**

07.09.2004

20:00. POLARSTERN goes in open water along the ice margin which is situated at the distance about 1 mile from the ship.

09.09.2004.

14.00. Some ice floes are slightly brownish at their bottom – traces of algal

13.09.2004

08:00. Strong algal bloom (bottoms of ice floes are brown).

14:00. Strong algal bloom at the bottom of ice floes.

20:00. Strong algal bloom at the bottom of ice floes.

14.09.2004

08.00. Strong algal bloom at the bottom of ice floes.

14:00. In small ice floe (5 m) there is log (wood) 1.5 long.

16.09.2004.

08:00. Rare small peaces of rotten ice (1–2 m).

17.09.2004.

16:00. Weak algal bloom at the bottom of ice floes.

19.09.2004.

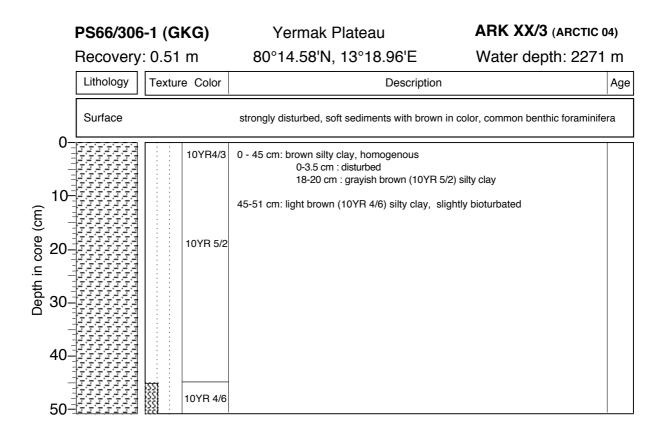
08:00. Weak algal bloom at the bottom of ice floes.

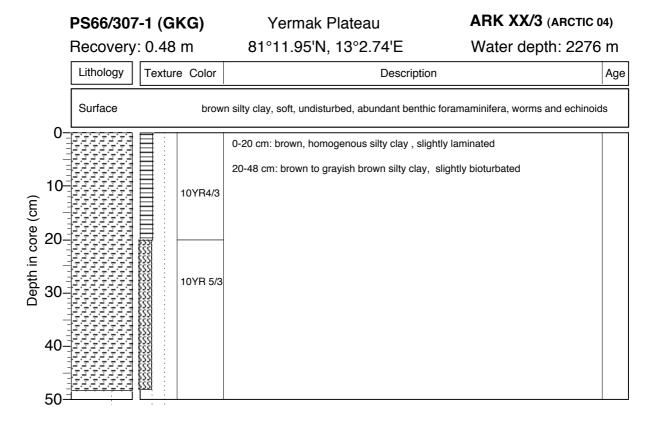
14: 00. Weak algal bloom at the bottom of ice floes.

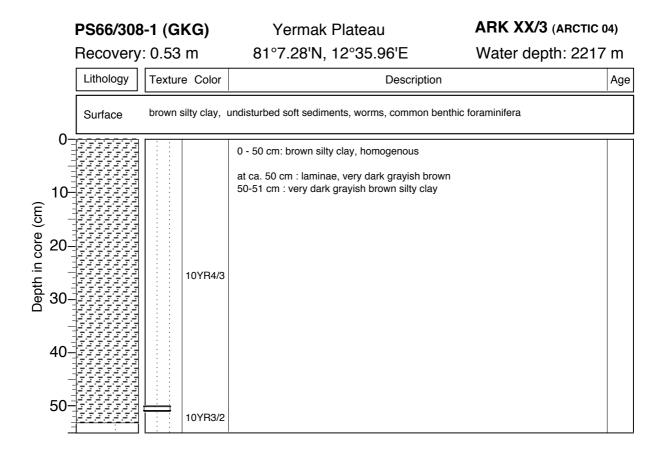
20.09.2004.

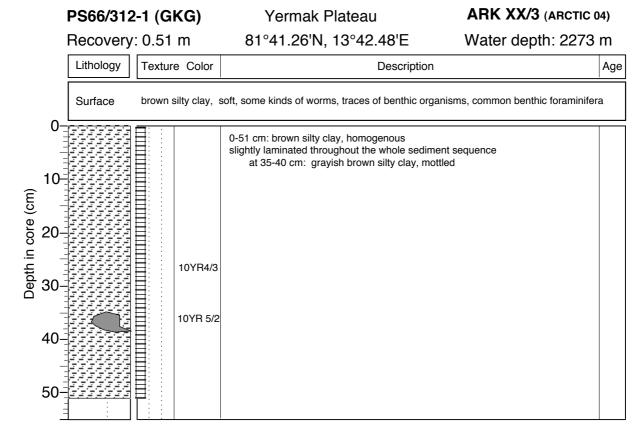
14:15. Ice 40%, 10% are sediment-laden ("dirty"), but the sizes of ice floes are small (5–15 m) and landing of helicopters on them is impossible.

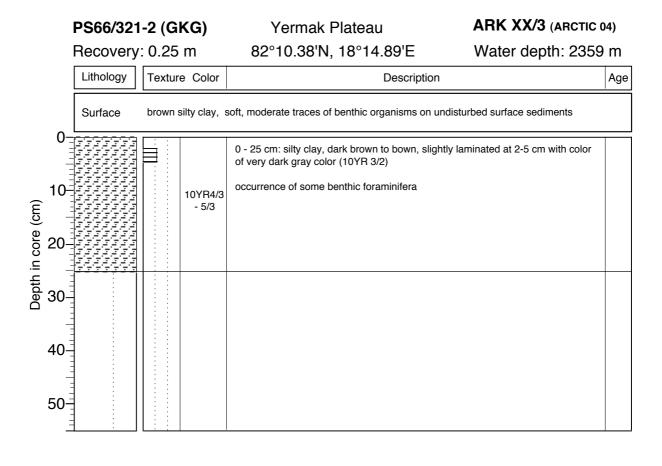
## **12.3.1** Graphical Core Description of Giant box cores S. Nam

