

27. DANIAN CALCAREOUS NANNOFOSSIL SUCCESSION AT SITE 738 IN THE SOUTHERN INDIAN OCEAN¹

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ABSTRACT

An apparently complete Danian section was recovered at ODP Site 738 on the southern Kerguelen Plateau. Calcareous nannofossils are abundant and moderately preserved in the section. A number of taxa common in middle or low latitudes, such as *Braarudosphaera*, *Biscutum? romeinii*, *Biscutum? parvulum*, *Cyclagelosphaera*, *Octolithus multipus*, and *Toweius petalopus* are absent at Site 738. On the other hand, a bloom of *Hornibrookina* occurs at Site 738 only slightly (15 cm) above the Cretaceous/Tertiary boundary as defined by the iridium peak. Species of *Chiasmolithus* and *Prinsius* are very abundant. This gives the nannofossil assemblages distinct high-latitude characteristics and suggests significant latitudinal thermal gradients in the Danian oceans. A Danian nannofossil zonation for the Antarctic region is proposed, which utilizes traditional markers and several nontraditional markers, i.e., the first occurrences of *Hornibrookina*, *Prinsius martinii*, and *Chiasmolithus bidens*, and the last occurrence of *Hornibrookina teuriensis*. Quantitative analyses of the calcareous nannofossil assemblages from Site 738 reveal four steps of rapid floral changes in the early Danian before relatively stable nannofloral conditions were reached at about 63.8 Ma.

INTRODUCTION

Large amounts of data have rapidly accumulated during the last several years on mass extinctions across the Cretaceous/Tertiary boundary. However, quantitative paleontological data for the entire Danian are still scarce, and examination of the repopulation processes of plankton and benthos in the early Paleocene oceans has yet to be carried out in different oceans, environments (shelf vs. oceanic), and latitudes. Many questions need to be addressed, for example, are high-latitude assemblages different from those in the middle or low latitudes immediately after the Cretaceous/Tertiary boundary and how do the latitudinal differentiations of assemblages change through time? What are the rates of diversification and how do they vary from one biogeographic region to another? Is there an "equilibrium" diversity beyond which diversification cannot progress? Is this equilibrium diversity reached synchronously?

Ocean Drilling Program (ODP) Leg 119 recovered an apparently continuous Cretaceous-Tertiary transition and a biostratigraphically complete Danian section at Site 738 on the southern Kerguelen Plateau, the highest latitude Danian sequence ever recovered in the Indian Ocean (Fig. 1). As complete calcareous Danian sections recovered from southern high latitudes are rare, with the only counterpart being that recovered at Site 690 (Leg 113) on Maud Rise in the Weddell Sea (Pospichal and Wise, 1990), the Danian section at Site 738 is especially valuable, serving as a southern anchor in the Indian Ocean for paleobiogeographic and paleoecological studies, and providing crucial information on the evolution of oceanic biota at high latitudes.

The first objective of this study is to provide calcareous nannofossil biostratigraphic information for the Danian section from Site 738. The second objective is to examine quantitatively calcareous nannofossil species abundance patterns in order to investigate the nature of calcareous nannoplankton evolution and diversification during the first several million years following the Cretaceous/Tertiary boundary extinction event. This should

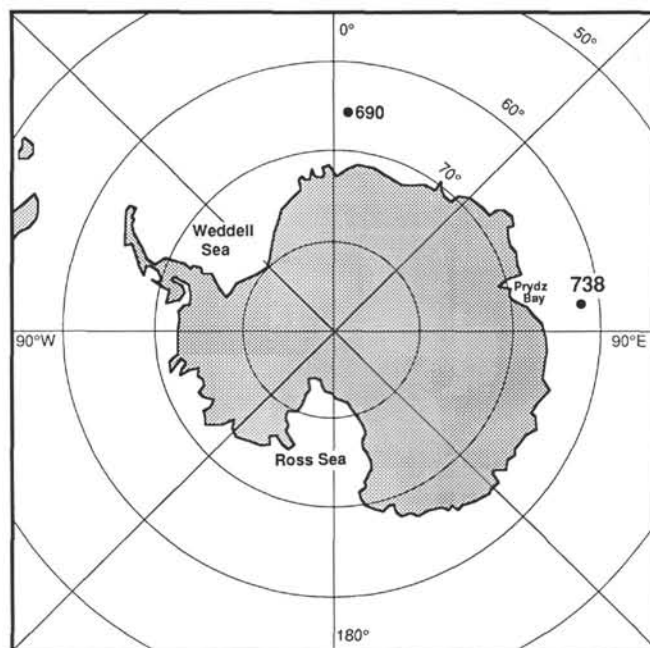


Figure 1. Location map of Site 738 and a comparison site, Site 690 (Leg 113).

shed some light on the paleoenvironment of the southern high latitudes in the early Paleocene. The calcareous nannofossil succession established here will serve as an important reference in the Southern Ocean. This study is the first to quantitatively analyze calcareous nannofossil assemblages of the entire Danian. For additional information on the Cretaceous/Tertiary boundary and the Paleocene calcareous nannofossils at Site 738 see Thierstein et al. (this volume) and Wei and Thierstein (this volume).

Pioneering work on calcareous nannofossils across the Cretaceous/Tertiary boundary interval was carried out by Bramlette and Martini (1964). Subsequently, a number of nannofossil biostratigraphic studies of the Cretaceous/Tertiary boundary and

¹ Barron, J., Larsen, B., et al., 1991. *Proc. ODP, Sci. Results*, 119: College Station, TX (Ocean Drilling Program).

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the early Paleocene have been completed, e.g., Edwards (1966) in New Zealand, Hay and Mohler (1967) in France, Perch-Nielsen (1969) in Denmark, Worsley (1974) in Alabama, Percival and Fisher (1977) in northern Spain, Romein (1977, 1979) in southern Spain, Thierstein and Okada (1979) in the North Atlantic, Perch-Nielsen (1981a) in Tunisia, Percival (1984) in the South Atlantic, Monechi (1985) in the North Pacific, Jiang and Gartner (1986) in Texas, Heck and Prins (1987) in the North Sea, and Varol (1989) in Turkey. Reports of southern high-latitude Cretaceous/Tertiary boundary and Danian nannofossils include Edwards (1973), and Pospichal and Wise (1990). Comprehensive reviews of calcareous nannofossil stratigraphy and assemblages across Cretaceous/Tertiary boundaries in various parts of the world were presented by Thierstein (1981), Perch-Nielsen et al. (1982), and Smit and Romein (1985). Haq and Lohmann (1976) and Haq et al. (1977, 1979) carried out quantitative biogeographic studies of calcareous nannofossils for selected time slices of the Danian. They reported that there was a latitudinal differentiation of nannofossil assemblages and that there were major latitudinal migrations of these assemblages as a response to major changes in climate.