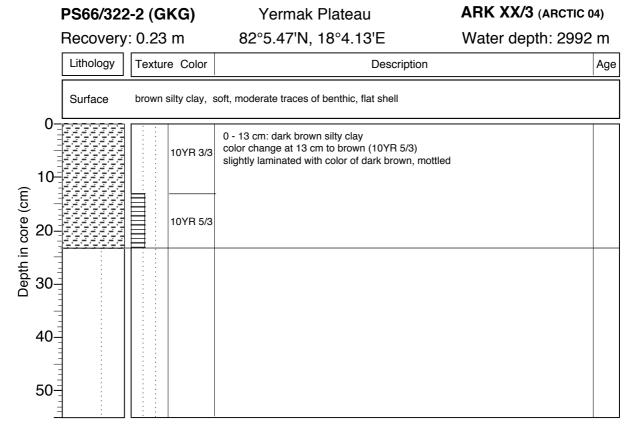


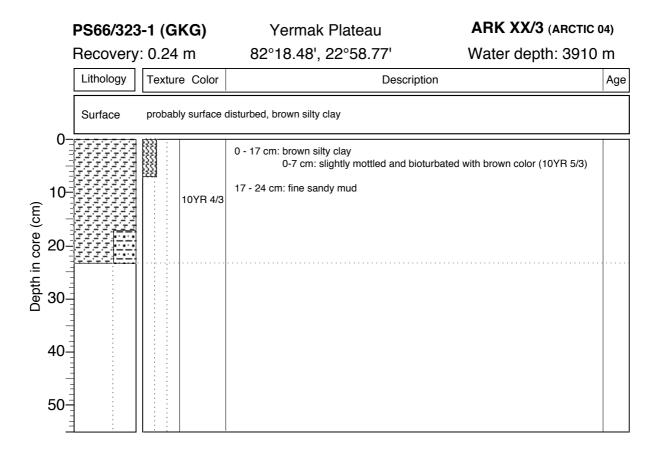


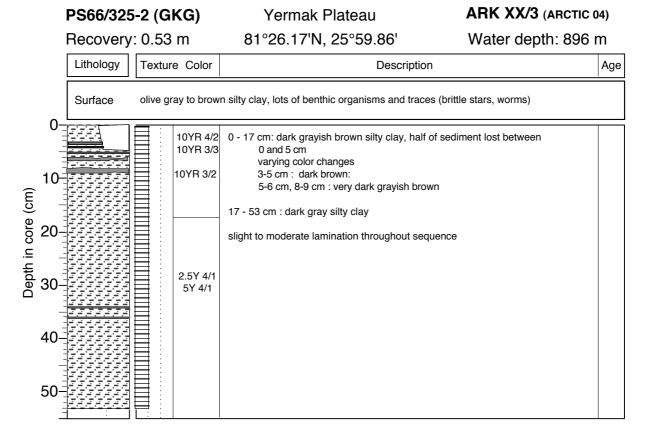


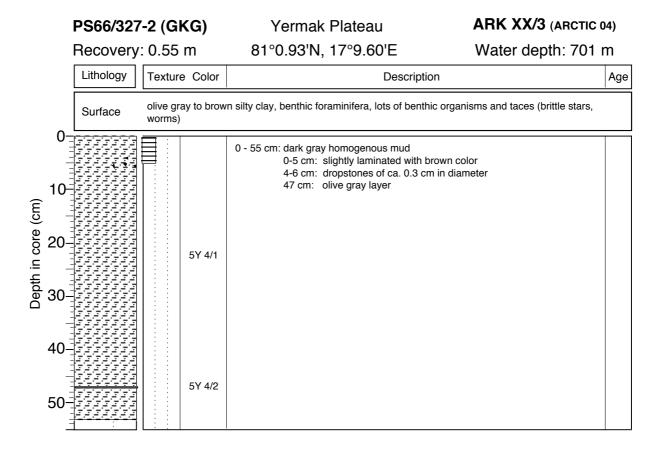


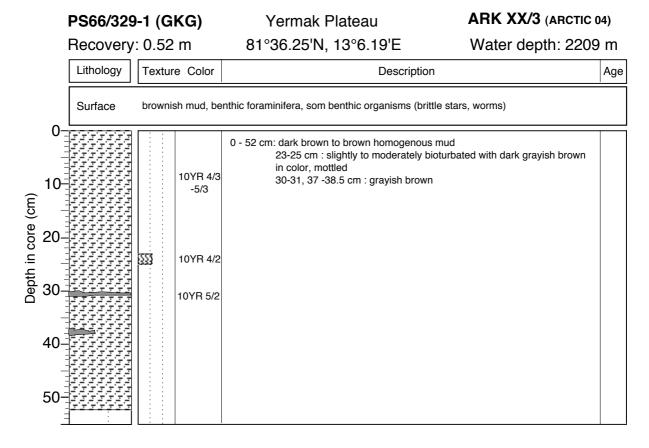


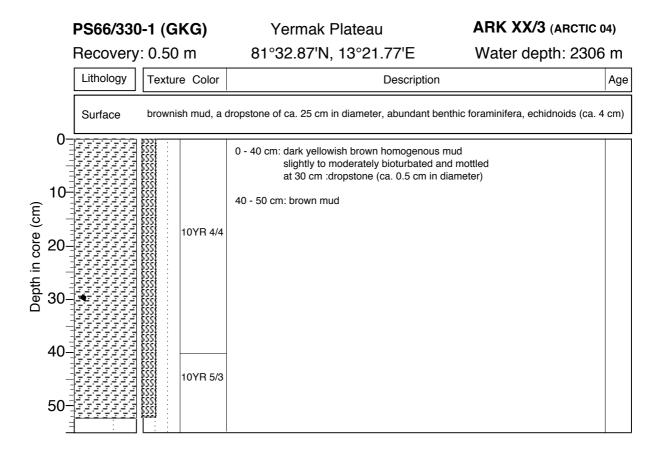


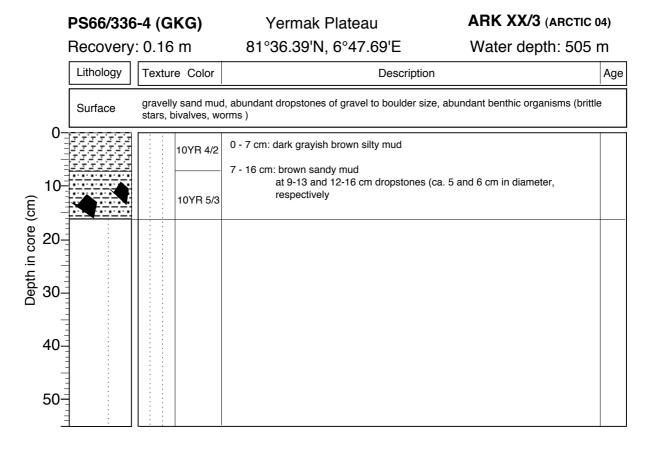






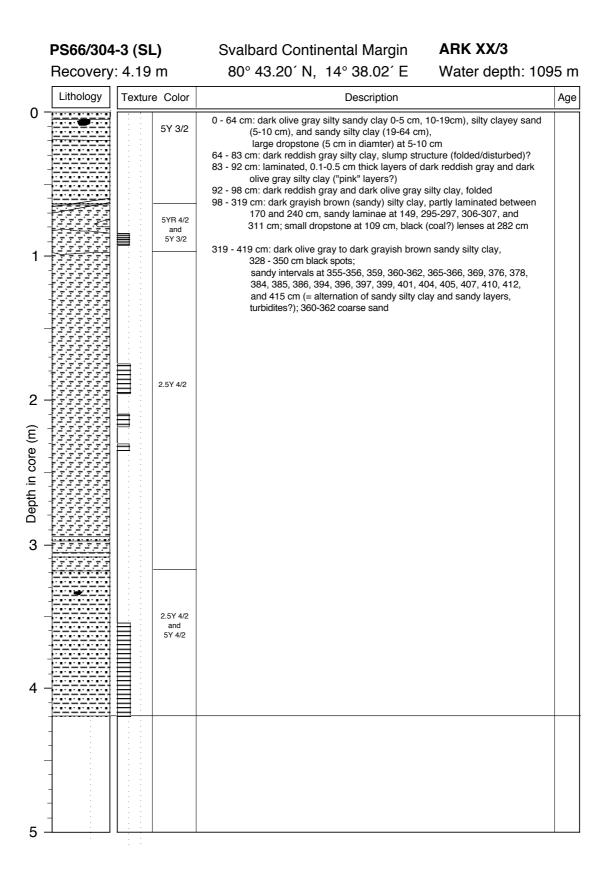






## 12.3.2 Graphical Core Description – Gravity Cores (SL) and Kastenlot Cores (KAL)

R. Stein



#### PS66/306-2

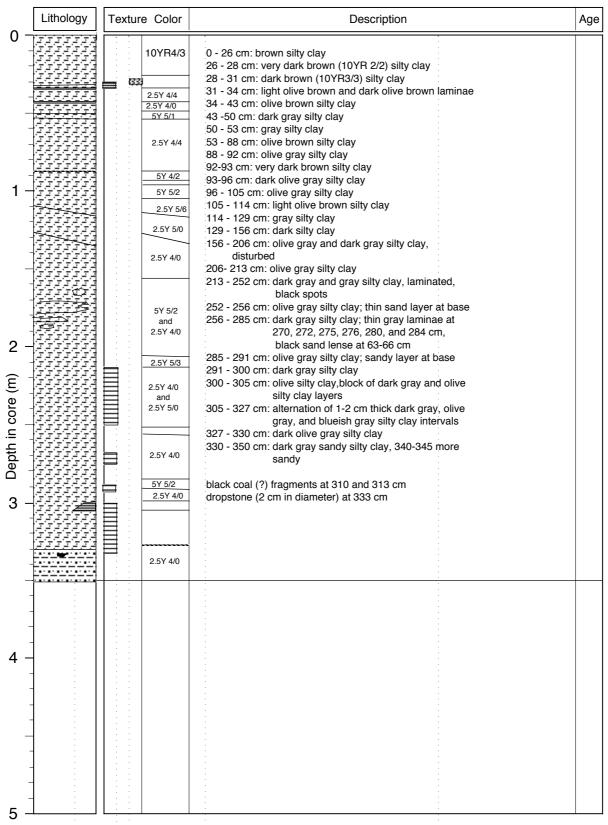
## Svalbard Continental Margin

### ARK XX/3

Recovery: 3.50 m

81° 14.56′ N, 13° 18.61′ E

Water depth: 2210 m



#### PS66/307-2 (SL)

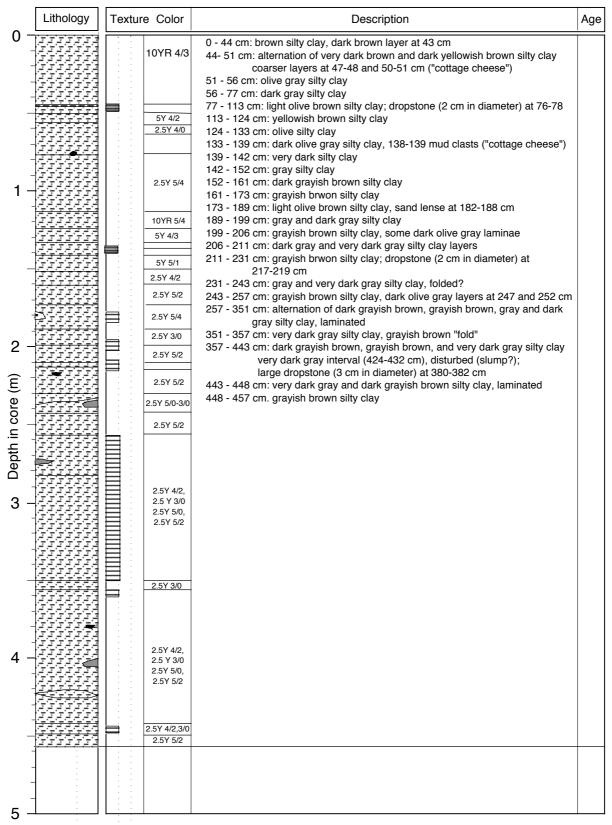
## Svalbard Continental Margin

ARK XX/3

Recovery: 4.57 m

81° 11.94′ N, 13° 02.68′ E

Water depth: 2275 m



#### PS66/308-3 SL

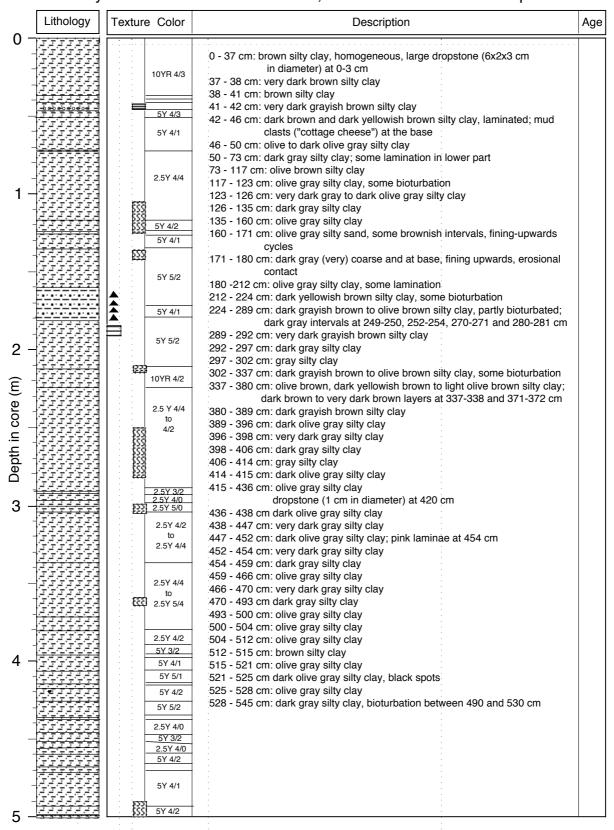
## Yermak Plateau

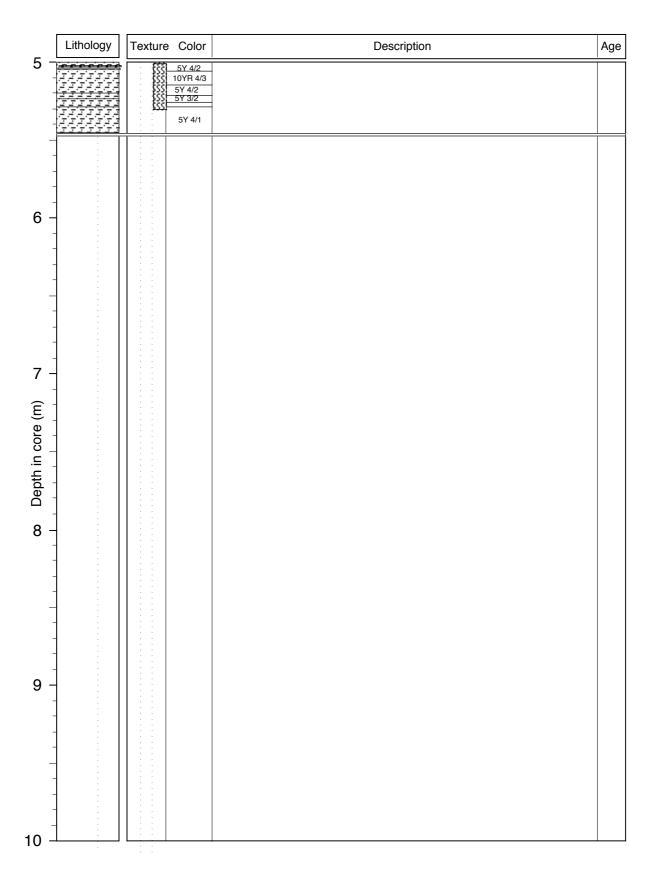
#### ARK XX/3

Recovery: 5.54 m

81° 07.30′ N, 12° 35.97′ E

Water depth: 2218 m





#### PS66/309-1 KAL

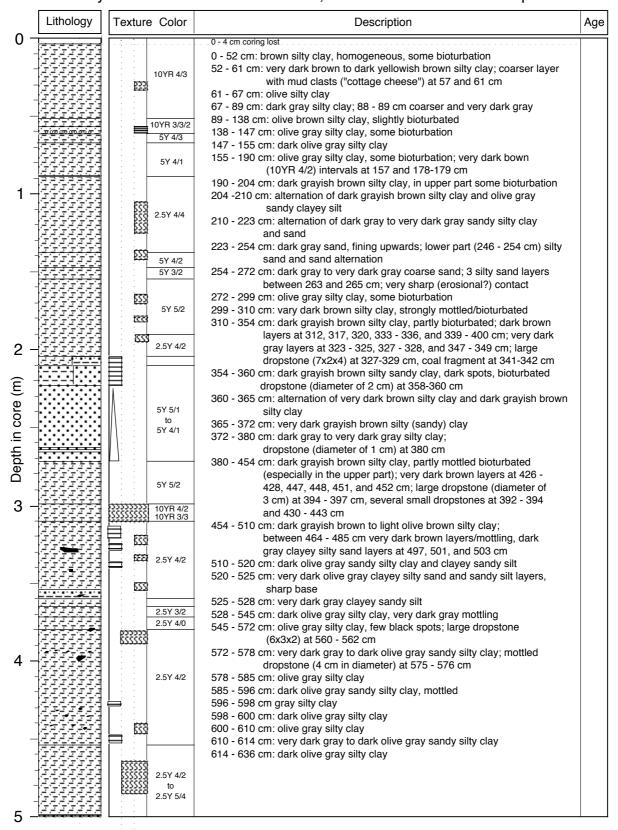
## Yermak Plateau

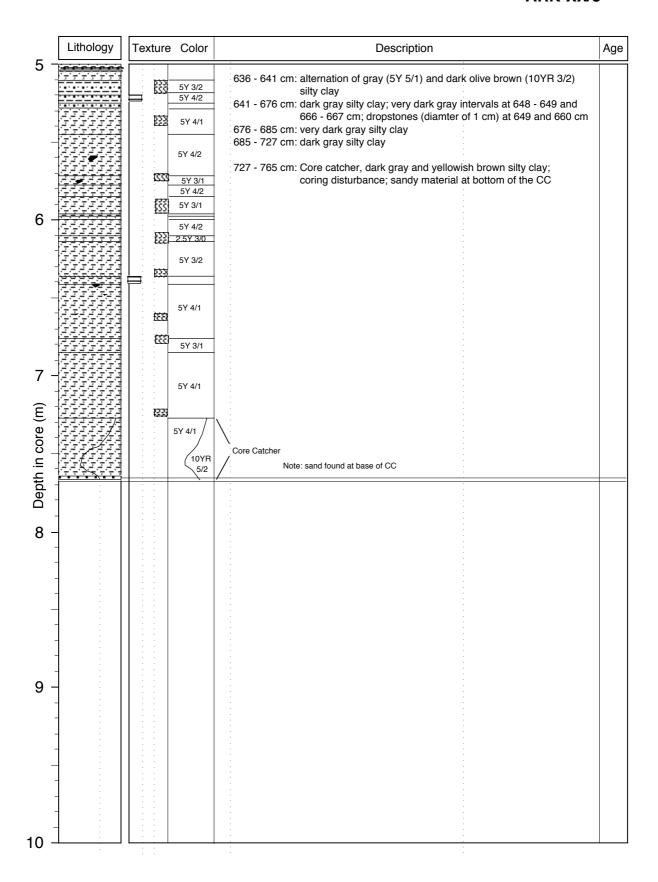
#### ARK XX/3

Recovery: 7.65 m

81° 11.22′ N, 12° 59.07′ E

Water depth: 2270 m





## PS66/311-3 (SL)

## Yermak Plateau

## ARK-XX/3

Recovery: 4.93 m

81° 41.78′ N, 13° 28.20′ E Water depth: 2192 m

	Lithology	Texture Color	Description	Age
0 -		: : 10YR4/	0-25 cm: brown silty clay 25-27 cm: very dark brown silty clay, lamination	
		10YR3/2 -5/4	27-34 cm: yellowish brown and dark brown silty clay, coarser darker layers at 28 and 31 cm ("cottage cheese")	
-		2.5Y4/0	34-68 cm: dark gray silty clay 68-73 cm: gray silty clay 73-103 cm: grayish brown silty clay	
	-19999999 - <del>199999</del> 9	2.5Y5/0	103-112 cm: dark olive gray silty clay, small dropstone (0.5 cm) at base 112-118 cm: gray silty clay	
		2.5Y5/2	118-125 cm: (light) gray silty clay 125-189 cm: grayish brown silty clay, brownish layer at 153-154 cm 189-200 cm: gray silty clay	
1 -		5Y3/2 2.5Y5/0	200-204 cm: gray silty clay (brown at top and base) 204-220 cm: gray silty clay; small dropstone at 204 cm	
		2.5Y6/0	220-245 cm: grayish brown silty clay 245-318 cm: (dark) grayishbrown silty clay;	
_		555 555 555 555 555 555 555 555 555 55	gray and dark gray silty clay lense at 257-259 cm (burrow?); small mud clasts at 268 and 272 cm; brownish layers at	
		2.5Y5/2	318-334 cm: dark gray silty clay 334-346 cm: gray silty clay	
			346-354 cm: light gray silty clay 354-393 cm: grayish brown silty clay; brownish layers at 372, 375, and	
2 -		2.5Y5/0 2.5Y5/2	378 cm	
<u> </u>		2.5Y5/0	400 4FF am (doub) gravial brown although a ma history bation	
ore (m		2.5Y5/2		
ω Depth in core (m)		2.5Y4/2		
		2.5Y4/0		
		2.5Y5/0		
- - -		2.5Y6/0 2.5Y5/2		
4 -		5Y4/2		
_		\$55 2.5Y4/2		
· ·		5Y4/2		
5 -	End of core			

#### PS66/312-2 SL

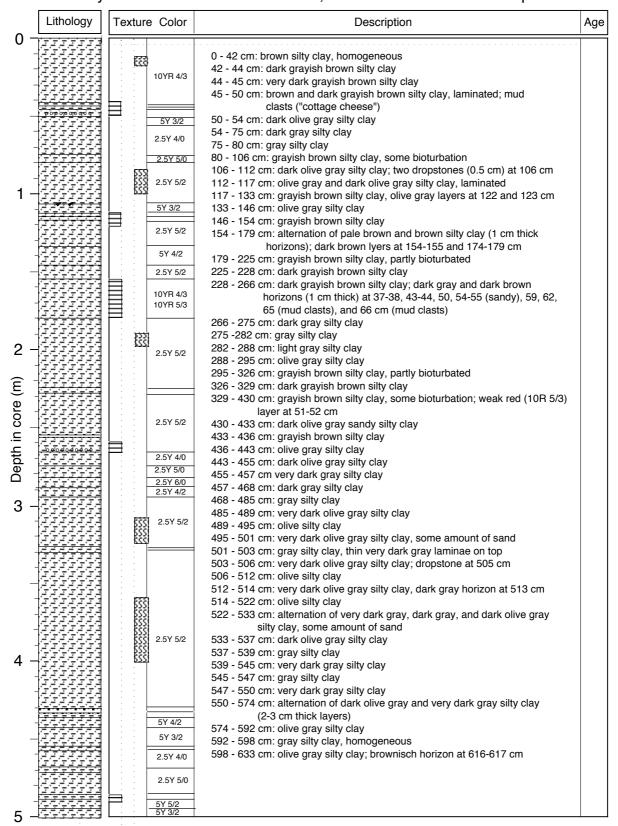
## Yermak Plateau

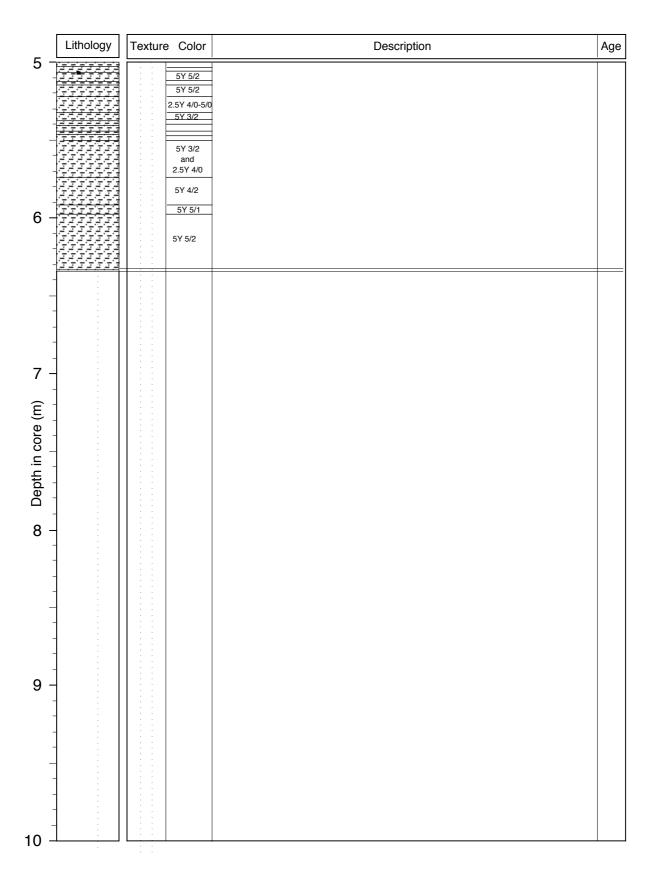
#### ARK XX/3

Recovery: 6.33 m

81° 41.24′ N, 13° 42.65′ E

Water depth: 2275 m





#### PS66/313-1

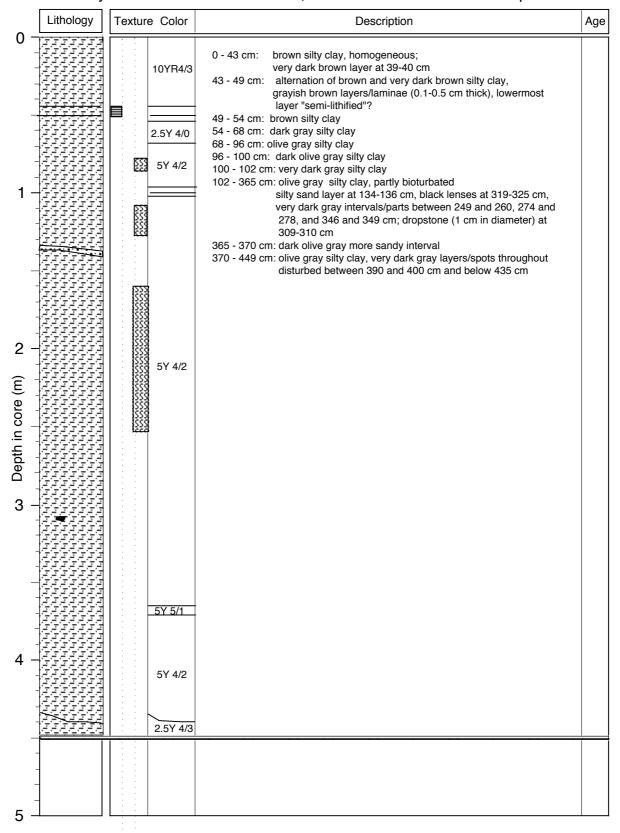
## Yermak Plateau

#### ARK XX/3

Recovery: 4.49 m

81° 45.56′ N, 14° 16.31′ W

Water depth: 2298 m



#### PS66/318-4

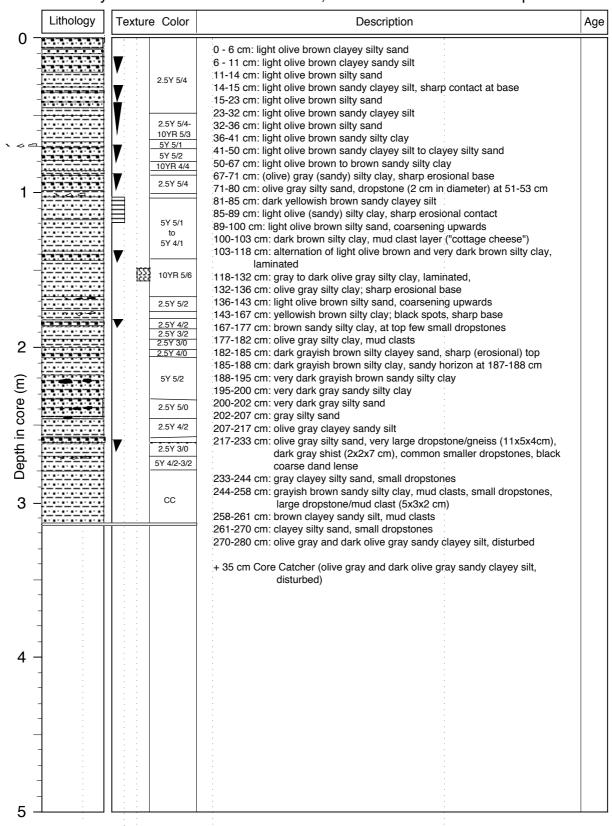
Yermak Plateau

#### ARK XX/3

Recovery: 3.15 m

81° 18.60′ N, 12° 02.71′ E

Water depth: 1262 m



#### PS66/319-1

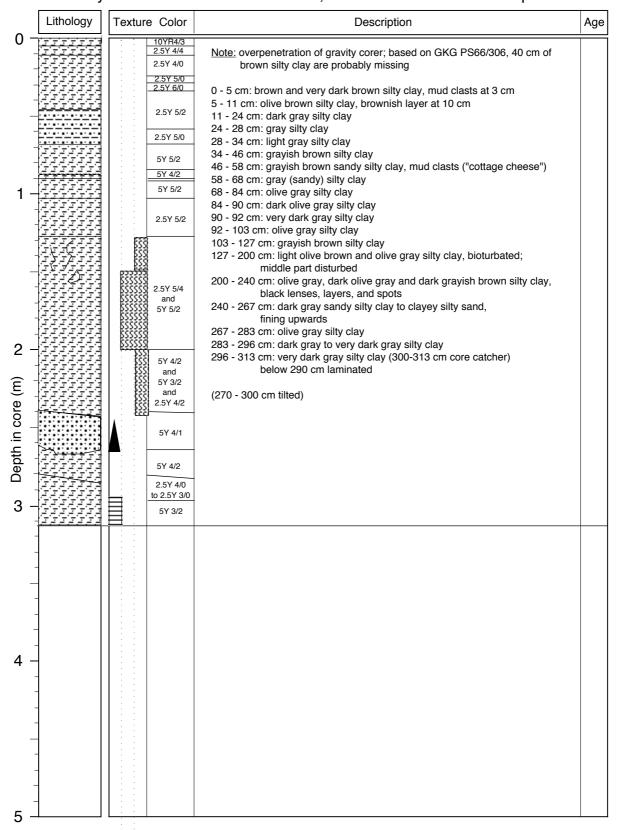
## Yermak Plateau

#### ARK XX/3

Recovery: 3.13 m

81° 24.49′ N, 13° 48.79′ E

Water depth: 2178 m



#### PS66/321-4 KAL

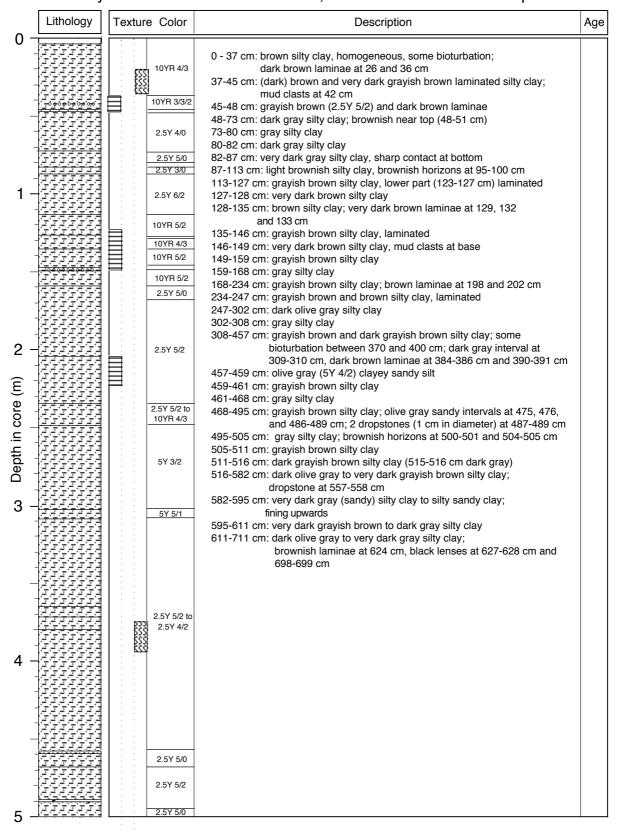
## Yermak Plateau

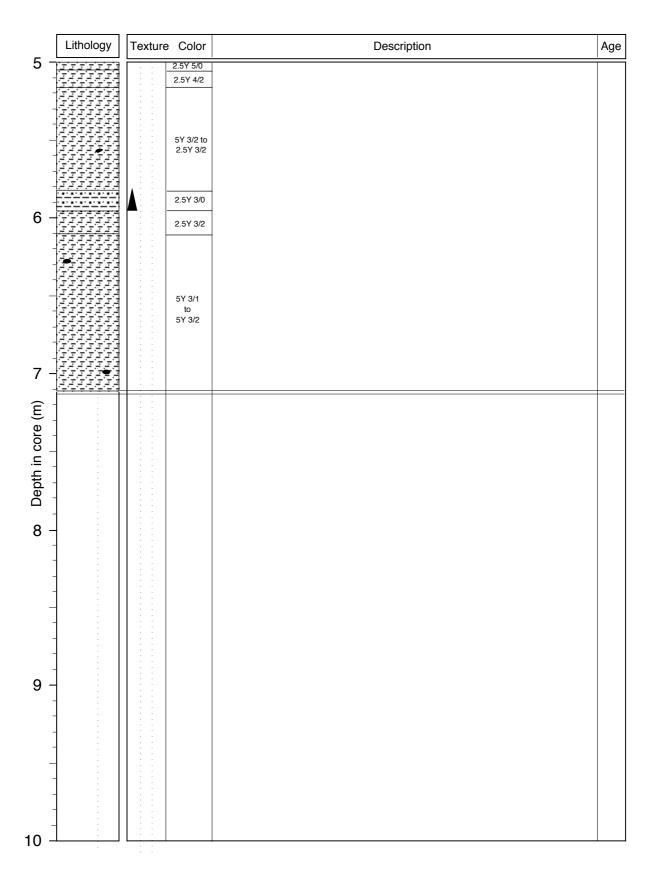
#### ARK XX/3

Recovery: 7.11 m

82° 10.27′ N, 18° 13.47′ E

Water depth: 2360 m



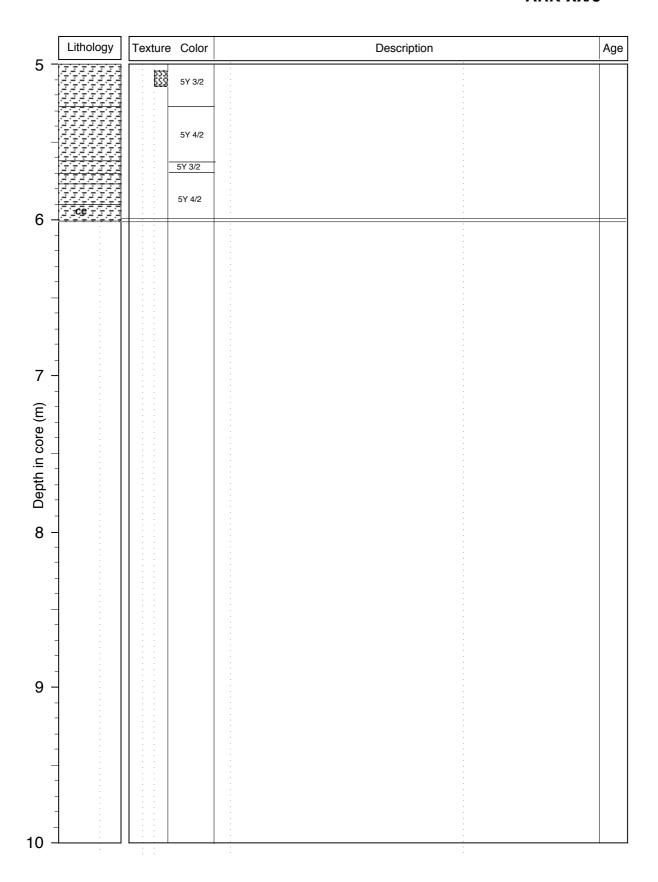


## PS66/322-3 SL

## Yermak Plateau ARK XX/3

Recovery: 6.00 m 82° 05.40′ N, 18° 05.25′ E Water depth: 3028 m

_	Lithology	Texture Color	Description	Age
0 -	777777 777777	10YR 4/3	0-14 cm: brown silty clay, some sand	
		10YR 3/2		
		10YR 4/2		
_			42-60 cm: dark olive gray silty clay	
		5Y 3/2	60-66 cm: dark olive gray silty sandy clay 66-80 cm: dark olive gray silty clay	
	777777		80-86 cm: dark gray silty clay	
		2.5Y 4/0	86-101 cm: gray silty clay 101-107 cm: olive gray silty clay, some sand	
	- 7 7 7 7 7 7 7	2.5Y 5/0	107-138 cm: dark olive gray (sandy) silty clay to silty sandy clay,	
1 -		5Y 4/2	fining upwards 138-145 cm: dark olive clay silty clay; sharp base and top	
		5Y 3/2	145-147 cm: olive gray silty clay	
			147-149 cm: olive gray silty sandy clay	
			149-154 cm: olive gray silty clay 154-161 cm: olive gray silty sandy clay	
_			161-172 cm: olive gray silty clay; brownish interval at 170-172 cm	
		5Y 4/2	172-179 cm: dark gray silty clay 179-184 cm: dark olive gray silty clay to clayey sandy silt, fining upwards	
	1333333	2.5Y 4/0		
		<b>5</b> Y 3/2	186-188 cm: dark olive gray silty clay	
2 -			188-200 cm: olive gray sand 200-207 cm: olive gray silty sand	
_		5Y 4/2	207-222 cm: olive gray silty clay; brownish laminae at 219 and 220 cm	
Ē	-333333		222-226 cm: dark gray silty clay 226-232 cm: (dark) gray silty clay	
<u>_</u>		2.5Y 4/0	232-269 cm: olive gray silty clay; more sandy layer at 249 cm	
Depth in core (m)	1999999		269-273 cm: dark olive gray silty sand	
_ ⊑	333333	5Y 4/2	273-280 cm: dark olive gray silty clay 280-300 cm: dark gray silty clay	
Ę	333333		300-309 cm: olive gray silty sandy clay	
eb(	44444	5Y 3/2	309-319 cm: olive gray silty clay 319-322 cm: very dark gray silty clay	
ш		5Y 4/1	322-329 cm: very dark gray clayey sandy silt	
3 -	3333333		329-338 cm: olive gray silty clay 338-340 cm: dark gray silty clay	
		5Y 4/2	340-350 cm: dark gray slity clay	
		5Y 3/1	330-365 cm: olive gray silty clay, sharp (erosional?) base	
		5Y 4/2	365-400 cm: dark olive gray sandy silty clay 400-405 cm: dark olive gray sandy silty clay	
_		5Y 4/1	405-438 cm: very dark gray silty clay	
		5Y 4/2	438-453 cm: dark gray silty clay; brownish laminae at 444 and 449 cm 453-480 cm: dark olive gray silty clay	
			480-496 cm: olive gray silty clay	
		5Y 3/2	496-500 cm: dark gray silty clay 500-527 cm: dark olive gray silty clay	
4			527-563 cm: olive gray silty clay	
4 -			563-569 cm: dark olive gray silty clay	
		5Y 3/1	569-600 cm: olive gray silty clay; dark gray layer at 577 cm	
		57/4/		
-		5Y 4/1	_	
	1999999	5Y 4/2		
	1999999			
		5Y 5/2	1	
5 -		5Y 4/1	-	

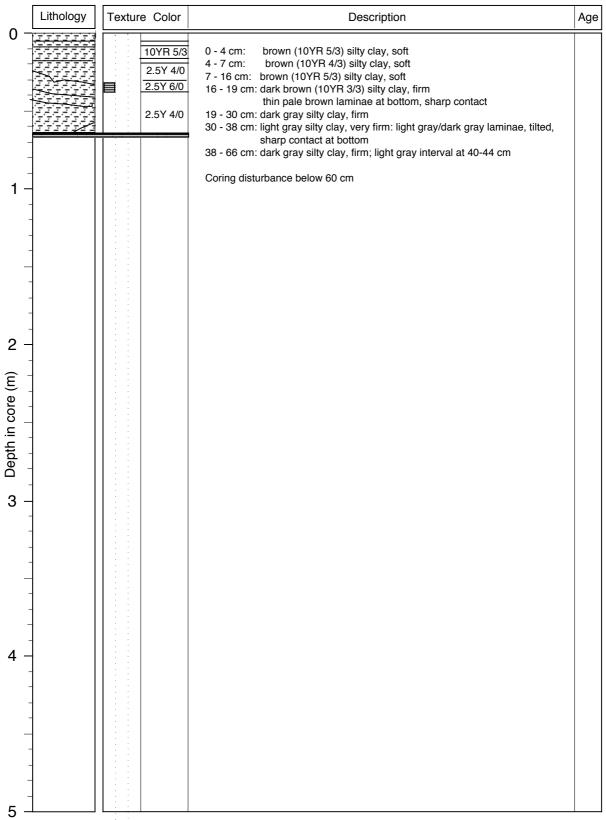


### PS66/323-3

## Yermak Plateau

### ARK XX/3





#### PS66/325-3 SL

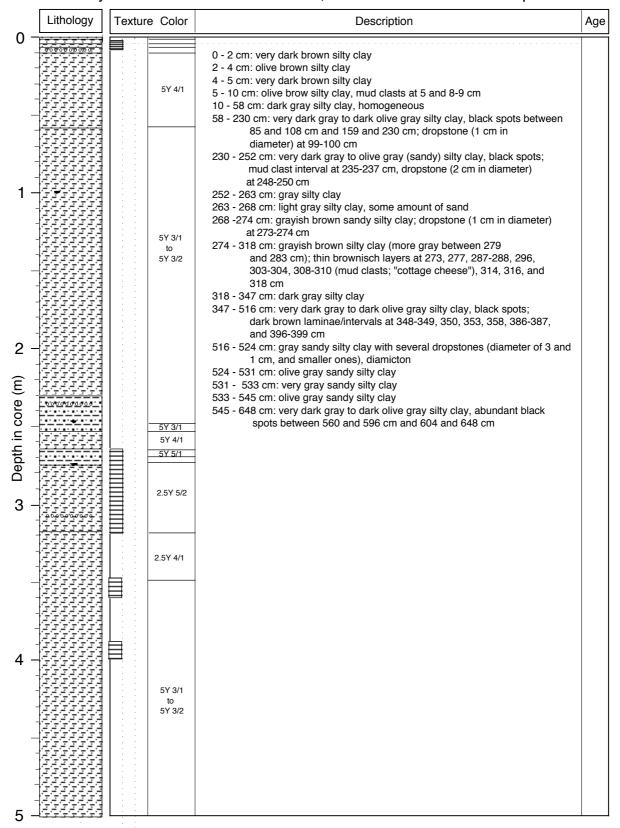
## Svalbard Continental Margin

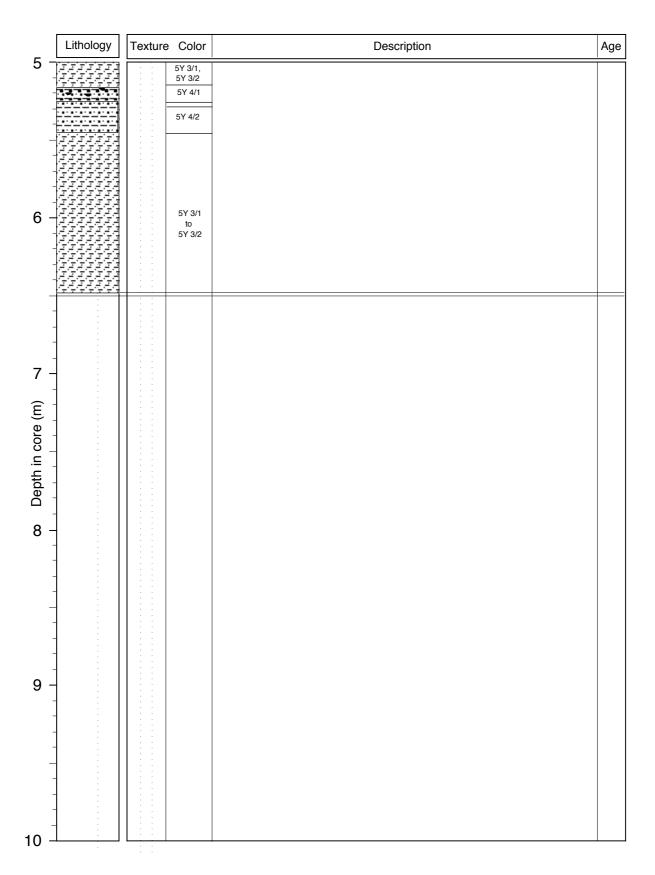
ARK XX/3

Recovery: 6.48 m

81° 26.18′ N, 25° 59.86′ E

Water depth: 896 m





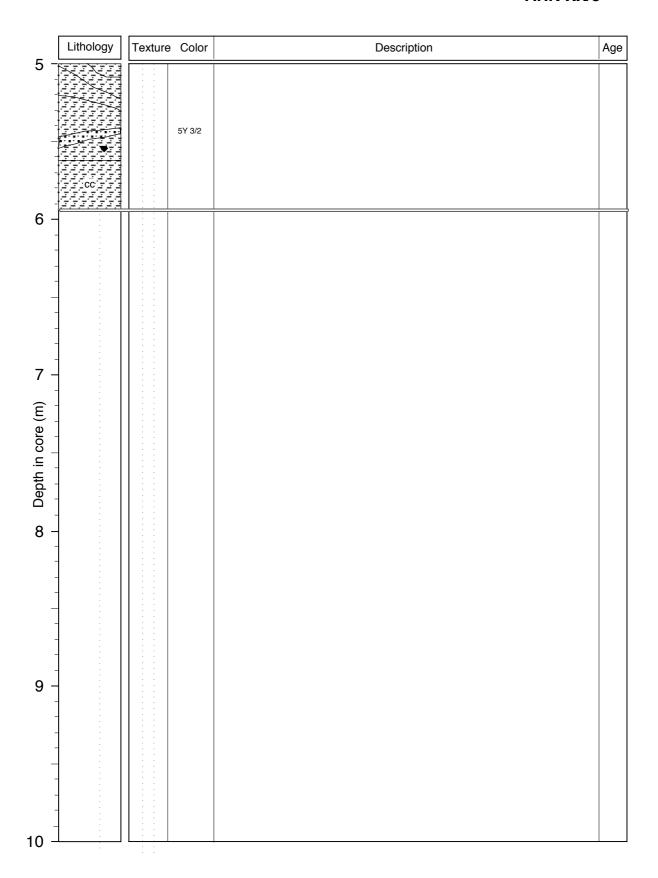
## **PS66/327-3 KAL** Recovery: 5.92 m

# Svalbard Continental Margin 81° 00.92′ N, 17° 09.39′ E

## ARK XX/3

E Water depth: 701 m

	Lithology	Texture Color	Description	
0 -	Lithology	Texture Color 5Y 5/2 5Y 4/2	Description  0 - 22 cm: dark olive gray sandy silty clay, upper part disturbed 22 - 56 cm: dark olive gray silty clay, brownish horizons at 38-53 cm 56 - 200 cm: olive gray silty clay, brownish intervals of few cm in thickness, faulted 200 - 300 cm: dominantly silty clay with sandy silty clay intervals, colour changes from olive gray, dark gray, very dark gray, brown, to brownish gray; interval between 265 and 285 cm folded?; small dropstones at 210-240, 272-278, and 291-297 cm 300 - 400 cm: dark olive gray silty clay, brownisch intervals (tilted/faulted) 400 - 500 cm: olive gray, brown, dark gray, dark grayish brown (sandy) silty clay, faulted/tilted; large dropstones at 407-418 (2x6x10 in diameter), 474-478 (3 cm in diameter), 478-480 (1x3x5), and 480-482 cm (8x2x3) 500 - 562 cm: dark olive gray silty clay, brown horizons, faulted; clayey sandy silt at 543-552 cm; dropstone (3 cm in diameter) at 555-557 cm 562 - 592 cm: core catcher (dark olive gray silty clay)	Age
ω Depth in core (m) ω		5Y 4/2, 10YR 4/3 2.5Y 4/0 2.5Y 3/0	302 - 392 cm. Core calcrer (dark onve gray silly day)	
- - 4 -		5Y 3/2		
5 -		5Y 3/2, 10YR 4/2 2.5Y 4/2		