MATERIAL AND METHODS

ODP Site 738 is located at 62°42.543'S, 82°47.245'E, in 2252 m of water. The site lies on the southern part of the Kerguelen Plateau, north of the modern Antarctic Divergence and near the southern limit of influence of the Antarctic Circumpolar Current. Site 738 was probably near sea level at 90 Ma based on the shallow marine and volcanoclastic rocks recovered from the basement, and represented an outer shelf environment (400–600 m) in the late Maestrichtian according to benthic foraminifers (Barron, Larsen, et al., 1989). During the Danian, Site 738 was probably in about 1000 m of water assuming a uniform rate of subsidence.

The Danian section at Site 738 consists of white to greenish-gray calcareous chalk with chert nodules and fragments. The chalk is partially silicified. Major components of the sediment include micrite, calcareous nannofossils, foraminifers, and clay. The bottom of the Danian (Section 119-738C-20R-5, 96 cm, to -20R-5, 83 cm) is laminated claystone, with distinct laminae on a millimeter scale. An iridium peak occurs at the base of this laminated section (Schmitz, this volume). For detailed description of the section, we refer the readers to Barron, Larsen, et al. (1989).

Samples used in this study were taken from Cores 119-738C-17R to -20R at about 50-cm intervals, which represent an average sampling frequency of about 0.05 m.y.. Smear slides were made directly from the unprocessed samples to avoid losing smaller-sized specimens. Scanning electron microscope (SEM) stubs were also made directly from the unprocessed samples. The SEM was used to identify the small nannofossil species in about a dozen samples, mostly from the lower Danian.

The relative abundances of the calcareous nannofossil species were first estimated with a light microscope at a magnification of about $\times 1600$ and recorded as follows: V = very abundant, more than 10 specimens per field of view; A = abundant, 1–10 specimens per field of view; C = common, 1 specimen per 2–10 fields of view; F = few, 1 specimen per 11–50 fields of view; R = rare, 1 specimen per 51–300 fields of view; N = not found in 300 fields of view. These semiquantitative data are used to construct the range chart from which nannofossil biostratigraphy is established. Preservation of assemblages is indicated by G = good, little overgrowth and/or etching; M = moderate, overgrowth and/or etching are apparent, but identification at species level is generally not impaired; P = poor, significant overgrowth and/or etching, identification of some species is impaired.

Calcareous nannofossil species recorded in this study are listed in Table 1, and most of these species are documented with SEM and light micrographs in Plates 1 through 4 of this paper. Bibliographic references for these species can be found in Loeblich and Tappan (1966, 1968, 1969, 1970a, 1970b, 1971, 1973), van Heck (1979a, 1979b, 1980a, 1980b, 1981a, 1981b, 1982a, 1982b, 1983), or Steinmetz (1985a, 1985b, 1986, 1987a, 1987b, 1988a, 1988b, 1989).

For the quantitative study of the nannofossil assemblages, about 300 specimens of calcareous nannofossils were counted along random traverses of each slide. Rare species generally are not encountered in these first 300 counts. This does not hinder the usefulness and interpretation of the data, because our primary goal here is to examine major components in the samples, which significantly affect the population dynamics. Paleoenvironmental information is extracted from these major components. Species less than a few percent or abundance change less than a few percent are not considered statistically significant in this study.

In order to make it practical to count the species for large numbers of samples and to obtain consistent results, it was necessary to group some taxa. We counted *Chiasmolithus bidens* and *Chiasmolithus danicus* as a group during the collection of the quantitative data. These two species differ by the presence of axial lines in the cross bars of the former species and the two

Table 1. Calcareous nannofossils considered in this paper. In alphabetical order of generic epithets; plate and figure numbers refer to illustrations in this paper.

<i>Ahmuelerella octoradiata</i> (Gorka) Bukry, 1969
<i>Arkhangelskiella cymbiformis</i> Veskshina, 1959
<i>Biscutum constans</i> (Gorka) Black, 1967. Pl. 1, Fig. 9; Pl. 2, Fig. 2
<i>Biscutum? parvulum</i> Romein, 1979
<i>Biscutum? romeinii</i> Perch-Nielsen, 1981
<i>Chiasmolithus bidens</i> (Bramlette and Sullivan) Hay and Mohler, 1967. Pl. 4, Figs. 5, 6
<i>Chiasmolithus danicus</i> (Brotzen) Hay and Mohler, 1967. Pl. 4, Figs. 3, 4, 7
<i>Coccolithus pelagicus</i> (Wallich) Schiller, 1930. Pl. 4, Fig. 8
<i>Coccolithus subpertusus</i> (Hay and Mohler) Wei and Pospichal, n. comb.
<i>Cretarhabdus conicus</i> Bramlette and Martini, 1964
<i>Cretarhabdus surirellus</i> (Deflandre) Reinhardt, 1970
<i>Cribrosphaerella daniae</i> Perch-Nielsen, 1977
<i>Cribrosphaerella ehrenbergii</i> (Arkhangelsky) Deflandre, 1952. Pl. 4, Fig. 10
<i>Cruciplacolithus edwardsii</i> Romein, 1979. Pl. 3, Figs. 4, 7
<i>Cruciplacolithus primus</i> Perch-Nielsen, 1977
<i>Cruciplacolithus tenuis</i> (Stradner) Hay and Mohler in Hay et al., 1967
<i>Cyclagelosphaera reinhardtii</i> (Perch-Nielsen) Romein, 1977
<i>Eiffellithus turrisseiffelii</i> (Deflandre) Reinhardt, 1965
<i>Fasciculithus tympaniformis</i> Hay and Mohler in Hay et al., 1967
<i>Fasciculithus ulii</i> Perch-Nielsen, 1971
<i>Gartnerago obliquum</i> (Stradner) Noël, 1970
<i>Hornibrookina teuriensis</i> Edwards, 1973. Pl. 2, Figs. 5–7
<i>Kamptnerius magnificus</i> Deflandre, 1959
<i>Lithraphidites carniolensis</i> Deflandre, 1963
<i>Lucianorhabdus cayeuxii</i> Deflandre, 1959
<i>Markalius apertus</i> Perch-Nielsen, 1979. Pl. 1, Fig. 11
<i>Markalius inversus</i> (Deflandre) Bramlette and Martini, 1964. Pl. 1, Fig. 10; Pl. 2, Figs. 3, 4; Pl. 4, Fig. 2
<i>Micula staurophora</i> (Gardet) Stradner, 1963
<i>Neochiastozygus concinnus</i> (Martini) Perch-Nielsen, 1971. Pl. 1, Fig. 7
<i>Neochiastozygus modestus</i> Perch-Nielsen, 1971
<i>Nephrolithus frequens</i> Gorka, 1957
<i>Octolithus multiplus</i> (Perch-Nielsen) Romein, 1979
<i>Prediscosphaera cretacea</i> (Arkhangelsky) Gartner, 1968
<i>Prediscosphaera spinosa</i> (Bramlette and Martini) Gartner, 1968
<i>Prediscosphaera stoveri</i> (Perch-Nielsen) Shaff and Stradner, 1971
<i>Prinsius bisulcus</i> (Stradner) Hay and Mohler, 1967. Pl. 3, Figs. 1–3, 8
<i>Prinsius dimorphosus</i> (Perch-Nielsen) Perch-Nielsen, 1977. Pl. 1, Figs. 1–6, 8
<i>Prinsius martinii</i> (Perch-Nielsen) Haq, 1977. Pl. 3, Fig. 1
<i>Reinhardtites anthophorus</i> (Deflandre) Perch-Nielsen, 1968
<i>Toweius pertusus</i> (Sullivan) Romein, 1979. Pl. 3, Figs. 5, 6, 9; Pl. 4, Fig. 1
<i>Zygodiscus sigmoides</i> Bramlette and Sullivan, 1961. Pl. 2, Figs. 2, 9
<i>Zygodiscus spiralis</i> Bramlette and Martini, 1964