### PS66/329-3 (SL)

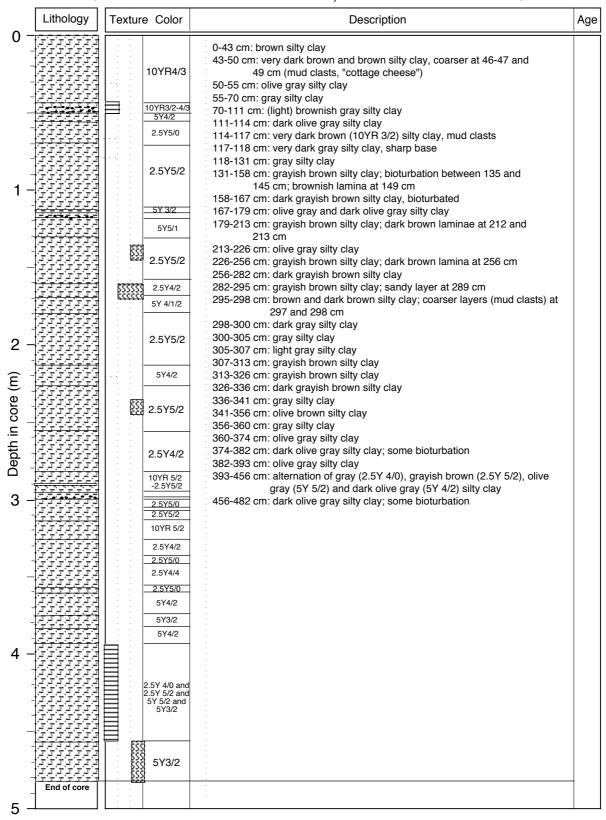
### Yermak Plateau

#### ARK-XX/3

Recovery: 4.82 m

81° 36.31′ N, 13° 06.13′ E

Water depth: 2211 m



### PS66/330-2

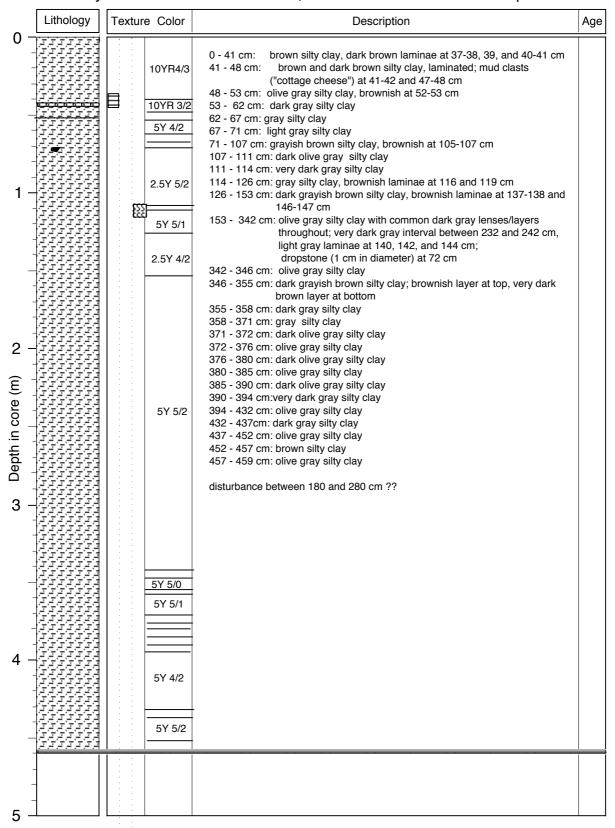
## Yermak Plateau

### ARK XX/3

Recovery: 4.54 m

81° 32.88′ N, 13° 21.71′ W

Water depth: 2305 m



### PS66/333-4

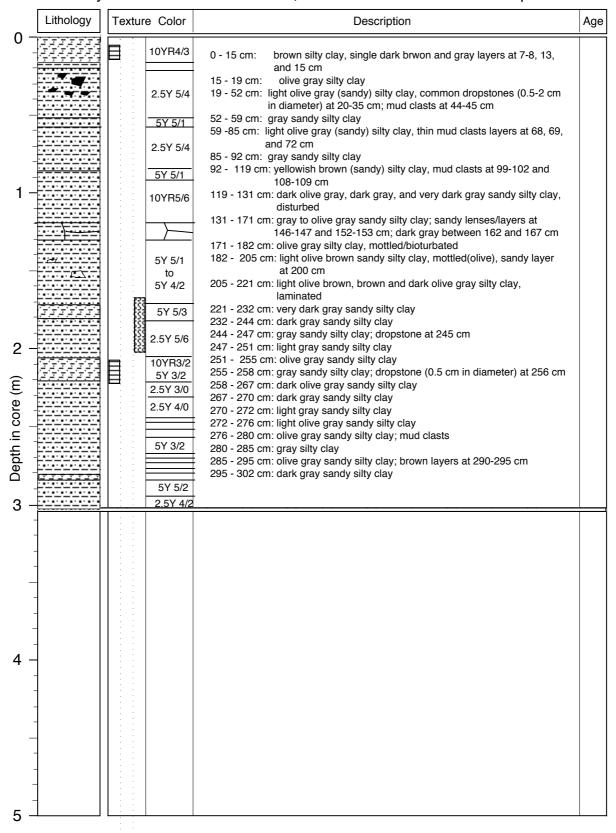
## Yermak Plateau

### ARK XX/3

Recovery: 3.02 m

81° 56.42 N, 08° 30.17′ W

Water depth: 726 m



### PS66/333-5

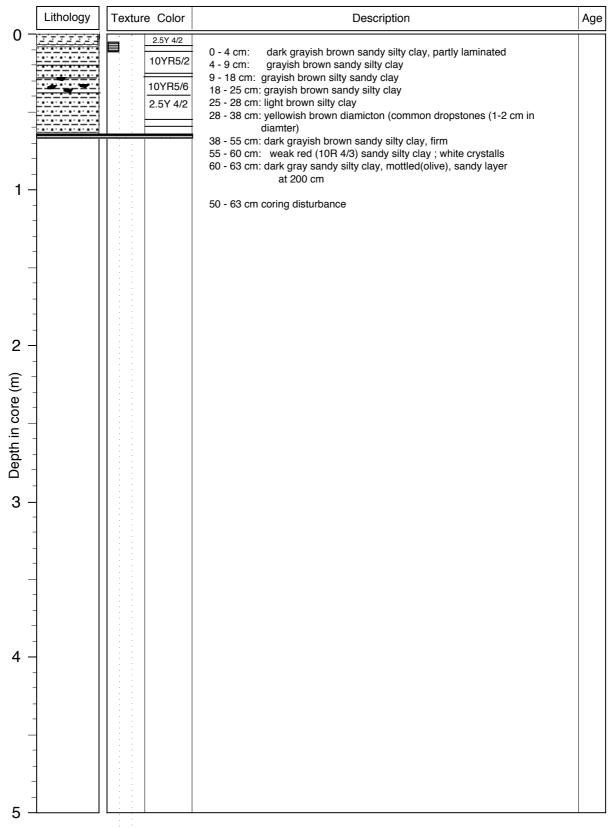
# Yermak Plateau

### ARK XX/3

Recovery: 0.63 m

81° 56.10′N, 08° 31.50′ W

Water depth: 719 m



PS66/335-2 (SL)

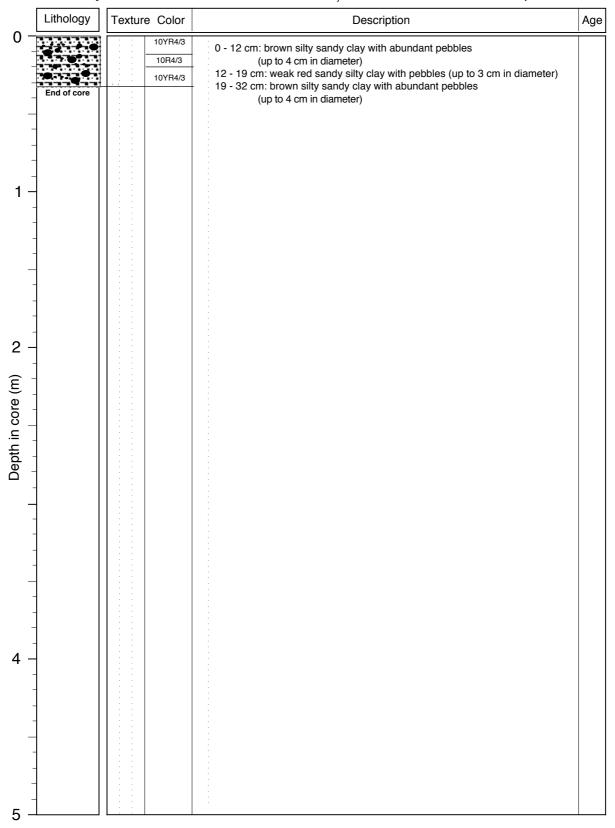
Yermak Plateau

ARK-XX/3

Recovery: 0.32 m

81° 36.05′ N, 06° 34.06′ E

Water depth: 658 m



## PS66/341-1 SL

## Fram Strait

### ARK XX/3

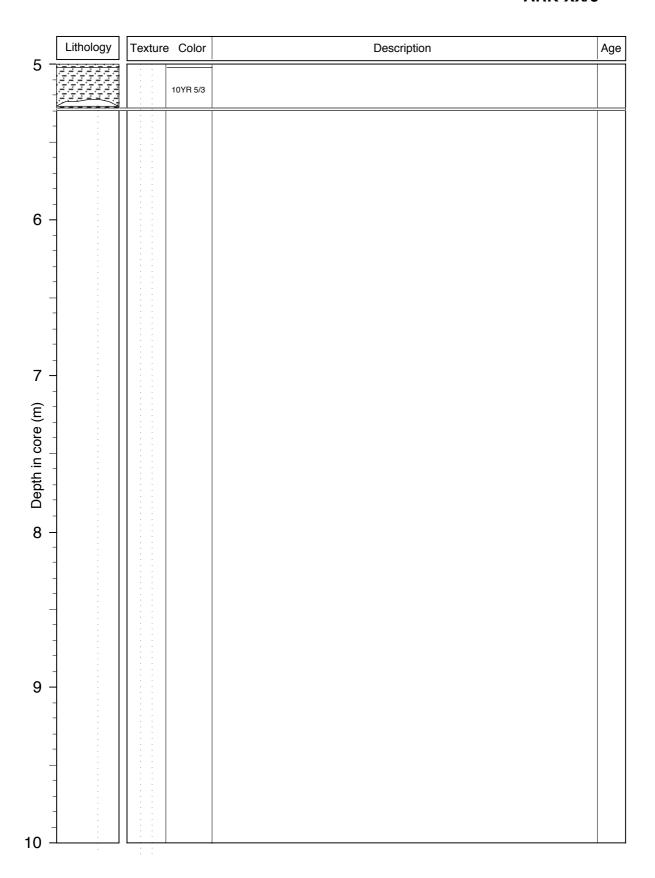
Recovery: 5.28 m

79° 44.07′ N, 00° 44.97′ E

Water depth: 2822 m

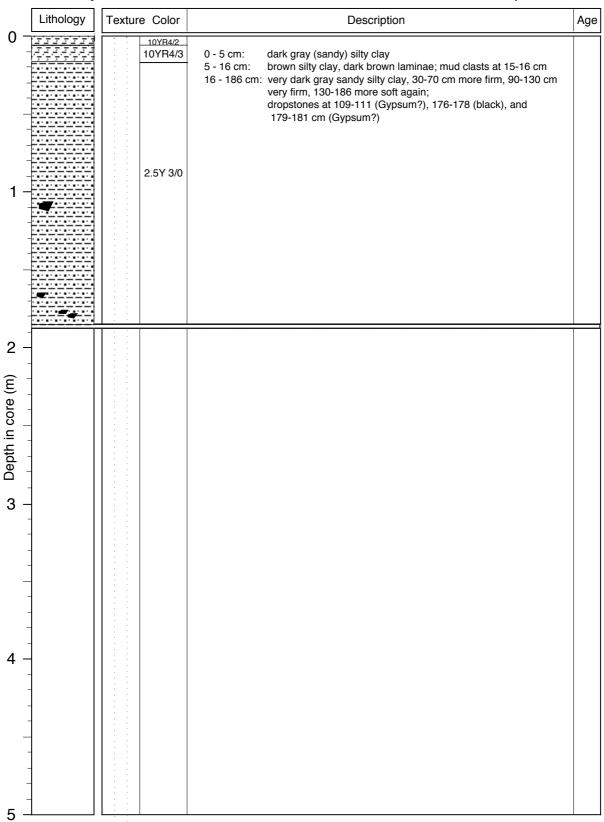
_	Lithology	Texture Color	Description	Age
0 -		10YR 4/3	0 - 30 cm: brown silty clay, homogeneous 30 - 35 cm: very dark brown silty clay; mud clasts at 34-35 cm	
	7666666		35 - 37 cm: yellowish brown silty clay	
	- কুকুকুকুকুকু -	10YR 3/2 10YR 5/4	37 - 41 cm: grayish brown (2.5Y 5/2) and dark brown (10YR 3/3) silty clay	
	122222	2.5Y 4/2	41 - 48 cm: dark grayish brown silty clay; dark brown lenses	
-	333333	2.5Y 4/1	48 - 61 cm: dark gray silty clay	
	777777	2.5Y 5/1	61 - 66 cm: gray silty clay	
	1222222	2.5Y 5/2	66 - 86 cm: grayish brown silty clay 86 - 92 cm: dark gray silty clay, olive mottling (bioturbation)	
			92 - 109 cm: grayish brown silty clay	
		2.5Y 4/1	109 - 111 cm: clayey silty sand, fining upwards; common foraminifer Pyrgo	
1 -	<del>1</del> 222223	2.5Y 5/2	111 - 128 cm: grayish brown silty clay, olive mottling (bioturbation)	
		A :	128 - 132 cm: brown and dark brown silty clay, laminated	
	1555555	555 10YR 5/3	132 - 136 cm: dark grayish brown silty clay 136 - 142 cm: brown and dark brown silty clay	
	<del>- [222222]</del>	2.5Y 4/2	142 - 159 cm: dark grayish brown silty clay; dark brown sandy lenses	
		10YR 3/2	at 146-149 cm	
-	-6666664	2.5Y 4/2	159 - 242 cm: grayish brown silty clay, mottled/bioturbated	
	1272727	\$25	242 - 252 cm: dark gray silty clay	
	-19999999	SSS SSS SSS SSS SSS SSS SSS SSS SSS SS	252 - 271 cm: dark gray silty clay 271 - 284 cm: olive gray silty clay	
	-6666664	ŞŞŞ	284 - 292 cm: light olive brown silty clay	
		\$25	292 - 307 cm: dark olive gray silty clay	
2 -	-199999999	\$55 10YR 5/2	307 - 313 cm: dark gray silty clay	
	1666666	ŞŞŞ	313 - 328 cm: light brownish gray silty clay	
<del>_</del>		\$ SS	328 - 343 cm: dark olive gray silty clay; very dark gray intervals at 333-335 and 342-343 cm	
Ξ	-222223	1777	343 - 364 cm: olive gray silty clay	
Depth in core (m)	-656566		364 - 397 cm dark gray silty clay	
8 -	777777	10YR 4/1	397 - 410 cm: dark grayish brown silty clay	
.⊑	-222223	5Y 4/1	410 - 431cm: dark gray silty clay	
닭	-333333	222	431 - 449 cm: olive gray silty clay 449 - 457 cm: dark brown and grayish silty clay	
)e	1666666	\$\$\$ 5Y 4/2	457 - 464 cm: dark gray silty clay	
_		2.5Y 5/4	464 - 468 cm: dark olive gray silty clay, very dark gray horizons at 464	
3 -	1222223	5Y 3/2	and 467 cm	
•	- <del>22222</del>	2.5Y 4/0	468 - 471 cm: gray silty clay	
	-666666	2.5Y 6/2	471 - 477 cm: grayish brown silty clay 477 - 481 cm: dark gray silty clay	
			481 - 483 cm: gray silty clay	
	19999999	5Y 3/2	483 - 490 cm: olive gray silty clay	
_	-555555		590 - 495 cm: dark gray silty clay	
		5Y 4/2	495 - 502 cm: gray silty clay, dropstone (1cm in diameter) at 498 cm 502 - 522 cm: brown silty clay	
			our our only day	
	-6666664	5Y 4/1	(516 - 528 cm: core catcher; 522-528 cm empty)	
	1666666			
4 -				
•	3333333	10YR 4/2		
	444444	2.5Y 4/0		
	777777			
	1222223	5Y 5/2		
_	<u> </u>			
	<del>                                     </del>	10YR 4/2 5Y 4/1		
	444444			
		10YR 5/2		
	555555	5Y 5/2		
5 -		2.5Y 4/0 2.5Y 5/0		
5				

### ARK-XX/3



### PS66/345-1 East Greenland Continental Shelf ARK XX/3

Recovery: 1.86 m 79° 03.58′ N, 11° 15.11′ W Water depth: 318 m



## PS66/345-2 East Greenland Continental Shelf ARK XX/3

Recovery: 1.51 m 79° 03.58′ N, 11° 16.03′ W Water depth: 313 m

		2.5Y 3/2 2.5Y 4/4 2.5Y 3/1		dropstone (diameter of very dark gray sandy s common dropstones of (core catcher): very da	n sandy silty clay n silty clay, more sandy layer v 2 cm) at 73-74 cm ilty clay, firm, f 1 - 7 cm in diameter (> tillite rk gray sandy silty clay, firm, gypsum (one large "crystall" of	)
			127 - 151 CF	common "crystalls" of g	gypsum (one large "crystall" of	
		2.5Y 3/1				
	: :		:			
					: : : : :	
:						

## 12.4 Geological data tables (Smear-slide data) N. Kukina

Т	able 12	<b>2.4.1a</b> : Βι	ılk mi	neralogy of	sediment	s from Co	re PS66/	304-3, bas	ed on s	mear-slide estimates
Depth	Quartz	Feldspar	Mica	Terr.	Volcanic	Clay	Org.	Fe-	Other	Sediment type
				carbonate	glass	minerals	rem ains	hydroxide	Heavy	
0 cm	28,7	16,7	8,3	4,6	2,8	3,7	4,6	7,4	23,2	dark gray sand
10 cm	31	13,3	5,3	10,6	0,1	4,4	8,8	3,5	23	dark gray sandy silt
19 cm	30,2	14,3	6,1	9,5	1,9	10,5	5,2	2,3	20	gray sandy silt
45 cm	25,1	11,1	4,3	5,1	1	30,2	8,2	1,1	13,9	dark olive gray sandy silt
80 cm	20,3	8,9	8,3	0,2	0	36	15	2,3	9	dark olive gray sandy silty clay
100 cm	25,6	11,6	7,2	0	0	29	11,1	5,3	10,2	dark olive gray sandy silty clay
119 cm	33,1	18,2	5,2	6,3	2,9	2,1	4,9	2,3	25	dark gray w ith black silty clay
186 cm	33,9	11,6	4,9	5,9	2,1	7,6	4,4	3,1	26,5	dark gray silty clay
196 cm	34,2	12,6	5,1	5,3	2,2	9,8	6,3	3,3	21,2	dark grayish brown silty clay
219 cm	41	13,2	4,6	11,5	4,6	2,3	2,1	3,5	17,2	dark grayish brown silty clay
281 cm	33,5	13,5	3,2	12,3	4,3	9,3	3,9	6,8	13,2	very dark gray sandy silt
295 cm	39,9	10,5	7,2	10,6	3,3	2,1	2,1	1,1	23,2	dark grayish brown sand
306 cm	38,9	10,2	6,8	11,3	3,5	2,9	2,6	2,1	21,7	dark grayish brown sand
319 cm	34,6	11,9	3,1	12,1	5	12,1	2,1	7,6	11,5	dark gray silty clay
360 cm	26,1	10,2	2,1	5,9	0	26	9,1	2,1	18,5	dark gray sand
397 cm	36,8	16,2	2,9	2,3	0	2,9	3,9	9,9	25,1	gray sand
410 cm	40,2	18,6	3,1	5,3	0	1,9	2,3	2,9	25,7	gray sand

Table 12.4.1b:	Heavy minera	ls in sediment	ts from Core PS	866/304-3, base	ed on smear-slic	de estimates
Depth	Pyroxene	Amphibole	Epidote	Garnet	Black ores	Opaque
0 cm	24.6	33.3	11.1	5.3	19.4	6.3
10 cm	31.3	25	6.3	6.3	18.8	12.3
19 cm	33.2	24.1	11.2	5.9	16.9	8.7
45 cm	29.6	29.8	6.2	11.3	18.5	4.6
80 cm	33.2	27.2	3.2	12.5	15.6	8.3
100 cm	33.9	25.1	3.1	13.6	16.3	8
119 cm	32	22	5.4	11.2	11.9	17.5
186 cm	35.2	33.5	6.9	13.9	5.3	5.2
196 cm	34.2	33.9	6.2	11.3	6.9	7.5
219 cm	30.3	37.1	5.4	10.2	7.2	9.8
281 cm	20.1	38.5	1.1	13.5	7.2	19.6
295 cm	20.1	39.8	2	13.8	6.5	17.8
306 cm	21.1	41.2	3.2	12.6	8.2	13.7
319 cm	20.9	37.9	3.9	11.9	9.7	15.7
360 cm	29.9	44.8	3.6	8.2	8.2	5.3
397 cm	26.9	48.2	4.1	8.2	8.4	4.2
410 cm	30.5	32.5	8.9	8.5	14.2	5.4
419 cm	31.5	32.6	9.8	7.9	5.2	13