species overlap in stratigraphic range (see for example, Perch-Nielsen, 1985). Where the two taxa overlap in range, it was difficult to count the two species separately because the cross bars of many specimens are not well preserved to allow the identification of axial lines. However, it was possible to determine the first occurrence of *C. bidens* based on a few well preserved specimens which show axial lines in the cross bars.

Similarly, all *Crucioplacolithus* were counted as a group. This group consists predominantly of *C. tenuis* and *C. primus* with few specimens of *C. edwardsii*. Relative abundances of these species can be estimated according to the relatively well preserved specimens, and the relative abundances are recorded in the range chart.

A broad species concept of *Coccolithus pelagicus* is used in this study. It may include *C. pelagicus* (Wallich) Schiller, *C. cavus* Hay and Mohler, and *C. eoplagicus* (Bramlette and Riedel) Bramlette and Sullivan. It appears that all these species are synonymous.

Prinsius bisulcus and *Prinsius martinii* differ mainly by size, but the size difference has never been well defined. Where the two species co-occur, it is impossible to differentiate them consistently. Consequently, these two taxa were treated as one during the collection of the census data.

All species of *Hornibrookina* were counted as a group. This includes *H. edwardsii* Perch-Nielsen and *H. teuriensis* Edwards. *Hornibrookina edwardsii* is taken as a junior synonym of *H. teuriensis* (see taxonomic note below for detailed discussion). No *Hornibrookina australis* was found at Site 738.

The abundance of *Thoracosphaera* was not counted but estimated roughly and recorded in the range chart. The reason is that most specimens are broken, and it was not practical to count broken pieces when it was difficult to determine what percentage of the whole specimens they represent.

Due to small size (2–4 μm) and poor preservation of *Prinsius dimorphosus*, this species was not counted. Only semiquantitative abundance estimates were recorded on the range chart.

All Cretaceous species were counted as one group, because individual species are rare and sporadic in the Danian section. These species are most likely to be redeposited, as discussed in the following text.

Because grouping and deletion of some species were necessary in this study, the limitations of the census data should be recognized and the interpretation of the results should take these factors into account.

Polar ordination analysis was made to determine how the assemblages change in relation to an axis defined by two end-member samples, one from the top of the section (Sample 119-738C-17R-1, 15–16 cm, 341.45 mbsf), and the other one from the bottom of the Danian section (119-738C-20R-5, 35–36 cm, 376.55 mbsf). Polar ordination is an axial method for ordinating samples according to their similarity values along an axis defined by two end-member samples. A detailed description of this method has been given by Cisne and Rabe (1978).

In order to examine how the nanoplankton diversity changes through the Danian at Site 738, diversity indices were calculated for each sample. The measure used corresponds to the entropy concept of information theory proposed by Shannon (1949). In a paleoecological context, this measure is often referred to as the Shannon-Weaver diversity index (H), which is calculated as

$$H = - \sum P_i (\log P_i)$$

where P_i is the probability of the occurrence of the i^{th} species. Since P_i is a parameter of the population and not the sample, it must be estimated by (n_i/n) , where n_i is the number of specimens of the i^{th} species, and n is the total number of specimens counted in that sample.

STRATIGRAPHIC DISTRIBUTION PATTERNS OF CALCAREOUS NANNOFOSSILS

Stratigraphic distribution patterns of calcareous nannofossils from Sample 119-738C-20R-CC (379.9 mbsf) to Sample 119-738C-17R-1, 15 cm (341.15 mbsf), are presented in Table 2. Samples 119-738C-20R-5, 96.1 cm, and -20R-CC contain only Mesozoic calcareous nannofossil species, including the last Mesozoic zonal marker, *Nephrolithus frequens*, but no Cenozoic species, and should be placed in the uppermost Maestrichtian Zone NC23 of Roth (1978). A sharp peak of iridium concentration has been found in Sample 119-738C-20R-5, 96.0–96.2 cm (377.16 mbsf) (Schmitz, this volume), and this level is taken as the Cretaceous/Tertiary boundary. This chemically defined Cretaceous/Tertiary boundary is consistent with the nannofossil data, because the first Tertiary nannofossil marker, *Biantholithus sparsus*, was found only 13 cm above.

The next easily recognized datum is the first occurrence of *Hornibrookina* in Sample 119-738C-20R-5, 82.0 cm (377.02 mbsf). This genus increases in abundance from few to abundant in less than 50 cm after its first appearance. The first *Crucioplacolithus tenuis* was found in Sample 119-738C-20R-5, 2 cm (376.22 mbsf). Abundant *Prinsius dimorphosus* were also first found in this sample. However, no *Biscutum? romeinii*, *Biscutum parvulum* or *Toweius petalosus* were found in this hole; thus the subdivision of NP1 as prescribed by Perch-Nielsen (1981) is not possible at Site 738. The first *Coccolithus pelagicus* was found in Sample 119-738C-20R-4, 117 (375.87) cm. The first occurrence (FO) of this species was found at the same level as the FO of *Prinsius dimorphosus* in Tunisia (Perch-Nielsen, 1981a), one sample lower than the latter species event in Spain (Romein, 1979), and at Site 577 (Leg 86) in the northwest Pacific (Monechi, 1985), one-half zone lower than the latter datum in Turkey (Varol, 1989). As at Site 738, *C. pelagicus* occurs stratigraphically higher than *P. dimorphosus* at Site 690 in the Weddell Sea (Pospichal and Wise, 1990). This may suggest that *C. pelagicus* preferred warmer waters than *P. dimorphosus* and that the latitudinal temperature gradient in the early Paleocene caused the latitudinal diachroneity of the first occurrence of either one or both species.

The FO of *Chiasmolithus danicus* was observed in Sample 119-738C-19R-3, 85 cm (366.35 mbsf). According to Heck and Prins (1987), *C. danicus* evolved from *Crucioplacolithus edwardsii*, and the former species is only slightly younger than the latter. This sequence was also observed at Site 738 as well as at Site 690 in the Weddell Sea (Pospichal and Wise, 1990).

The FO of *Prinsius martinii* was located in Sample 119-738C-18R-CC (360.50 mbsf). This datum was easy to pick because the species is abundant even at the beginning of its stratigraphic range. The first *Toweius pertusus* was also found in the same sample, and this species event may serve as a provisional secondary marker. The traditional marker species, *Ellipsolithus macellus*, was not found at Site 738. This is not surprising because the species is quite delicate and is usually not found in deep-sea sediments.

The next species event is the FO of *Chiasmolithus bidens* in Sample 119-738C-18R-1, 35 cm (351.25 mbsf). The first *Fasciculolithus tympaniformis* and *Prinsius bisulcus* were found in the next higher sample (Sample 119-738C-17R-CC). *Fasciculolithus tympaniformis*, as well as other species of *Fasciculolithus*, are rare and sporadic in this part of the section at Site 738, thus making its application as a marker in the Antarctic region difficult.

Rare *Hornibrookina teuriensis* were present up to Sample 119-738C-16R-5, 97–98 cm, whereas *Helolithus kleinpellii* was not found until Sample 119-738C-15R-CC (see Wei and Thierstein, this volume). The LO of *H. teuriensis* is used as a zonal

Table 2. Distribution of calcareous nannofossils, Hole 738C.

Age	Zones of this study	Core, section, interval (cm)	Depth (mbsf)	Abundance	Preservation	<i>Biantholithus sparsus</i>	<i>Chiasmolithus bidens</i>	<i>Chiasmolithus danicus</i>	<i>Coccolithus pelagicus</i>	<i>Coccolithus supertus</i>	<i>Cruciplacolithus edwardsii</i>	<i>Cruciplacolithus primus</i>	<i>Cruciplacolithus subrotundus</i>	<i>Cruciplacolithus tenuis</i>	<i>Fasciculolithus tympaniformis</i>	<i>Fasciculolithus ulii</i>	<i>Hornibrookina teuriensis</i>	<i>Lapideacasis</i> sp.	<i>Markalius apertus</i>	<i>Markalius inversus</i>	<i>Neochiastozygus concinnus</i>	<i>Neochiastozygus modestus</i>
Selandian	NA6	17-1, 15	341.45	A	M	A											R		R			
		17-1, 66	341.96	A	M	A											R		R			
		17-1, 115	342.45	A	M	A			R			R					R		R			
		17-2, 15	342.95	A	M	C			R								R		R			
		17-2, 66	343.46	A	M	A			R								R		R			
		17-2, 115	343.95	A	M	A									R		R		R			
		17-3, 15	344.45	A	M	A											R		R			
		17-3, 66	344.96	A	M	A											R		R			
		17-3, 115	345.45	A	M	A											R		R			
		17-4, 15	345.95	A	M	A				R									R			
		17-4, 66	346.46	A	M	A				C									R			
		17-4, 115	346.95	A	M	C											R		R			
		17-5, 15	347.45	A	M	A			R								R		R			
		17-5, 66	347.96	A	M	A											R		R			
		17-5, 115	348.45	A	M	A			R	R					R		R		R			
		17-6, 15	348.95	A	M	A											R		R			
		17-6, 66	349.46	A	M	A									R		R		R			
		17-6, 115	349.95	A	M	A											R		R			
		17-7, 15	350.45	A	M	A											R		R			
		17-CC	350.90	A	M	C					F				R		R		R		F	
Danian	NA5	18-1, 35	351.25	A	M	C		C	R								R		R		F	
		18-1, 85	351.75	A	M	A		A	R								R		R		R	
		18-2, 35	352.75	A	M	A		A	R								R		R		R	
		18-2, 85	353.25	A	M	A		A	F								R		R		R	
		18-2, 135	353.75	A	M	A		A	F								R		R		R	
		18-3, 35	354.25	A	M	A		A	F								R		R		R	
		18-3, 85	354.75	A	M	A		A	R								R		R		R	
		18-3, 135	355.25	A	M	A		A	R								R		R		R	
		18-4, 35	355.75	A	M	C		C	R		R						R		R		R	
		18-4, 135	356.75	A	M	C		C	R		R						R		R		R	
		18-5, 35	357.25	A	M	C		C	R		R						R		R		R	
		18-5, 85	357.75	A	M	C		C	R		R						R		R		R	
		18-5, 135	358.25	A	M	C		C	F		R						R		R		R	
		18-6, 30	358.70	A	M	C		C	F		R						R		R		R	
		18-6, 85	359.25	A	M	C		C	F		R						R		R		R	
		18-6, 135	359.75	A	M	C		C	F		R						R		R		R	
	NA4	18-7, 35	360.25	A	M	F		C	F								R		R		R	
		18-CC	360.50	A	G	C		C							F		R		R		R	
		19-1, 35	360.85	A	M	A		A	F		R				F		R		R		R	
		19-1, 85	361.35	A	G	A		A	F						F		R		R		R	
		19-2, 35	362.35	A	M	A		A	F		F				F		R		R		R	
		19-2, 85	362.85	A	M	A		C	F		R				F		R		R		R	
		19-2, 135	363.35	A	M	C		C	R		R				F		R		R		R	
		19-3, 35	363.85	A	G	C		C	R		R				F		R		R		R	
	NA3	19-3, 85	364.35	A	G	C		C	F						C		R		R		R	
		19-3, 135	364.85	A	G	C		C	A		F				A		R		R		R	
		19-CC	370.20	A	G	A		A	R		R				A		R		R		R	
		20-1, 46	370.66	A	G	A		A	R		R				A		R		R		R	
		20-1, 98	371.18	A	G	A		A	R		R				A		R		R		R	
		20-2, 42	372.12	A	G	A		A	R		R				A		R		R		R	
		20-2, 138	373.08	A	G	A		A	R		R				A		R		R		R	
		20-3, 17	373.37	A	G	A		A	R		R				A		R		R		R	
		20-3, 70	373.90	A	G	A		A	R		R				A		R		R		R	
		20-3, 137	374.57	A	G	A		A	R		R				A		R		R		R	
		20-4, 20	374.90	A	M	A		A	R		R				A		R		R		R	
		20-4, 117	375.87	A	M	A		A	R		R				A		R		R		R	
	NA2	20-5, 2	376.22	A	M	A		A	R		R			R	C		R		R		R	
		20-5, 35	376.55	A	M	A		A	R		R			R	C		R		R		R	
	NA1	20-5, 82.0	377.02	R	P	R		R						R	A		R		R		R	
		20-5, 83.5	377.04	R	P	R		R						R	A		R		R		R	
		20-5, 95.9	377.16	C	P	R		R						r			R		R		R	
Maestrichtian	NC23	20-5, 96.1	377.16	F	P																	
		20-CC	379.90	C	M																	

Note: Abundance is characterized: V, very abundant; A, abundant; C, common; F, few; R, rare. For preservation, G, good; M, moderate; P, poor. Lower-case letters indicate reworked specimens.

marker in this study. The zonal scheme is discussed in detailed below.