Table	12.4.2	a: Bulk mi	ineral	ogy of sedi	ments from	Core PS	66/308-3,	based or	n smear-sli	de estimates
Depth	Quartz	Feldspar	Mica	Terr.	Biog.	Volcanic	Clay	Org.	Fe-	Other Heavy
				carbonate	carbonate	glass	minerals	remains	hydroxide	,
0cm	25,3	10	7	1,2		0	36,2	8,2	2,6	9,5
42cm	20	8,2	5,6	0	0	0	2,1	2,3	54,8	7
48cm	20,5	10	10	2,1	4,2	0	40	5,3	1	6,9
54cm	20	11,3	8,6	1,6	2,2	0	34	12	2	8,3
154cm	30	11	7,6	0	0	0	32,1	10	4,1	5,2
173cm	40	15	6,1	2,3	3	6,5	0	0,5	1,2	25,4
180cm	40	12,5	5,8	3,6	1,6	2,2	11,1	3	6,8	13,4
225cm	30,5	8,9	5,2	0	0,2	2,6	40,6	2,2	2,6	7,2
254cm	30	11,3	2,1	6	3,2	2,6	30	4	2,2	8,6
295cm	33,2	11,2	2,6	8,2	4,1	2,3	25	2,1	1,2	10,1
354cm	30	12	5,6	6,8	4,1	0	24	2,2	2,3	13
454cm	31,6	11,2	5,9	7,5	3,9	0	23,9	2,2	3	10,8

Depth	Pyroxene	Amphibole	Epidote	Chlorite	Garnet	Black ore	Opaque	Sulphides
0cm	44,9	26,3	8,2	0	6,6	10	4	C
42cm	26,9	45,6	0	2,1	8,9	8,3	8,2	C
48cm	25,5	44,6	0	0	8,9	12,9	8,1	C
54cm	30,3	45,9	0	0	6,9	10	6,9	C
154cm	35,2	36,1	6,9	1,6	3,2	8,9	8,1	C
173cm	39	18,2	11	6,9	2,1	12,3	8,4	2,1
180cm	10,9	18	2,6	0	3,6	51	10	3,9
225cm	22,6	39,8	1	0	6,5	17,9	8,9	3,3
254cm	20,1	40,2	2,6	0,2	5,6	18,2	9	4,1
254cm	12,9	40,5	3,1	0	7	20	15	1,5
295cm	12,6	41,2	4,1	0,5	6,5	22,6	12,5	C
354cm	22,3	39,5	2,6	0	8,4	10,3	16,9	C
454cm	43,6	22,6	11,2	4,2	2	10	5	1,4
513cm	45,6	20,3	15,5	6,9	2,3	4,2	3,1	2,1

Depth	Quartz	Feldspar	Mica	Terr.	Volcanic	Clay	Org.	Fe-	Other	Sediment type
				carbonate	glass	minerals	rem ains	hydroxide	He avy	
5 cm	29,3	10,2	1,8	8,5	1,7	18,6	13,6	0,5	15,8	dark brow n silty clay
34cm	25,4	9,6	2,7	9,6	0,9	21,1	9,6	0,8	20,3	dark brow n silty clay
59cm	30,1	9,9	2,1	6,3	1,1	3,2	7,2	29,1	11	very dark brown sandy silty clay
70cm	31,7	13,2	1,1	5,5	3,3	11	9,9	1,2	23,1	dark gray silty clay
88cm	36,5	7,7	2,9	9,6	1,9	1,9	9,6	2,9	27	dark grayish brown silty clay
96cm	24,1	7,9	2,5	7,5	1,1	15,1	12,3	13,6	15,9	dark grayish brow n silty clay
128cm	28,2	8,1	4,2	7,9	1,3	15	5,6	18,2	11,5	dark grayish brown silty clay
151cm	20	7,2	2,1	6,7	0,2	2,1	2,1	47,4	12,2	very dark brown sandy silty cla
154cm	29,5	12,4	2,9	6,7	0	1,1	5,9	20,9	20,6	gray sandy silt
162cm	30,2	15,2	2,6	4,5	3,3	15,2	10,3	2,5	16,2	very dark gray sandy silt
198cm	40,9	21,3	3,1	6,9	1,1	2,1	2,9	2,1	19,6	dark grayish brow n silty clay
214cm	37,4	22,5	4,6	2,1	3,4	1,5	5,8	2,1	20,6	dark gray sand
235cm	35,8	20,5	3,8	2,6	2,5	2,1	6,2	8,2	18,3	dark gray sand
268cm	26,9	14,6	3,9	2,3	1,2	2,2	4,1	18,9	25,9	dark gray sand
328cm	39,1	18,9	5,1	2,3	2,1	1,1	3,4	4,8	23,2	olive brown silty clay
336cm	29,1	15,2	7,9	0,9	2,1	12,9	7,2	6,9	17,8	dark brow n silty clay
348cm	38,4	21,1	5,1	1,6	0	2,6	5,1	6,9	19,2	dark gray silty clay
369cm	21,9	13,4	5,6	3,2	0	23,9	6,9	8,6	16,5	dark gray silty clay
377cm	21,1	19,3	6,2	9,2	0	21,2	3,6	4,6	14,8	olive brown silty clay
116cm	22,6	19,6	4,2	10,3	0	22,9	3,3	5,1	12	olive brown silty clay
198cm	32,1	24,1	5,6	2,1	0	3,3	1,1	12,1	19,6	dark gray sandy silt
522cm	36,5	12,5	2,3	0,2	1,6	6,9	0,9	24,3	14,8	dark gray sandy silt
526cm	44,6	22,1	2,9	0,5	3,5	3,8	1,9	3,6	17,1	dark grayish brow n silty clay
30cm	40,9	21,3	3,1	6,7	0,9	2,2	2,8	2,1	19,6	dark grayish brow n silty clay
582cm	40	25	4,6	5,9	0	2,8	1,1	2,6	18	brown silty clay
87cm	34,6	28	4,2	1,2	0	3,3	5,9	6,9	15,9	dark brown silty clay
97cm	35,9	22,1	8,9	5,1	0	6,6	2,1	1,2	18,1	dark grayish brown silty clay
323cm	26	15	6,6	9,4	2,1	25,8	3,3	1,1	10,7	dark grayish brown silty clay
46cm	15,9	8,3	1,2	3,3	0	33,9	4,1	3,1	30,2	gray silty clay
370cm	27,9	8,9	5,9	2,1	0	11	4,1	5,1	35	dark gray silty clay
83cm	40	20	8	0	0	2,3	1,1	1	27,6	dark grayish brown silty clay
86cm	34,3	15	5,6	0	0	1,1	0,2	4,9	38,9	dark brown silty clay (kakao)
393cm	33,5	20	7,9	0	0	29	0,5	2,1	7	dark grayish brown silty clay

Depth	Pyroxene	Am phibole	Epidote	Garnet	Black ores	Opaque	Sedim ent type
5 cm	29,4	35,3	8,8	11,8	8,9	5,8	dark brow n silty clay
34cm	20,6	36,1	5,3	12	13,2	12,8	dark brow n silty clay
59cm	32,2	28,1	6,5	8,8	12,5	11,9	very dark brown sandy silty clay
70cm	25,8	35,5	12,9	6,5	16,1	3,2	dark gray silty clay
88cm	25	35,7	14,3	3,6	12	9,4	dark grayish brown silty clay
96cm	26,9	33,1	11,2	2,1	11,2	25,5	dark grayish brown silty clay
128cm	31,2	34,8	9,5	8,2	12,1	4,2	dark grayish brown silty clay
151cm	30,2	38,2	8,4	12,1	8,9	2,2	very dark brown sandy silty clay
154cm	31,7	31,8	9,2	8,2	15,6	3,5	gray sandy silt
162cm	34	36,2	9,2	6,9	10,2	3,5	very dark gray sandy silt
198cm	30,1	36,2	12,1	6,1	9,2	6,3	dark grayish brown silty clay
214cm	37,1	27,8	14,2	9,1	3,6	8,2	dark gray sand
235cm	36	27,1	10,2	8,6	12,2	5,9	dark gray sand
268cm	31,5	28,6	11,3	7,9	12,5	8,2	dark gray sand
328cm	38,1	25,1	21,8	6,2	6,2	2,6	olive brown silty clay
336cm	35,1	28,1	15,2	11,7	9,9	0	dark brow n silty clay
348cm	36,5	26,9	14,2	12,1	9,2	1,1	dark gray silty clay
369cm	36,6	26,1	9,5	13,6	12,3	1,9	dark gray silty clay
377cm	33,1	33,5	4,9	5	12,3	11,2	olive brown silty clay
416cm	33,2	33,2	5,2	6,2	11,1	11,1	olive brown silty clay
498cm	32,6	33,6	6,8	7,5	12,9	6,6	dark gray sandy silt
522cm	25,6	33,2	7,9	8,2	13,6	11,5	dark gray sandy silt
526cm	30,1	29,9	6,2	8,2	18,2	7,4	dark grayish brown silty clay
530cm	33,5	21,9	5,2	12,9	17,2	9,3	dark grayish brown silty clay
582cm	35,8	22,1	8,6	8,2	18,6	6,7	brown silty clay
587cm	22,6	38,5	8,7	7,9	11,2	11,1	dark brow n silty clay
597cm	36,5	31,2	12,8	6,2	8,2	5,1	dark grayish brown silty clay
623cm	35,9	28,2	15,9	2,3	11,6	6,1	dark grayish brown silty clay
646cm	36,1	22,3	11,3	8,9	7,2	14,2	gray silty clay
670cm	33,6	26,9	9,2	6,3	6,9	17,1	dark gray silty clay
683cm	33,6	22,1	13,6	11,2	12,1	7,4	dark grayish brown silty clay
686cm	28,9	19,6	11,3	7,2	1,8	31,2	dark brow n silty clay (kakao)
693cm	26,9	18,9	9,8	6,3	12,3	25,8	dark grayish brown silty clay
715cm	24,5	23,9	9,1	12,4	8,2	21,9	gray silty clay

PS66/312-2SL	Quartz	Feldspar	Mica	Terr.	Biog.	Volcanic	Clav	Org.	Fe-	Other
1 000/012 202	Quuitz	Telaspai	miou	carbonate	carbon	glass	minerals	remains		Heavy
0 cm	25,3	10,2	7,3	4,3	1,3	2,1	25,3	8,6	5,6	10
32 cm	20,1	10,3	7,2	3,2	6	4,1	24,3	7,2	8	9,6
132 cm	30,2	18,9	5,2	12,3	6	3,9	8,3	5,2	2	8
228 cm	22,7	15	4	11,9	5	3,5	11,2	5,2	5	16,5
329 cm	23	14,7	4	12	3	4,1	10	3,6	5,6	20
432 cm	30	12	7,9	8	6	1	15	2,1	3	15
533 cm	30,5	15,2	3,6	0	0	4	7,5	3,2	15,9	20,1

Depth	Pyroxene	<b>Am phibole</b>	Epidote	Chlorite	Garnet	Black ore	Opaque	Sulphides
0 cm	47,5	19	15,6	5,2	3,5	5,2	4	0
32 cm	44,2	18	14	5,2	3	8,8	6,8	0
132 cm	40	33	7,9	1,2	5,6	8,6	3,7	0
228 cm	44	29	8,2	2,6	4,1	8,1	4	
329 cm	35,9	20,6	11,3	2,6	4,8	9,9	8,8	6,1
432 cm	44,2	32,5	8,9	0	2,1	6,2	2,3	3,8
533 cm	30,5	15,2	6,1	8,1	4,2	10,1	15,2	10,6
633 cm	40,3	20	5	0	3,5	18	10,3	2,9

Table 12.	<b>4.5a</b> : Bulk	mineralogy of	sediment	s from Core PS66	6/318-4, based o	n smear-slide es	stimates
Depth	Quartz	Field spar	Mica	Clay minerals	Org. remains	Fe-hydroxide	Other Heavy
4 cm	20	12	6,2	33,8	8,5	4,2	15,3
13 cm	26,3	13,2	7,5	11,4	10,8	9,6	21,2
58 cm	20,6	8,5	7,9	27	8,5	15,9	11,6
69 cm	21,3	9,6	7,6	25,3	6,3	16,2	13,7
101 cm	21,3	10,2	6,9	15,9	6,5	20,6	18,6
112 cm	29,3	16,2	7,9	5,6	4,9	14,9	21,2
124 cm	31,5	18,2	10,9	12,3	3,9	8,2	15
134 cm	35,9	21,3	11,2	6,9	3,3	6,5	14,9
139 cm	26,7	15,9	9,2	7,2	6,9	12,9	21,2
160 cm	40,2	22,5	7,2	15,9	3,6	2,1	8,5
172 cm	34,2	22,6	5,2	19	6,2	5,6	7,2
180 cm	28,6	18,2	13,9	6,9	7,1	4,2	21,1
189 cm	38,3	22,1	14,3	6,2	2,3	2,3	14,5
199 cm	42,1	15,9	9,2	8,2	2,6	3,5	18,5
206 cm	29	16,2	13,9	3,9	4,5	6,9	25,6
214 cm	18,9	7,4	4,4	35,1	8,2	12,5	13,5
222 cm	20,2	10,2	8,3	6,3	3,6	12,2	39,2
229 cm	33,2	18,2	7,9	22,9	5,4	2,3	10,1
238cm	32,9	18,5	7,7	21,5	5,2	3,8	10,4
246 cm	10,2	5,9	1,2	1,2	1,2	75,6	4,7
250 cm	29,2	18,5	6,9	20,1	2,3	2,5	20,5
255 cm	35,6	20,1	7,9	5,9	2,6	8,5	19,4
266 cm	36,6	20,1	8,9	5,9	1,8	8,8	17,9
278 cm	24,4	18,6	5,6	18,2	2,3	8,8	22,1

Depth	Pyroxene	Amphibole	Epidote	Garnet	Black ores	Opaque	
4 cm	26,9	40,2	5,9	3,8	13,7	9,5	
13 cm	28,6	30,2	8,3	6,2	16,5	10,2	
58 cm	25,3	36,8	6,8	6,9	16,3	7,9	
69 cm	24,6	38,2	5,3	8,9	18,3	4,7	
101 cm	26,3	40,5	6,2	9,1	11,3	6,6	
112 cm	20,3	40,6	5,2	8,5	15,2	10,2	
124 cm	24	40,2	3,2	11,9	16,2	4,5	
134 cm	24,8	45,6	4,2	11,3	9,9	4,2	
139 cm	22,1	36,6	3,2	11,2	15,3	11,6	
160 cm	24,3	46,3	5,2	8,2	7,9	8,1	
172 cm	20,3	48,2	4,3	5,6	14,2	7,4	
180 cm	20,1	50,2	3,6	12,6	9,2	4,3	
189 cm	18,9	46,3	4,2	11,3	10,2	9,1	
199 cm	27,3	31,2	6,9	13,6	11,2	9,8	
206 cm	28,6	29,3	8,9	12,5	12,5	8,2	
214 cm	23,1	33,8	4,2	7,1	8,9	22,9	
222 cm	5,5	13	0,5	3,6	6,9	70,5	
229 cm	28,9	44,5	6,3	9,5	4,3	6,5	
238cm	27,9	45,3	6,2	9,5	5,1	6	
246 cm	22,1	44,6	5,2	8,2	6,3	13,6	
250 cm	32,5	33,2	9,6	12,9	4,9	6,9	
255 cm	33,6	34,2	12,5	11,5	5,2	3	
266 cm	33,6	35,2	12,4	10,2	6,3	2,3	
278 cm	34,8	38,2	13,5	9,2	2,6	1,7	

Depth	Ouertz	Feldspar	Mica	Terr.	Biog.	Volcanic	Clay	Org.	Fe-	Other	
Depth	Quartz	reiuspai	Wilca						-		
				carbonate	carbonate	glass	minerals	remains	hydroxide	Heavy	
0cm	20	6	2,3	0	0	0	2,1	5,6	56	8	
5cm	30,2	10,6	6,3	0	0	0	32,1	3,1	10,2	7,5	
20cm	40,2	19,5	6	2,6	1,3	0	8,3	2,1	2	18	
50cm	40	15,6	8,3	18,2	8,1	0	4,3	1,2	0	4,3	
60cm	35	15	5,3	12,3	6,2	0	11,5	4,2	0	10,5	
80cm	36,2	14,2	6,1	11,2	7,2	0	9,6	3,1	0	12,4	
90cm	36,4	10	3,4	0,5	0,6	2,6	8,9	4,3	19,6	13,7	
100cm	32,1	9,9	7,3	1,2	1,3	2,1	35,1	2,1	2,1	6,8	
114cm	35,1	11,5	6,6	4,6	1,6	3,2	20,3	2,5	4	10,6	
200cm	40,5	18	8	0	0	0,2	15	1,3	9,6	7,4	
260cm	35	15	6,6	8,9	4,3	1,3	18	0,5	0	10,4	
280cm	35	11	3,2	10	6	1,5	10,1	1,2	2	20	

Depth	Pyroxene	Amphibole	Epidote	Chlorite	Garnet	Black ore	Opaque	Sulphides
0cm	40,3	20,3	8,4	4	3,9	8,2	10,3	4,6
5cm	44,6	14,3	12,3	8	3,2	4,3	11,3	2
20cm	35,6	18,9	13,5	3,2	4,9	12,8	4	7,1
50cm	22,6	36,9	3,4	0	8	8,2	5,6	15,3
60cm	44	20	11,3	4,9	4,1	5	5,2	5,5
80cm	44,2	26	12,3	5,3	3,1	4,9	3,1	1,1
90cm	33,6	20,1	4,9	2,1	5,6	16,9	10	6,8
100cm	41,2	32	8	0	4,2	6,1	4	4,5
114cm	32,9	25,6	4,6	5,6	2,1	15,9	10,3	3
200cm	41,1	15,2	18	8,2	6,2	7,3	4	0
260cm	33,8	21,1	12	2,1	3,9	18	2,3	6,8
280cm	40	18,6	10,1	4,6	1,9	10	3	11,8
300cm	45,2	35	5	2,2	5,6	2,3	0,5	4,2