ZONATION

A number of calcareous nannofossil zonations for the Danian have been proposed since the detailed work of Hay and Mohler (1967). The most commonly used zonations are summarized in Figure 2. However, none of the previous zonations can achieve a satisfactory biostratigraphic resolution at Site 738 because of the rarity or absence of some conventional marker species (i.e., *Ellipsolithus macellus*, *Fasciculithus tympaniformis*) at this extreme high-latitude site. On the other hand, some nonconventional datums (i.e., the FO's of *Hornibrookina*, *Prinsius martinii*, and *Chiasmolithus bidens*) are useful for southern high-latitude biostratigraphy. Based on data from Site 738 (Table 2 and Fig. 3) and other Southern Ocean sites, particularly Site 690 in the Weddell Sea (Fig. 4), a calcareous nannofossil zonation for the Antarctic region is proposed. Each zone is prefixed by "NA" (short for "nannofossil Antarctic") and numbered from the bottom of the Danian upward. The zonation is summarized in Figure 2. The interpretation for the magnetostratigraphy of Hole 738C is poor at this moment due to large sampling gaps (Sakai and Keating, this volume). Future resampling of the cores for paleomagnetic measurement should produce a very useful magnetostratigraphy for this relatively complete high-latitude Danian sequence, and the nannofossil zones established in the present study can then be well correlated to the geomagnetic time scale. Detailed descriptions of the nannofossil zones proposed follow.

Biantholithus sparsus Zone (NA1)

Base. FO of *Biantholithus sparsus* or the iridium peak.

Top. FO of *Hornibrookina*.

Author. Perch-Nielsen (1979), emended by Wei and Pospichal (this paper).

Remarks. The base of this zone corresponds to the Cretaceous/Tertiary boundary. Martini (1971) defined the Cretaceous/Tertiary boundary by the last occurrence of Cretaceous forms. This criterion usually cannot be successfully applied because of reworking of Cretaceous assemblages into the lower Danian. Edwards (1973) proposed to use the FO of *Conococcolithus panis* to mark the Cretaceous/Tertiary boundary. This species is, however, a junior synonym of *Biscutum constans* (see Romein, 1977, p. 260, for detailed discussion). The latter species is well known to occur in the Cretaceous. Both Romein (1979) and Perch-Nielsen (1981a) used the beginning of *Thoracosphaera* acme to delimit the Cretaceous/Tertiary boundary. Though Perch-Nielsen et al. (1982) claimed that the *Thoracosphaera* acme is universally related to the base of the Tertiary, this is not the case in some areas (e.g., Monechi and Thierstein, 1985; Pospichal and Wise, 1990; this study). Perch-Nielsen (1979) used the FO of *Biantholithus sparsus* to define the Cretaceous/Tertiary boundary. *Biantholithus sparsus* is the first true Tertiary species. The shortcoming of this marker for the Cretaceous/Tertiary boundary is that the species, as its name indicates, is generally very sparse. Consequently, it is difficult to determine precisely its evolutionary first appearance, and the boundary drawn by the FO of this species is often slightly higher than the chemically defined Cretaceous/Tertiary boundary. The iridium peak, if found, near and below the FO of *B. sparsus*, represents a more precise location of the Cretaceous/Tertiary boundary.

Hornibrookina first occurs just slightly above (14 cm) the Cretaceous/Tertiary boundary as defined by the iridium peak (Schmitz, this volume) at Site 738. This genus increases its abundance from few at its FO level to abundant less than 50 cm uphole. A similar phenomenon was also observed at Site 690 on Maud Rise, also an extreme southern high-latitude site (65°S). With the quest for finer stratigraphic resolution of the Cretaceous/Tertiary boundary interval, the FO of *Hornibrookina* offers an easily picked and useful datum in the southern high latitudes. Assuming a uniform sedimentation rate for the interval be-

Age	Martini, 1971 (general)	Okada and Bukry, 1980 (low latitude)	Romein, 1979 (Spain)	Perch-Nielsen, 1979 (North Sea)	Perch-Nielsen, 1981b (Tunisia)	This paper (Antarctic)
Paleocene	late Selandian	<i>H. kleinpellii</i>	<i>H. kleinpellii</i>	S2	<i>H. kleinpellii</i>	<i>H. teuriensis</i>
		NP5	CP4	<i>F. tympaniformis</i>	<i>F. tympaniformis</i>	
	early Danian	<i>F. tympaniformis</i>	<i>F. tympaniformis</i>	S1	<i>F. tympaniformis</i>	NA6
				D10		
		NP4	CP3	<i>E. macellus</i>	<i>E. macellus</i>	<i>C. bidens</i>
				D9		NA5
		<i>E. macellus</i>	<i>E. macellus</i>	<i>N. saepes</i>	<i>E. macellus</i>	<i>P. martinii</i>
		NP3	CP2	D8	<i>C. edwardsii</i>	NA4
				D7		
			<i>C. tenuis</i> s. str.	D6		
			<i>P. dimorphosus</i>	D5		
		<i>C. danicus</i> s.l.	<i>P. dimorphosus</i>	D4	<i>C. edwardsii</i>	<i>C. danicus</i>
Cret.	Maest.	NP2	CP1b	<i>C. primus</i>	<i>C. primus</i>	NA3
				D3		
		<i>C. tenuis</i>	<i>C. tenuis</i>	<i>C. tenuis</i>	<i>C. primus</i> (large)	<i>C. tenuis</i>
		NP1	CP1a	D2	<i>T. petalosus</i>	NA2
				D1	<i>C. ultimus</i>	
Cret.	Maest.	<i>A. cymbiformis</i> & Cretaceous forms	<i>M. mura</i> & Cretaceous forms	<i>B. sparsus</i>	<i>B. ? parvulum</i>	<i>Hornibrookina</i>
		<i>M. murus</i>	<i>M. murus</i>	<i>M. prinsii</i>	<i>B. ? romeinil</i>	<i>B. sparsus</i> or the iridium peak

Figure 2. Correlation of nannofossil zonations of various authors (after Perch-Nielsen, 1985).

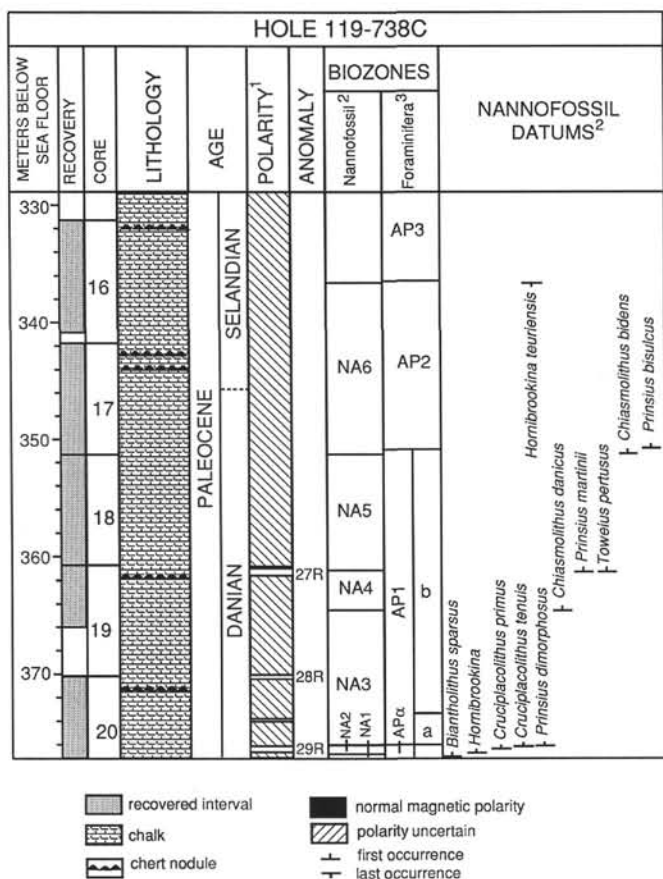


Figure 3. Lithostratigraphy and magnetobiostratigraphy of the Danian and lower Selandian sequence recovered in Hole 738C. Data are taken from: ¹ Sakai and Keating (this volume); ² Wei and Pospichal (this paper); ³ Huber (this volume).

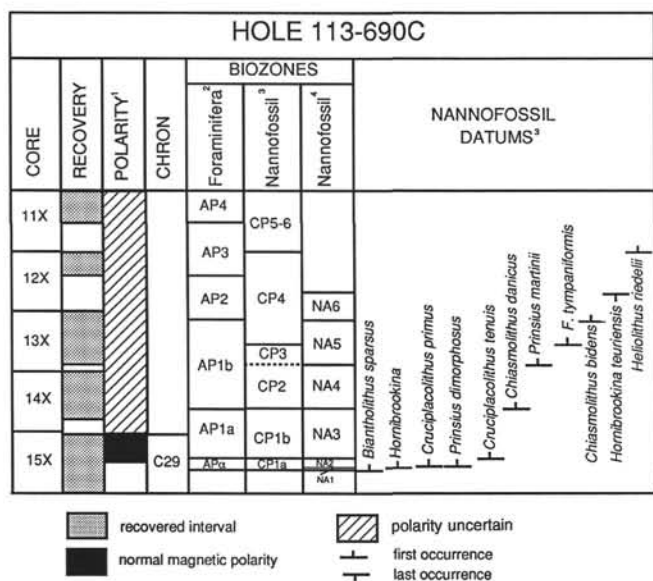


Figure 4. Magnetobiostratigraphy of the Danian and lower Selandian sequence recovered in Hole 690C, a comparison hole for Hole 738C. Data are taken from: ¹ Hamilton (1990); ² Stott and Kennett (1990); ³ Pospichal and Wise (1990); also, ³ Rezoned in the Antarctic zonation proposed in this paper.

tween the iridium peak (377.16 mbsf) and the FO of *Cruciplacolithus tenuis* (376.38 mbsf), and an age of 66.4 Ma for the iridium peak and 66.2 Ma for the FO of *C. tenuis* (Monechi and Thierstein, 1985) at Site 738, the FO of *Hornibrookina* would be 66.37 Ma.

Hornibrookina teuriensis Zone (NA2)

Top. FO of *Cruciplacolithus tenuis*

Base. FO of *Hornibrookina*

Author. Wei and Pospichal (this paper).

Remarks. This zone comprises most of Zone NP1 of Martini (1971) or Subzone CP1a of Okada and Bukry (1980). Perch-Nielsen (1981) subdivided Zone NP1 into five zones, using the FO's of *Biscutum? romeinii*, *Biscutum? parvulum*, *Cruciplacolithus primus* (small), and *Towelius petalosus*. Except for *C. primus*, none of these species were found at Site 738, probably due to exclusion of these species by low water temperatures at this Antarctic site. *Cruciplacolithus primus* first occurs within this zone. The FO of *Prinsius dimorphosus* can be used as a secondary marker to approximate that of *C. tenuis*.

Magnetobiostratigraphic data (Sakai and Keating, this volume) reveal that the FO of *C. tenuis* is within Subchron 29R at Site 738, a similar situation as in the Contessa section in Italy (Monechi and Thierstein, 1985) but different from results of other sites (Subchron 29N, see Berggren et al., 1985). An age of 66.2 Ma (Monechi and Thierstein, 1985) is tentatively used in this study for the FO of *C. tenuis*.

Cruciplacolithus tenuis Zone (NA3)

Top. FO of *Chiasmolithus danicus*

Base. FO of *Cruciplacolithus tenuis*

Author. Mohler and Hay in Hay et al. (1967), emend. Martini (1970).

Remarks. The FO of *C. danicus* has an age of 64.8 Ma at several DSDP sites (Berggren et al., 1985).

Chiasmolithus danicus Zone (NA4)

Top. FO of *Prinsius martinii*

Base. FO of *Chiasmolithus danicus*

Author. Martini (1970), emend. Edwards (1971).

Remarks. The base of this zone is the same as that of Zone NP3 of Martini (1971) or Zone CP2 of Okada and Bukry (1980). The top of the zone is different from the latter two zonations, where the FO of *Ellipsolithus macellus* is used instead. *Ellipsolithus macellus* is absent at Site 738, as at Site 690 in the Weddell Sea. This delicate species is usually not present in deep-sea sediments. Further more, the FO of *E. macellus* is quite diachronous in different regions (Wei and Wise, 1989). On the other hand, *Prinsius martinii* is rather cosmopolitan. Though many authors do not differentiate *Prinsius bisulcus* from *P. martinii* (e.g., Perch-Nielsen, 1977; Okada and Thierstein, 1979; Percival, 1984; Monechi, 1985), the first occurrence of the *P. bisulcus*—*P. martinii* group has been consistently found between the FO of *C. danicus* and the FO of *Fasciculolithus tympaniformis*. The FO of *P. martinii* has an age of 63.8 Ma in the Contessa section in Italy (Monechi and Thierstein, 1985).

Prinsius martinii Zone (NA5)

Top. FO of *Chiasmolithus bidens*

Base. FO of *Prinsius martinii*

Author. Edwards (1971), emend. Wei and Pospichal (this paper).