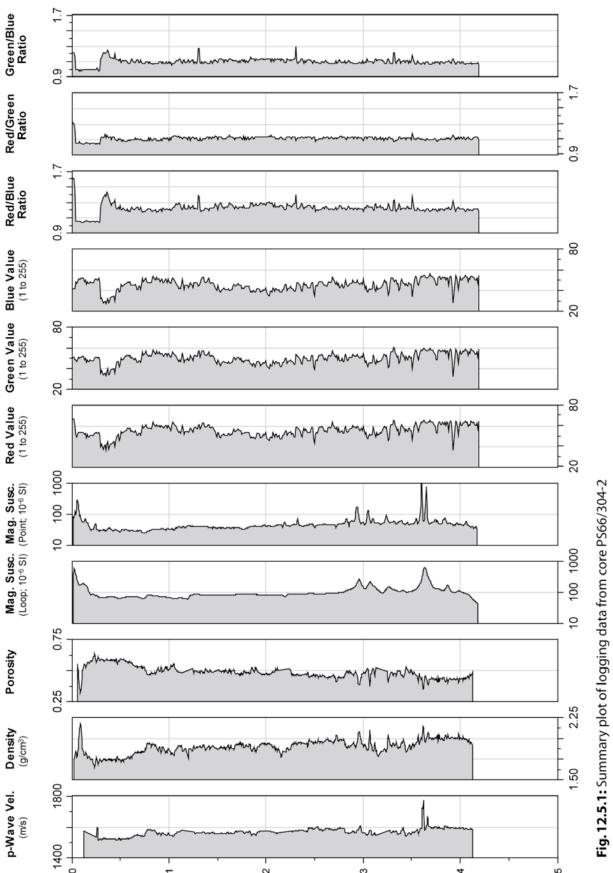
Depth	Quartz	Feldspar	Mica	Terr.	Biog.	volc.	Clay	Org.	Fe-	Other
				carbonate	carbonate	glass	minerals	remains	hydroxyde	Heavy
0 cm	26,3	11,2	5,6	2,1	0	0,1	30,2	11,1	8,1	5,3
4 cm	29,1	10,4	8,2	5,6	1,2	0	5,2	10,4	20,1	9,8
58 cm	36,5	14,2	7,4	12,2	4,9	2,3	4,6	4,9	3,9	9,1
159 cm	25	8	8,2	13,9	4,4	2,2	19	3	2,4	13,9
247 cm	35,8	15,3	4,2	11,1	3,5	4,6	12,2	2,6	2,1	8,6
347 cm	29,6	8,2	6,3	6,2	2,1	2,2	25	7,6	2,1	10,7
448 cm	30	10	6,8	7,9	2	2	24,2	8,1	2,1	6,9
548 cm	37	15,9	5,1	12,6	0,5	4,6	5,2	4,1	8,9	6,1

Depth	Pyroxene	Amphibole	Epidote	Chlorite	Garnet	Black ore	Opaque
cm	38,1	15,6	8,1	0	8,9	17,2	12,1
4 cm	37,1	22,1	2,1	0	12,5	15	11,2
58 cm	50,2	19,6	10,2	4,2	5,1	6,2	4,5
59 cm	47,6	21,6	12,6	5,3	4,1	4,6	4,2
7 cm	40,1	20,3	11,3	5,2	7,1	6,6	9,4
17 cm	39,5	19,5	12,1	2,9	6,6	8,2	11,2
48 cm	40,2	27,7	8,8	4,2	7,6	4,9	6,6
48 cm	40,3	23,1	10,2	4,2	6,2	6,2	9,8

Depth	Quartz	Feldspar	Mica	Terr.	Biog.	Clay	Org.	Fe-	Other	Sediment type
				carbonate	carbonate	minerals	re m ains	hydroxide	Heavy	
21 cm	30,2	10,5	7,1	5,1	0	20,3	7,2	7,1	12,5	olive brown sandy silty clay
89 cm	33,1	12,6	7,4	0	5,1	20	12	0	9,8	gray sandy silty clay
95 cm	25,1	11,1	12,4	0	0	4,6	8,1	30,9	7,8	yellow ish brown sand silty cla
122 cm	37,9	20,1	10	0	0	5,1	5,6	3,1	18,2	black spots sandy silty clay
165 cm	40	20	8	3,1	2,2	10,8	5,6	3,9	6,4	dark olive gray sandy silty cla
194 cm	30,2	10,2	8,5	0,8	4,4	8,1	8,6	15,9	13,3	olive sandy silty clay
214 cm	35,9	19,1	4,2	0	1,1	5,1	10,5	3,1	21	olive gray sandy silty clay
231 cm	38,1	20,1	8,1	0	0	10,6	6,1	9,5	7,5	very dark gray sandy silty cla
284 cm	37,2	19,2	10.7.1	8	5,9	4	10	2,1	13,6	gray sandy silty clay

Table 12	2. <b>4.8b</b> : He	avy minerals	in sedir	ments fror	n Core	PS66/333-4	4, based o	n smear-	slide estimates
Depth	Pyroxene	<b>Am phibole</b>	Epidote	Chlorite	Garnet	Turm aline	Black ore	Opaque	Sediment type
21 cm	40,2	28,1	9,6	2,1	2,3	1	11,6	5,1	olive brown sandy silty clay
89 cm	29	36,6	4,1	0	9,1	0	9,1	12,1	gray sandy silty clay
95 cm	33,9	34,6	11,3	2,1	2,2	1,1	12,3	2,5	yellow ish brown sand silty clay
122 cm	20,5	20,4	10,1	0,9	4,8	1,6	30,2	11,5	black spots sandy silty clay
165 cm	40,2	22,3	8,9	2,2	3,6	1,1	11,2	10,5	dark olive gray sandy silty clay
194 cm	45,1	28,8	9,6	4,8	2,1	0	5,4	4,2	olive sandy silty clay
214 cm	38	30,5	9,5	5,2	7,1	1	6	2,7	olive gray sandy silty clay
231 cm	28	22,6	6,8	2,5	8,8	0,1	21	10,2	very dark gray sandy silty clay
284 cm	43,6	28,3	6,3	2,1	8,4	1,1	5,8	4,4	gray sandy silty clay

Summary plots of logging data for selected sediment cores 12.5 C. Gebhardt, F. Niessen, D. Penshorn, J. Rogenhagen, J. Schneider



Depth (m)