Remarks. In the high latitudes, the FO of *C. bidens* is closely associated with the FO of *Prinsius bisulcus* (> 5 μ m) and the LO of *Prinsius dimorphosus*. The latter two datums may serve as secondary markers. The FO of *C. bidens* has an age of about 63.2 Ma at DSDP Site 516 (Wei and Wise, 1989).

Chiasmolithus bidens Zone (NA6)

Top. LO of *Hornibrookina teuriensis*

Base. FO of *Chiasmolithus bidens*

Author. Wei and Pospichal (this paper).

Remarks. The LO of *H. teuriensis* has been found consistently to fall between the FO of *Fasciculolithus tympaniformis* and the FO of *Helioolithus klempellii* at a number of localities, including the southern Kerguelen Plateau (this study), the Weddell Sea (Pospichal and Wise, 1990), the southwest Atlantic (Perch-Nielsen, 1977), and Turkey (Varol, 1989). It seems that the LO of *H. teuriensis* has a relatively consistent age over large areas. On the other hand, *F. tympaniformis* is rather rare and sporadic in the lower part of its stratigraphic range at Site 738, making the

precise location of the FO of this species impossible. Consequently, the top of Zone NP4 of Martini (1971) or Zone CP3 of Okada and Bukry (1980) is difficult to determine.

RESULTS OF QUANTITATIVE ANALYSES

The relative abundances of individual taxa in percent are given in the Appendix and are graphically presented in Figure 5. Cretaceous species drop from 100% below the Cretaceous/Tertiary boundary to less than 10% of the total assemblage 4 m above the boundary. Individual Cretaceous taxa are sporadically present uphole. Based on these distribution patterns, we consider these specimens of Cretaceous species to be reworked. Elsewhere, Cretaceous species in Danian sediments are considered to have survived the Cretaceous/Tertiary crisis (e.g., Perch-Nielsen et al., 1982; Smit and Romein, 1985). The evidence for that interpretation came from stable isotope studies of Perch-Nielsen et al. (1982). These authors measured oxygen and carbon isotopes of bulk samples from the uppermost Cretaceous and the lowermost Tertiary, and found strongly differing signals across the Cretaceous/Tertiary boundary. They believed that Cretaceous coccoliths constitute the bulk of the calcite measured, and thus concluded that the Cretaceous coccoliths in the basal Danian actually lived in the Tertiary. However, the bulk samples from the lowermost Danian usually contain a large amount of micrite, juvenile foraminifers, and pieces of *Thoracosphaera*, and the isotopic signals from these would overshadow those of Cretaceous nannofossil species in the lowermost Danian. Thus, the isotopic results from bulk-sample measurements by Perch-Nielsen et al. (1982) do not offer unambiguous evidence that the Cretaceous nannofossils found in the basal Danian actually lived into the Tertiary.

Cruciplacolithus registers a sudden bloom in Sample 119-738C-20R-5, 2 cm (376.22 mbsf, Fig. 5, arrow 1), and makes up as much as 87% of the assemblage in less than 2 m uphole. This taxon maintains relatively high abundance up to about 360 mbsf, and is virtually absent (sporadic occurrence and less than 1%) above that. *Cruciplacolithus tenuis* is the dominant species in the genus at Site 738, whereas *Cruciplacolithus primus* is considerably rarer compared to low- or middle-latitude sites. Semi-quantitative data (Table 2) reveal that *Prinsius dimorphosus* begins to bloom virtually at the same time as *Cruciplacolithus* (Sample 119-738C-20R-5, 2 cm). This a similar phenomenon found at Site 690 in the Weddell Sea (Pospichal and Wise, 1990). The next significant change in population composition is the abrupt increase (from 5% to 49%) of *Coccolithus pelagicus* in Sample 119-738C-19R-CC (370.20 mbsf, Fig. 5, arrow 2). This species appears earlier but is relatively fewer in numbers (less than 6%). The third large change in population composition is the addition of a new species, *Chiasmolithus danicus*, in Sample 119-738C-19R-3, 35 cm (363.85 mbsf, Fig. 5, arrow 3). The relative abundance of this species reaches up to 77% shortly after it appears. The final large change of population composition is in Sample 119-738C-18R-CC (360.5 mbsf), where *Prinsius martinii* and *Toweius pertusus* first occur (Fig. 5, arrow 4). The former species dominates the assemblage and reaches 85% of the total assemblage.

Figure 6 shows the results of the polar ordination analysis. Sample 119-738C-20R-5, 35–36 cm, from the bottom of the Danian was chosen as one end point and Sample 119-738C-17R-1, 15–16 cm, from the top of the section was selected as the other end point in constructing the ordination axis onto which all samples were projected. The lower the value, the greater the similarity of the sample with Sample 119-738C-20R-5, 35–36 cm, and the greater the dissimilarity of the sample with Sample 119-738C-17R-1, 15–16 cm. Four steps of floral change from the oldest assemblage represented by Sample 119-738C-20R-5, 36–37 cm, to the youngest assemblage in Sample 119-738C-17R-1,

15–16 cm, are shown in Figure 6 (marked by arrows 1 through 4). These four steps are the same as those identified in Figure 5.

Changes of Shannon-Weaver diversity index values through time at Site 738 are presented in Figure 7. As pointed out in the preceding, because of the necessary grouping and deletion of some species in the census data as required by this study, the Shannon-Weaver diversity values given in the figure are conservative. However, it is clear that by about 371 mbsf (Zone NA3), the diversity has recovered to a level which is representative of the rest of the Danian section. The diversity index remains relatively stable above 360 mbsf, which may indicate relatively stable environmental conditions from late Danian to early Selandian time.

PALEOBIOGEOGRAPHIC PATTERNS OF SELECTED TAXA

Quantitative data on calcareous nannofossil assemblages of the Danian are still quite limited. Data for particular time-slices of the Danian have been presented by Haq and Lohmann (1976) and Haq et al. (1977, 1979). These data will be compared with those of the present study to reveal some paleobiogeographic patterns.

Except for two samples in Zone NA3, *Coccolithus pelagicus* has a relatively low abundance (a few percent) at Site 738. Haq et al. (1977, 1979) showed that *C. pelagicus* constitutes 35%–12% of the total assemblage from low to high (about 53°N) latitudes for Zone NP3 (= NA4). For this time interval, Site 738 yielded less than 10% *C. pelagicus*. All these data are consistent with a general trend of decreasing *C. pelagicus* abundance toward higher latitudes. This trend is the reverse of that in the modern ocean, where *C. pelagicus* is most abundant in cool waters of the northern high-latitudes (McIntyre and Bé, 1967). This may indicate that *C. pelagicus* has undergone an evolutionary shift in its ecological response, as suggested by Haq and Lohmann (1976).

Haq and Lohmann (1976) reported *Coccolithus subpertusus* (*Ericsonia subpertusa*) as high as 47% in the low latitudes and as low as a few percent in the high latitudes. At Site 738, *C. subpertusus* only occurs sporadically. It thus seems that *C. subpertusus* also has a general trend of decreasing abundance toward higher latitudes and can be considered as a warm-water species.