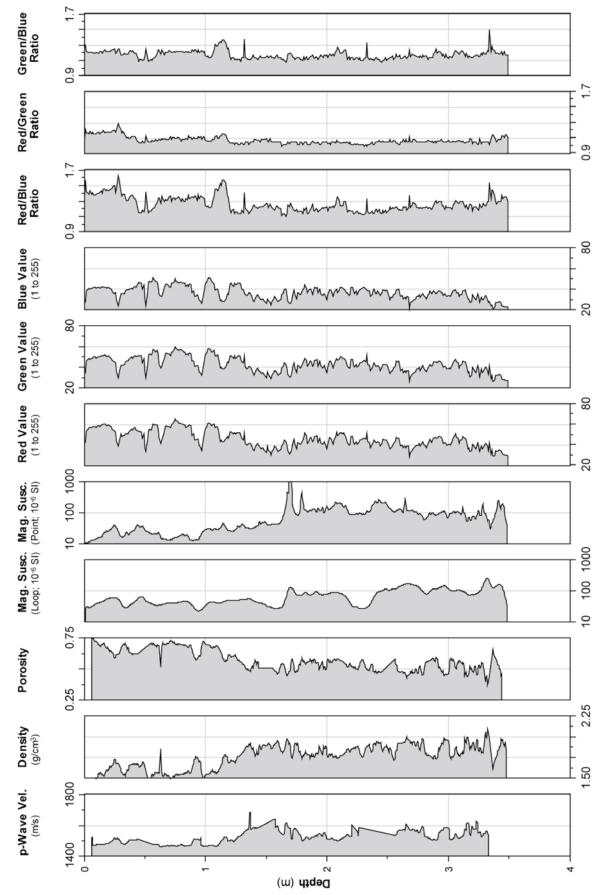


Fig. 12.5.2: Summary plot of logging data from core PS66/306-2

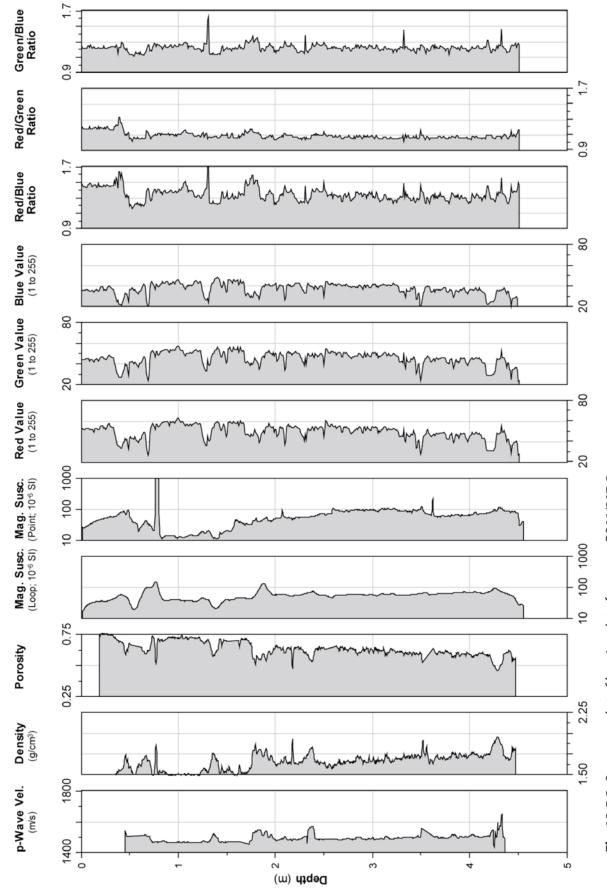


Fig. 12.5.3: Summary plot of logging data from core PS66/307-2

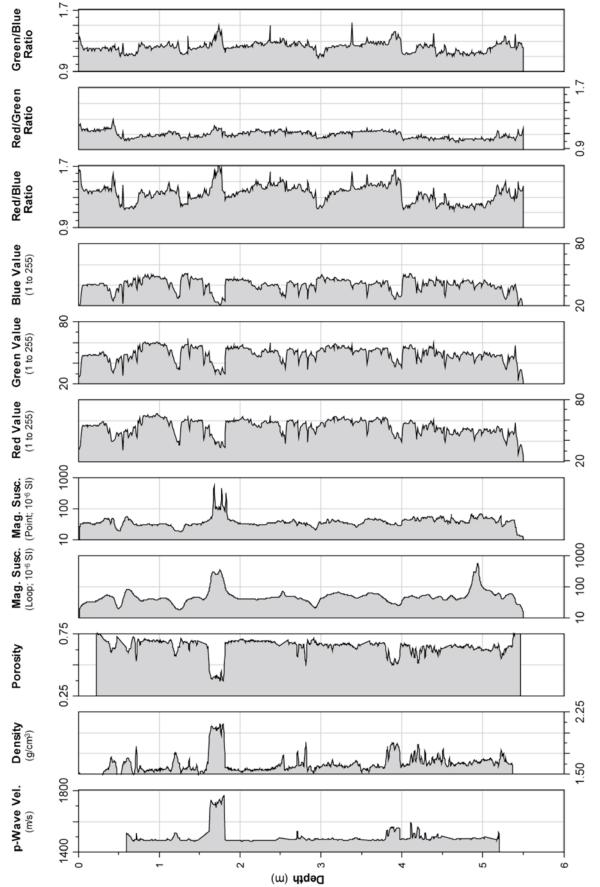


Fig. 12.5.4: Summary plot of logging data from core PS66/308-3

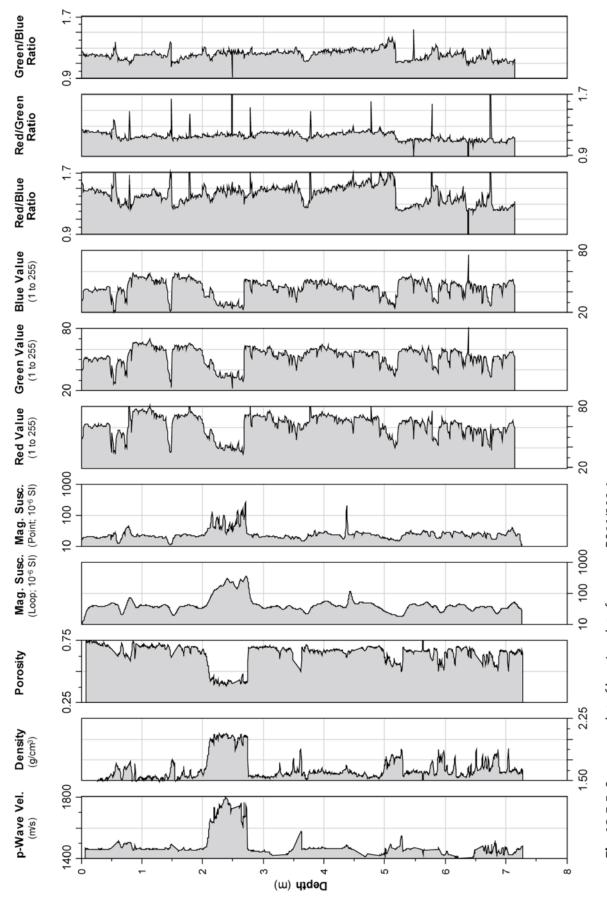


Fig. 12.5.5: Summary plot of logging data from core PS66/309-1

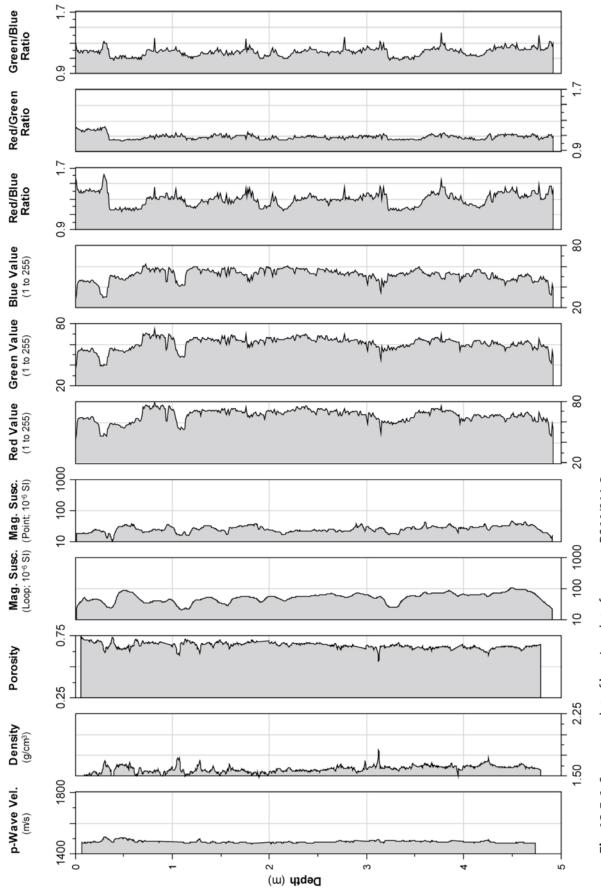


Fig. 12.5.6: Summary plot of logging data from core PS66/311-3

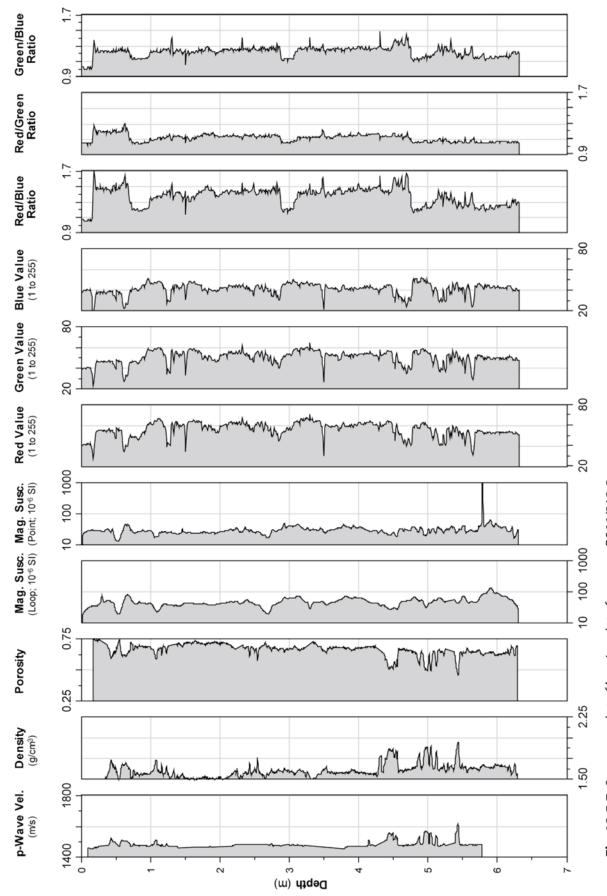


Fig. 12.5.7: Summary plot of logging data from core PS66/312-2

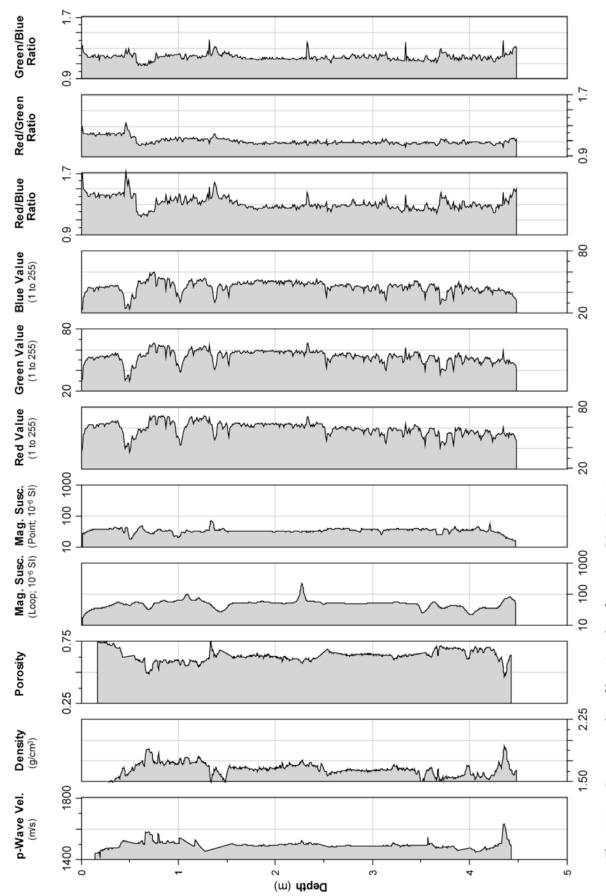


Fig. 12.5.8: Summary plot of logging data from core PS66/313-3

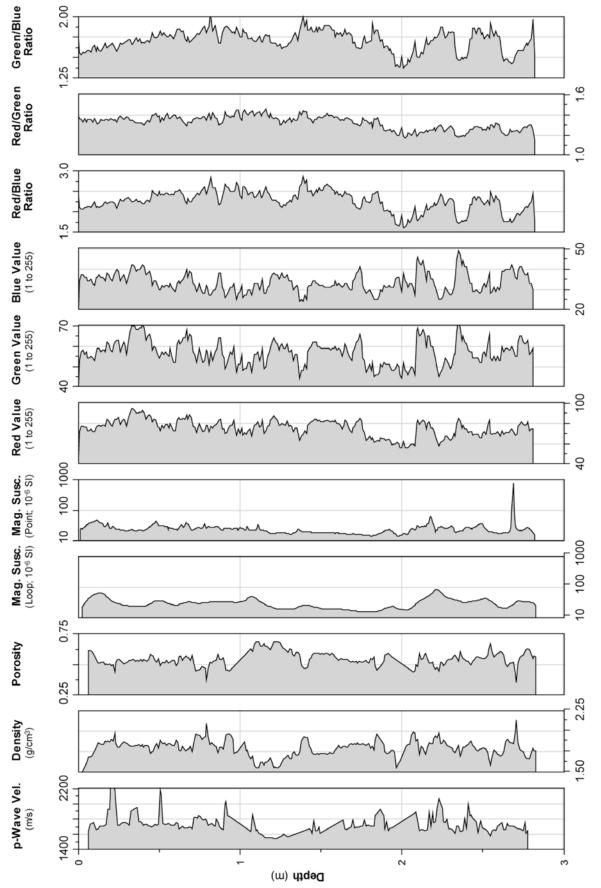


Fig. 12.5.9: Summary plot of logging data from core PS66/318-4

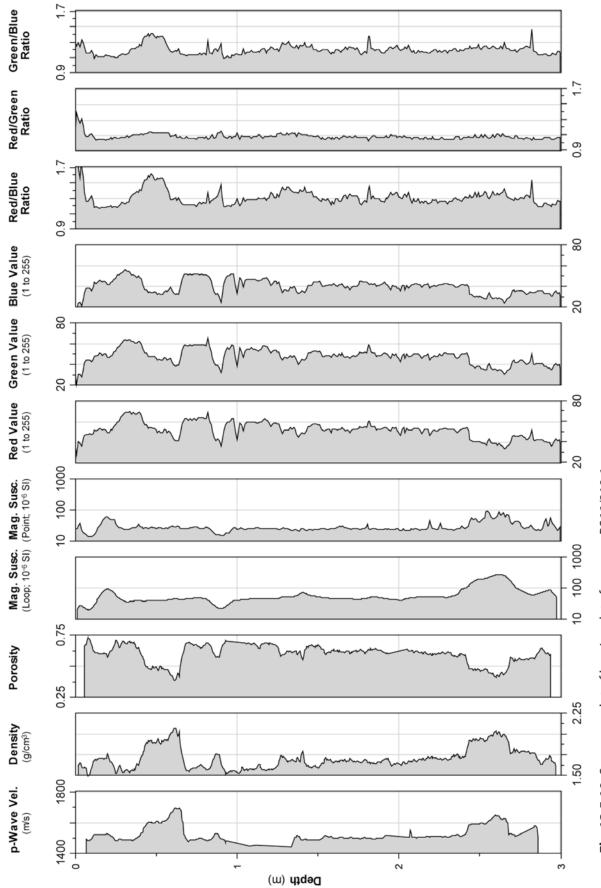


Fig. 12.5.10: Summary plot of logging data from core PS66/319-1

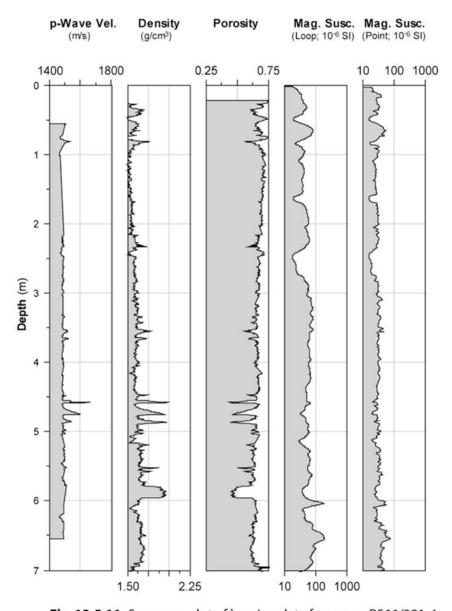


Fig. 12.5.11: Summary plot of logging data from core PS66/321-4

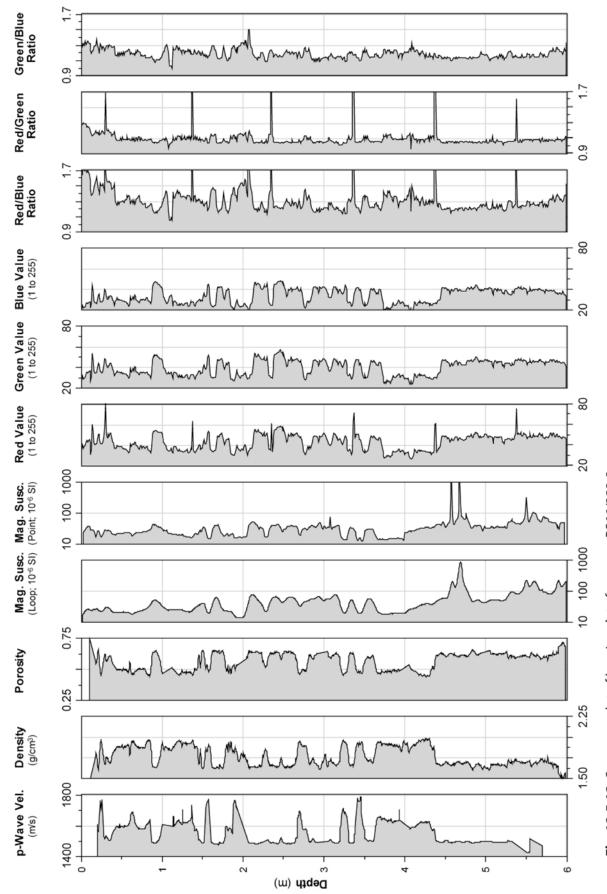


Fig. 12.5.12: Summary plot of logging data from core PS66/322-3

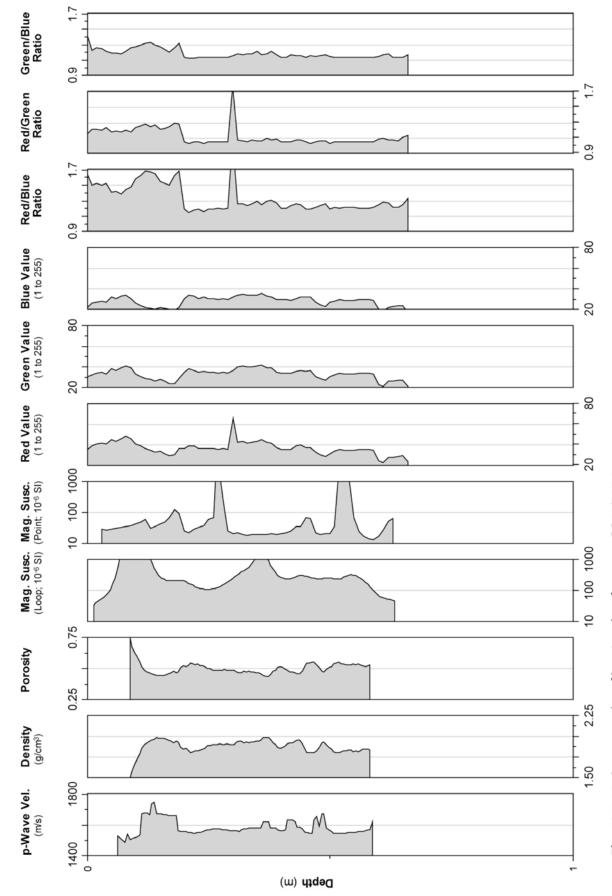


Fig. 12.5.13: Summary plot of logging data from core PS66/323-3

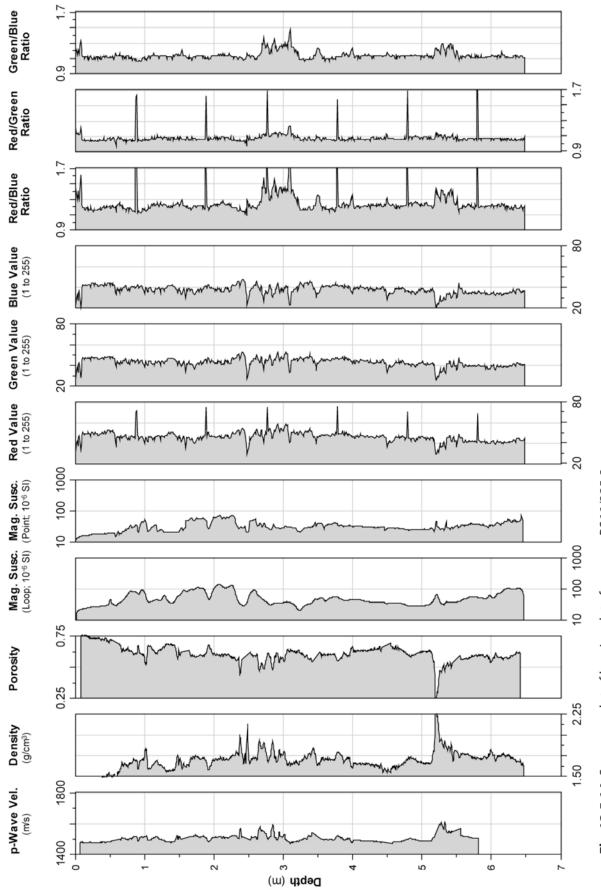


Fig. 12.5.14: Summary plot of logging data from core PS66/325-3

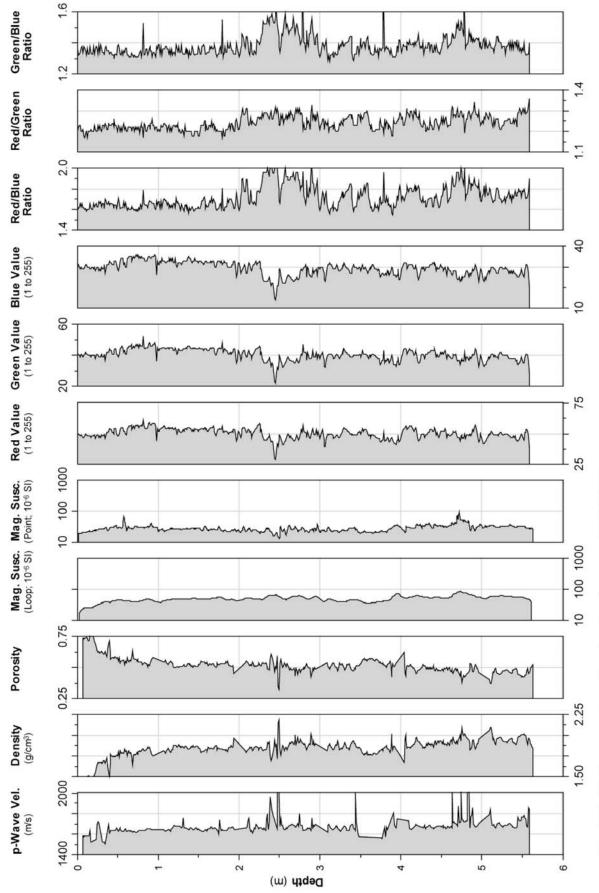


Fig. 12.5.15: Summary plot of logging data from core PS66/327-3

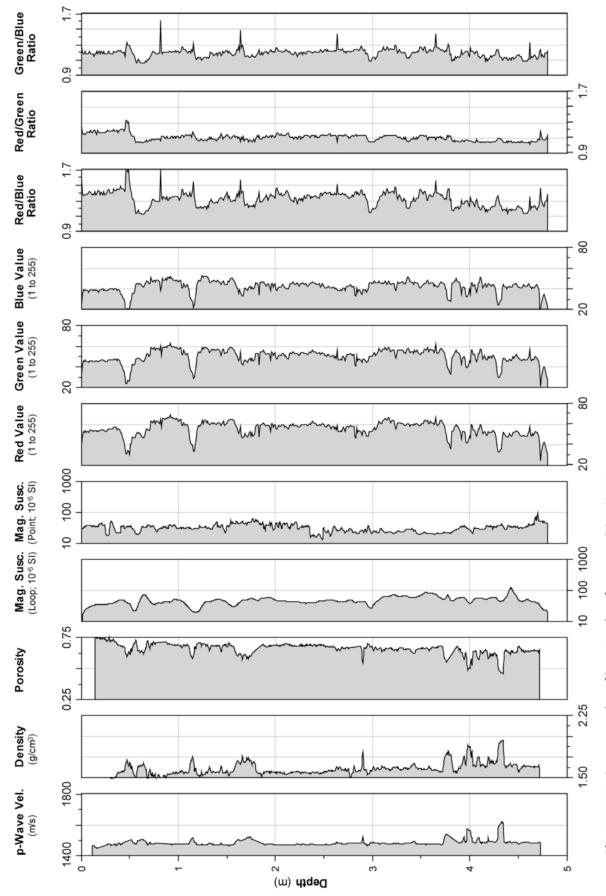


Fig. 12.5.16: Summary plot of logging data from core PS66/329-3

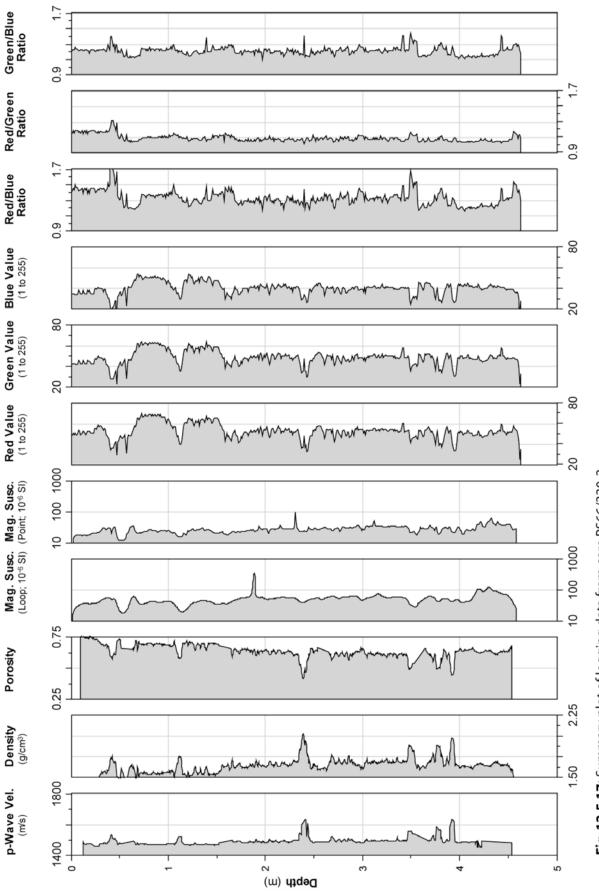


Fig. 12.5.17: Summary plot of logging data from core PS66/330-2

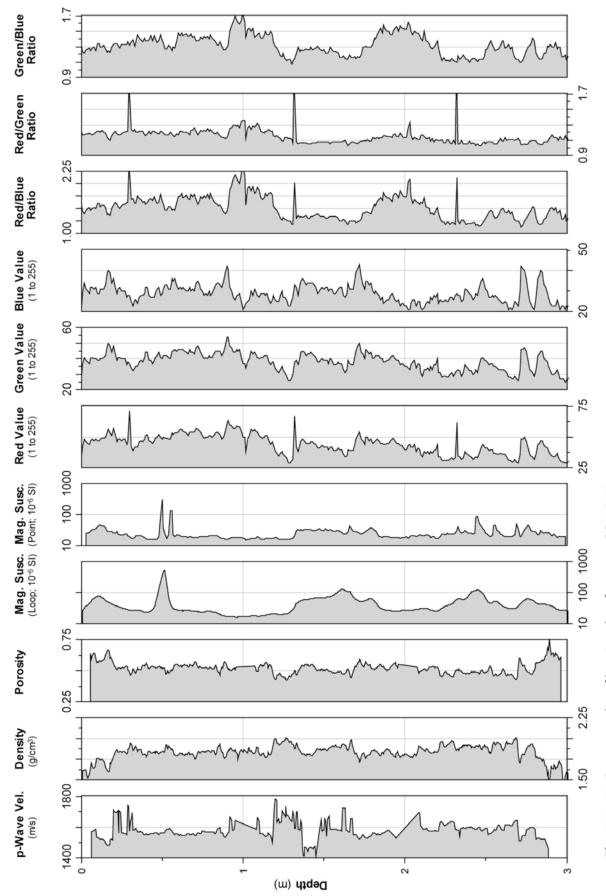


Fig. 12.5.18: Summary plot of logging data from core PS66/333-4

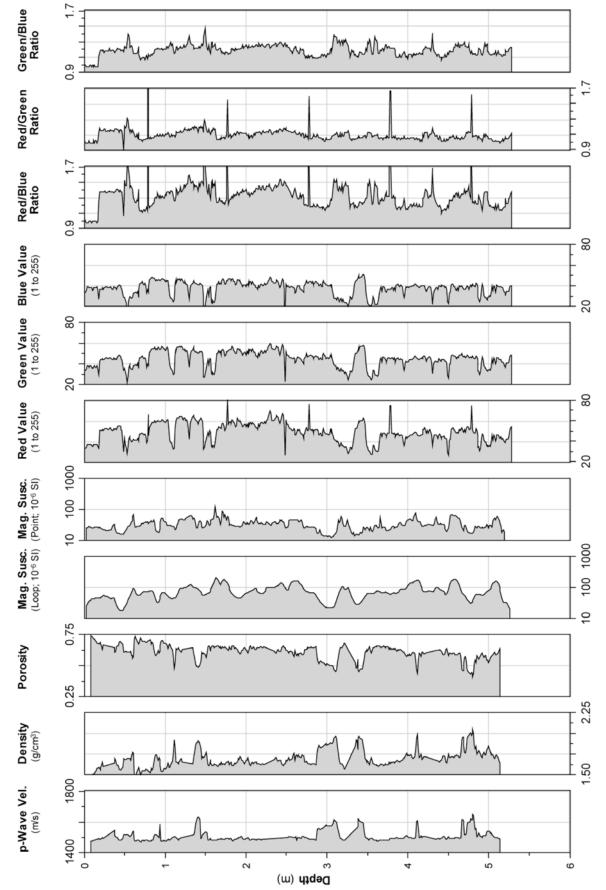


Fig. 12.5.19: Summary plot of logging data from core PS66/341-1

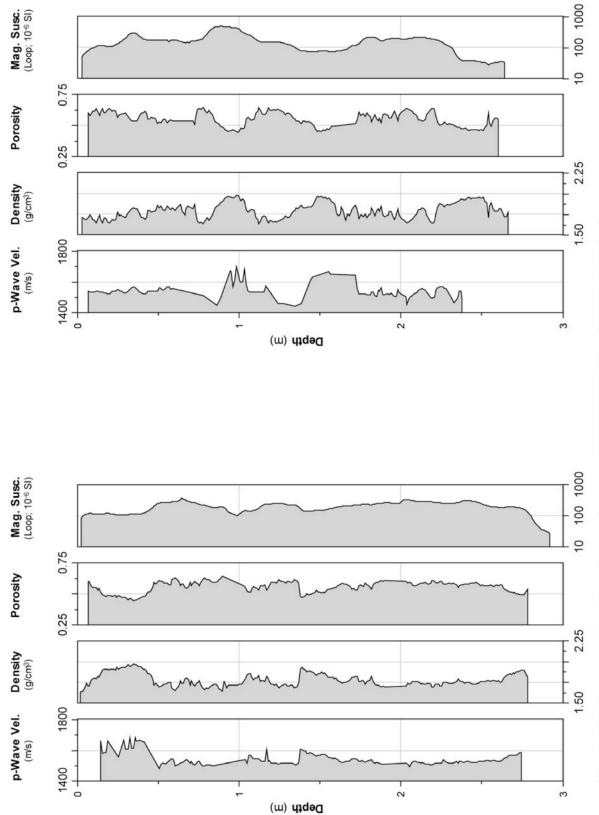


Fig. 12.5.20: Summary plots of logging data from cores PS66/343-1 (left) and PS66/343-2 (right)

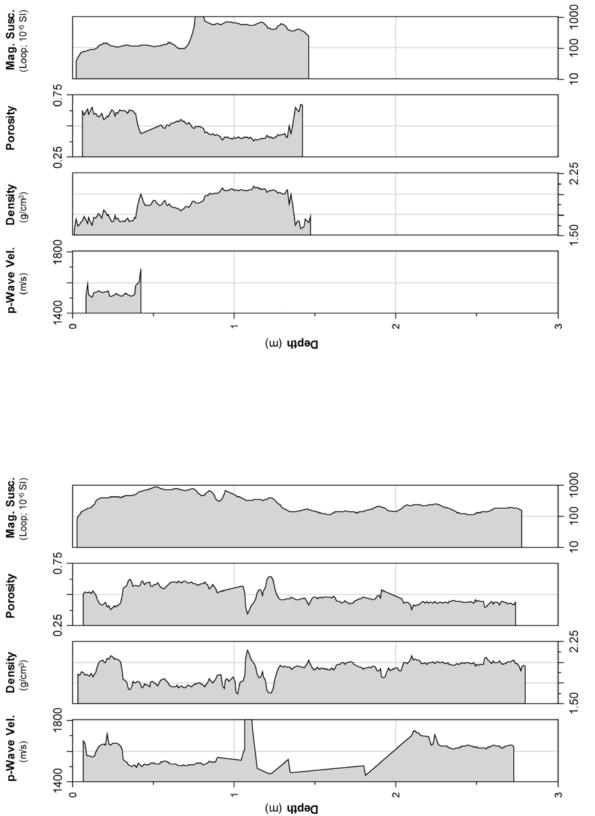


Fig. 12.5.21: Summary plots of logging data from cores PS66/343-3 (left) and PS66/343-4 (right)

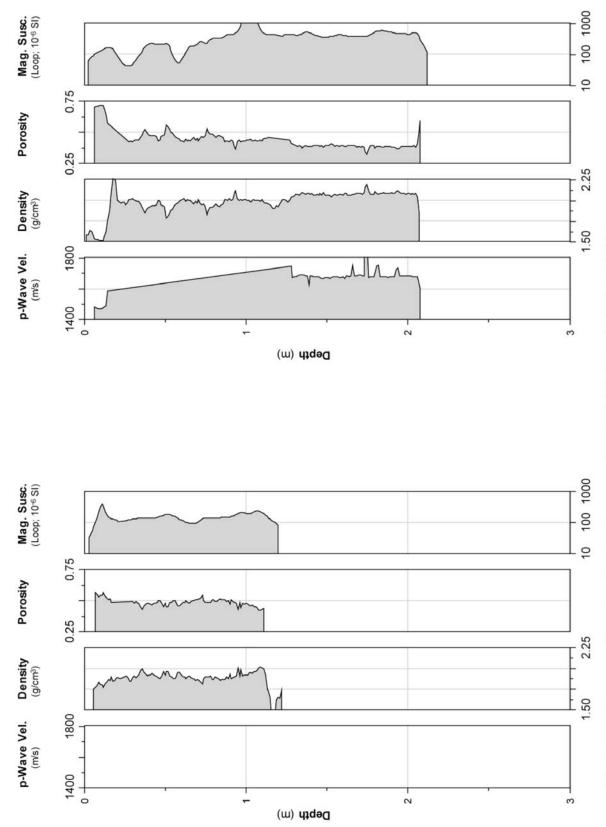


Fig. 12.5.22: Summary plots of logging data from cores PS66/343-5 (left) and PS66/343-6 (right)

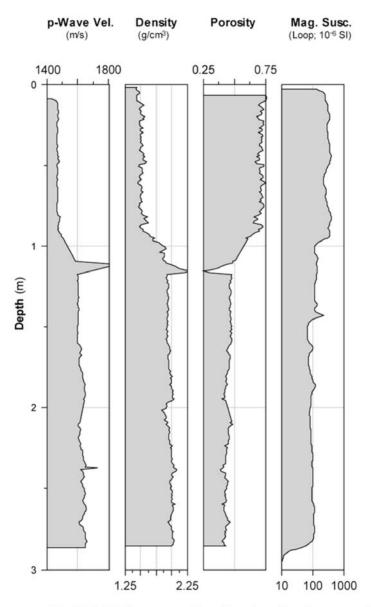


Fig. 12.5.23: Summary plot of logging data from core PS66/344-1

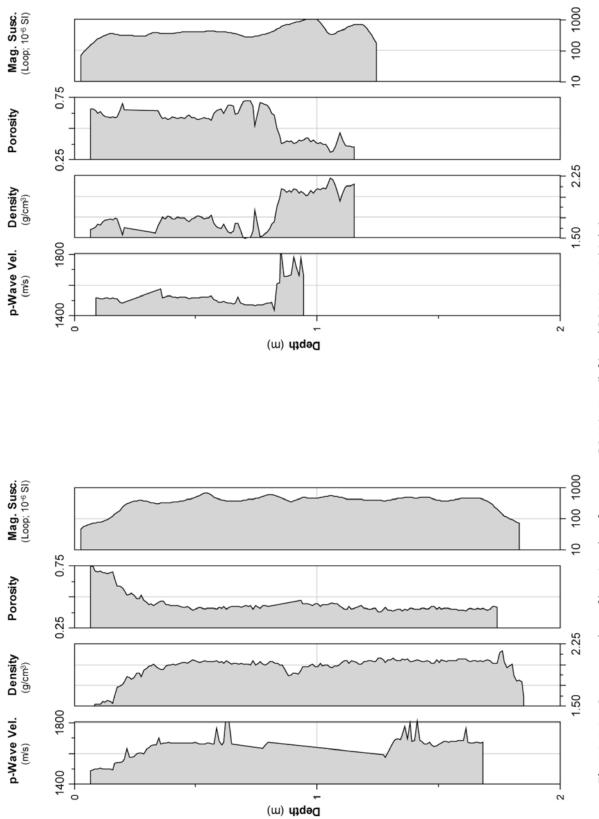
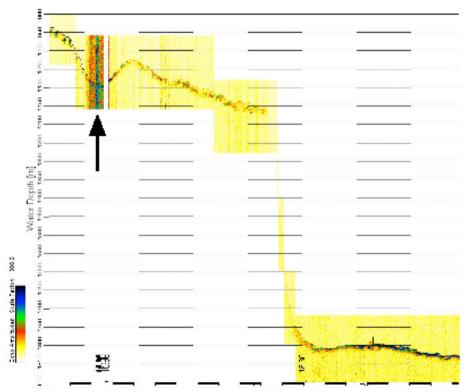


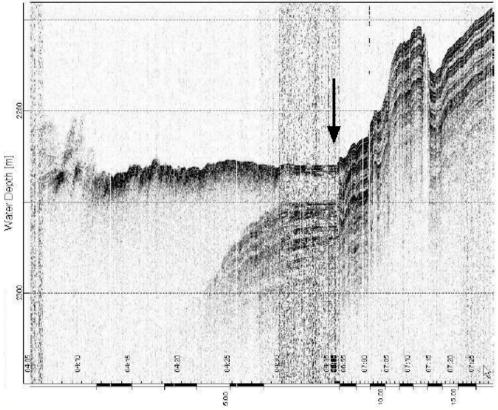
Fig. 12.5.24: Summary plots of logging data from cores PS66/345-1 (left) and PS66/345-2 (right)

## 12.6 Parasound profiles at coring stations

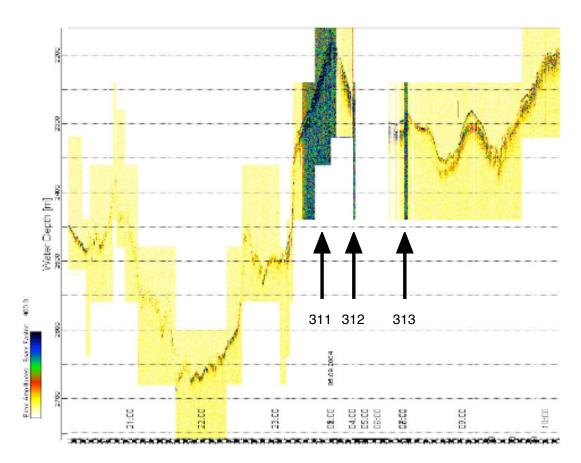
F. Niessen, J. Rogenhagen



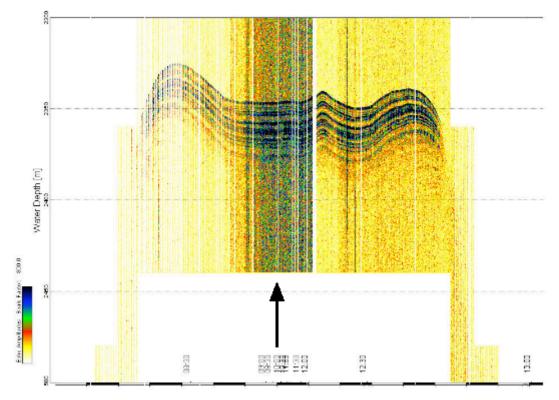
**Fig. 12.6.1**: Geological Situation at Station PS66/304 as recorded by PARASOUND. For details see Station List. Horizontal distance bar (black or white) is 1 km. Numbers are time UTC.



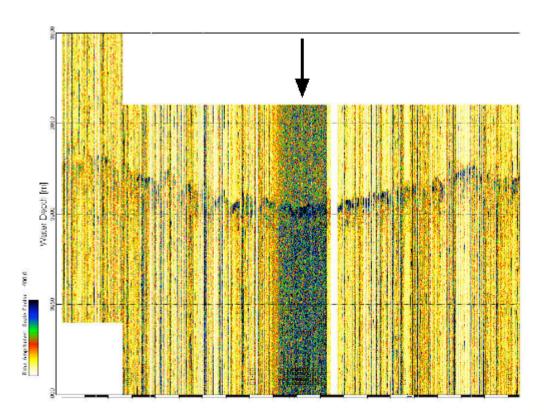
**Fig. 12.6.2**: Geological Situation at Station PS6/307 as recorded by PARASOUND. For details see Station List. Horizontal distance bar (black or white) is 1 km. Numbers are time UTC.



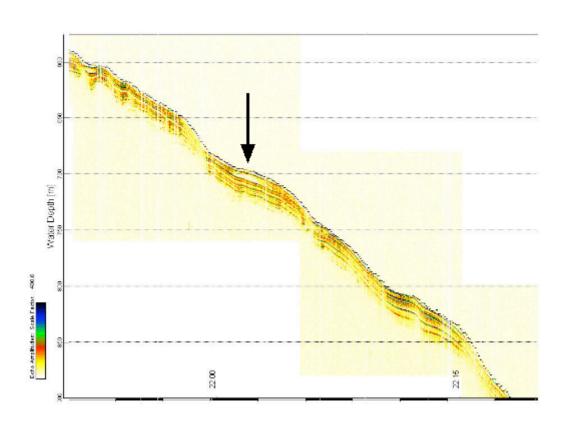
**Fig. 12.6.3**: Geological Situation at Stations PS66/311, /312, /313 as recorded by PARASOUND. For details see Station List. Horizontal distance bar (black or white) is 1 km. Numbers are date and time UTC.



**Fig. 12.6.4**: Geological Situation at Station PS66/321 as recorded by PARASOUND. For details see Station List. Horizontal distance bar (black or white) is 1 km. Numbers are time UTC.



**Fig. 12.6.5**: Geological Situation at Station PS66/323 as recorded by PARASOUND. For details see Station List. Horizontal distance bar (black or white) is 1 km. Numbers are time UTC.



**Fig. 12.6.6**: Geological Situation at Station PS66/327 as recorded by PARASOUND. For details see Station List. Horizontal distance bar (black or white) is 1 km. Numbers are time UTC.

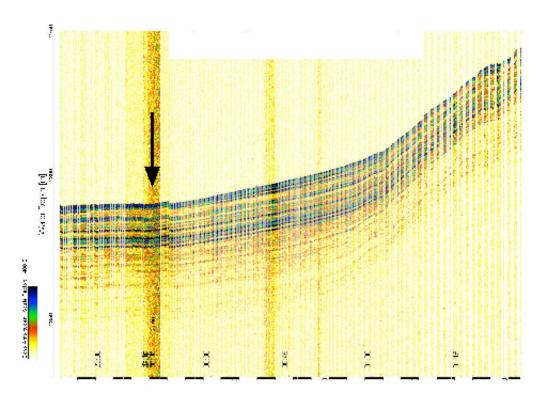
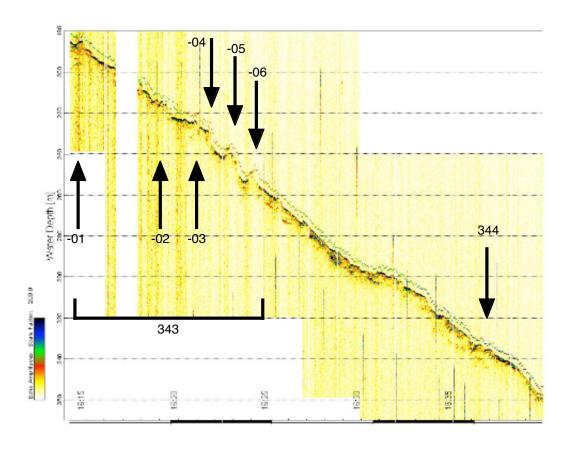
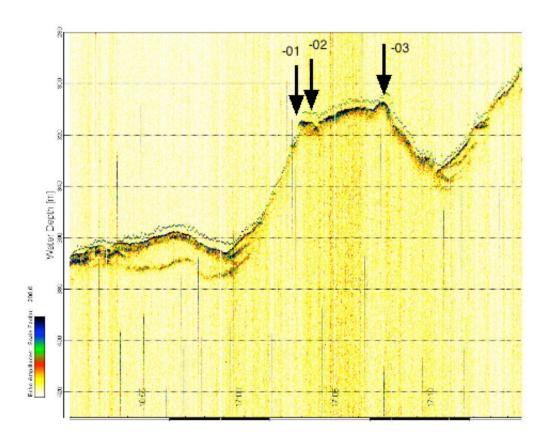


Fig. 12.6.7: Geological Situation at Station PS66/341 as recorded by PARASOUND. For details see Station List. Horizontal distance bar (black or white) is 1 km. Numbers are time UTC.



**Fig. 12.6.8**: Geological Situation at Stations PS66/343 and /344 as recorded by PARASOUND. For details see Station List. Horizontal distance bar (black or white) is 1 km. Numbers are time UTC.



**Fig. 12.6.9**: Geological Situation at Station PS66/345 (-01-03) as recorded by PARASOUND. For details see Station List. Horizontal distance bar (black or white) is 1 km. Numbers are time UTC.

#### 12.7 **List of Participants**

#### **Scientific Crew**

**AWI** Bahr Barbara Behr Yannik **AWI** Birnstiel Hendrik **AWI** 

HeliTransair Büchner Jürgen

Buldt Klaus **DWD** 

Feldt Oliver HeliTransair

Fricke Nikolas Gebauer André Gebauer Manfred Gebhardt Catalina Karsten Gößling Günther Daniel Jokat Wilfried Kukina Natalia Norbert Lensch Martens Hartmut Mende Ingo Nam Seung-il Niessen Frank Noffke Hannah Penshorn Dietmar

Platten Bettina Puhr Andreas Raabe Wolfgang Rathlau Rike Rogenhagen Johannes

Saraswat Rajeev Schäfer Christoph Schmidt-Aursch Michita Schneider Julia Frank Schoster Schroeder Max Shevchenko Vladimir Spengler Thomas Stein Ruediger (Chief scientist) Thiele Julia

Winkelmann Daniel Winkler **Andreas** Winter Stefan Yaroslava Yanina Yousif Khalaf

MPI AWI **DWD** AWI **AWI AWI AWI MMBI** AWI **AWI** MPI **KIGAM AWI AWI AWI** 

**AWI AWI AWI** AWI **FIELAX AWI AWI** AWI AWI **AWI AWI IORAS AWI AWI AWI AWI** 

> **MMBI** HeliTransair

HeliTransair

AWI

### **Participating Institutions**

AWI: Alfred-Wegener-Institut for Polar and Marine Research

Columbusstrasse 27568 Bremerhaven

Germany

IORAS P.P. Shirshov Institute of Oceanology

The Russian Academy of Sciences

36, Nakhimovsky Prospekt

Moscow 117997

Russia

KIGAM Korean Institute of Geoscience and Mineral Resources

30 Gajeong-dong, Yuseong-gu

305-350 Daejeon

Korea

MMBI Murmansk Marine Biological Institute

Vladimirskaya 17 Murmansk 183010

Russia

MPI Max Planck Institute Seewiesen

82319 Seewiesen

Germany

## Ship's Crew

01. 02. 03. 04. 05. 06. 07. 08. 09. 10. 11. 12. 13. 14. 15. 16. 17. 18. 19. 20.	Domke, Udo Grundmann, Uwe Farysch, Bernd Hartung, René Peine, Lutz Fallei, Holger Grigoleit, Urte Riess Delff, Wolfgang Ziemann, Olaf Kotnik, Herbert Muhle, Heiko Nasis, Ilias Verhoeven, Roger Muhle, Helmut Kahrs, Thomas Loidl, Reiner Reise, Lutz NN NN	Master 1.Offc. Ch. Eng. 2.Offc. 2.Offc. Doctor R.Offc. 1.Eng. 2.Eng. 2.Eng. Electr. FielaxElo FielaxElo FielaxElo FielaxElo FielaxElo Carpenter A.B. A.B.
21.	Winkler, Michael	A.B.
22.	Guse, Hartmut	A.B.
23.	Hagemann, Manfred	A.B.
24.	Schmidt, Uwe	A.B.
25.	Vehlow, Ringo	A.B.
26.	Bäcker, Andreas	A.B.
27.	Preußner, Jörg	Storek.
28.	Ipsen, Michael	Mot-man
29.	Voy, Bernd	Mot-man
30.	Elsner, Klaus	Mot-man
31.	Hartmann,Ernst-Uwe	Mot-man
32.	Grafe, Jens	Mot-man Cook
33. 34.	Müller-Homburg, Ralf-Dieter Völske, Thomas	Cooksmate
35.	Silinski, Frank	Cooksmate
36.	Jürgens, Monika	1.Stwdess
37.	Wöckener, Martina	Stwdss/KS
38.	Czyborra, Bärbel	2.Stwdess
39.	Silinski, Carmen	2.Stwdess
40.	Gaude, Hans-Jürgen	2.Steward
41.	Möller, Wolfgang	2.Steward
42.	Huang, Wu-Mei	2.Steward
43.	Yu, Chung Leung	Laundrym.
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