Prinsius martinii and *Prinsius bisulcus* are very abundant at Site 738. This is contrary to the conclusion of Haq (1976) that *Prinsius martinii* is a boreal species that was confined to the northern high latitudes based on his observation of rare *P. martinii* at Site 323 (Leg 35) in the high-latitude South Pacific Ocean. In addition to Site 738, *P. martinii* and *P. bisulcus* were found to be abundant in other southern high-latitude localities, such as the Falkland Plateau in the South Atlantic (Wise and Wind, 1977) and Maud Rise in the Weddell Sea (Pospichal and Wise, 1990). Therefore, the conclusion that the northern high latitudes were cooler than the southern high latitudes during the Danian based on the rare *P. martinii* at Site 323 (Leg 35) (Haq, 1976) should be revised in light of the data available now.

Elsewhere, blooms of *Thoracosphaera*, *Braarudosphaera*, *Cyclagelosphaera*, *Biscutum? romeinii*, *Biscutum? parvulum*, *Toweius petalosus*, and *Octolithus multiplus* have been reported in Zones NP1 and NP2 (Perch-Nielsen, 1981a; Romein and Smit, 1981; Thierstein, 1981). All of these blooms have been found only in low or middle latitudes and some may be of only local significance. However, at the southern high-latitude Site 738, *Thoracosphaera* is significantly less abundant than in the lower latitudes, and no bloom of this taxon was recognized. No specimen of *Braarudosphaera*, *Biscutum? parvulum*, *Biscutum? romeinii*, *Cyclagelosphaera*, or *Octolithus multiplus* was found at Site 738. These taxa are also absent or very scarce at Sites 689 and 690 in the Weddell Sea (Pospichal and Wise, 1990). Since

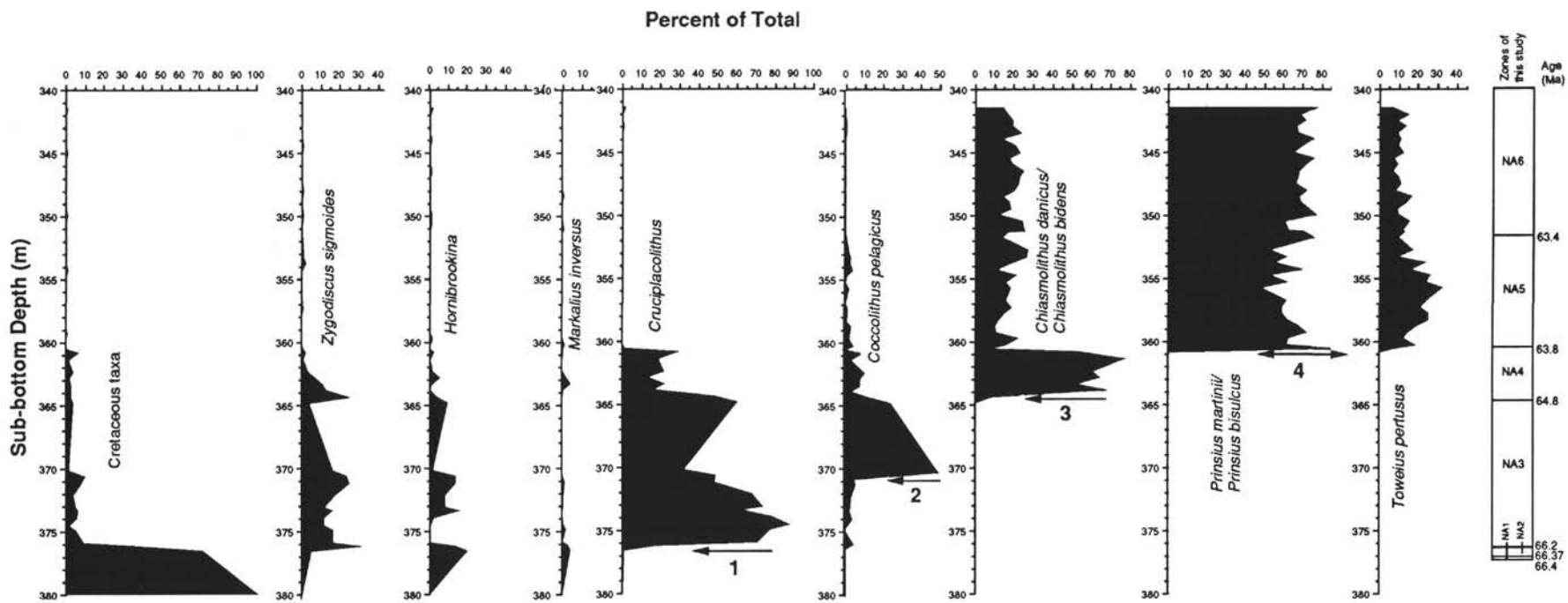


Figure 5. Changes of relative abundance of individual taxa through time at Site 738. Arrows 1 through 4 indicate four steps of large floral composition change. See zonation section for numerical age estimates.

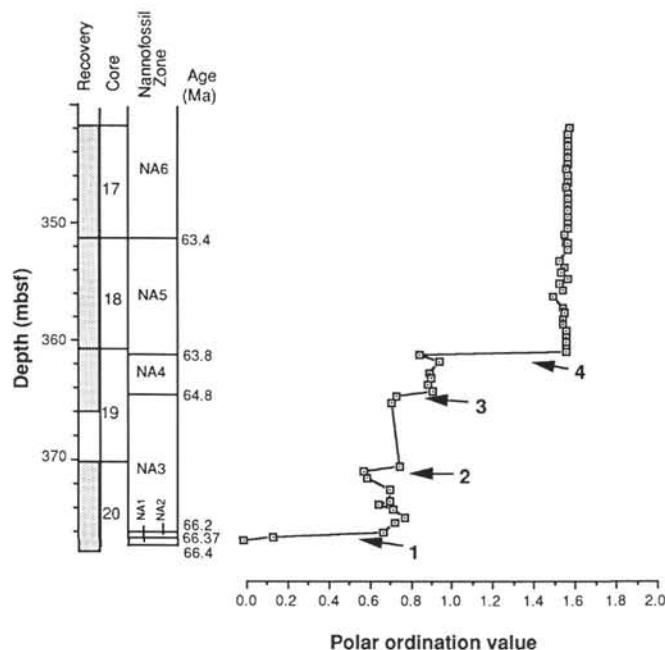


Figure 6. Polar ordination values of Site 738. The lower the value, the greater the similarity of the sample is in relation to Sample 119-738C-20R-5, 35–36 cm, and the greater the dissimilarity of the sample is in relation to Sample 119-738C-17R-1, 15–16 cm. Arrows 1 through 4 indicate four steps of large floral composition change. See zonation section for numerical age estimates.

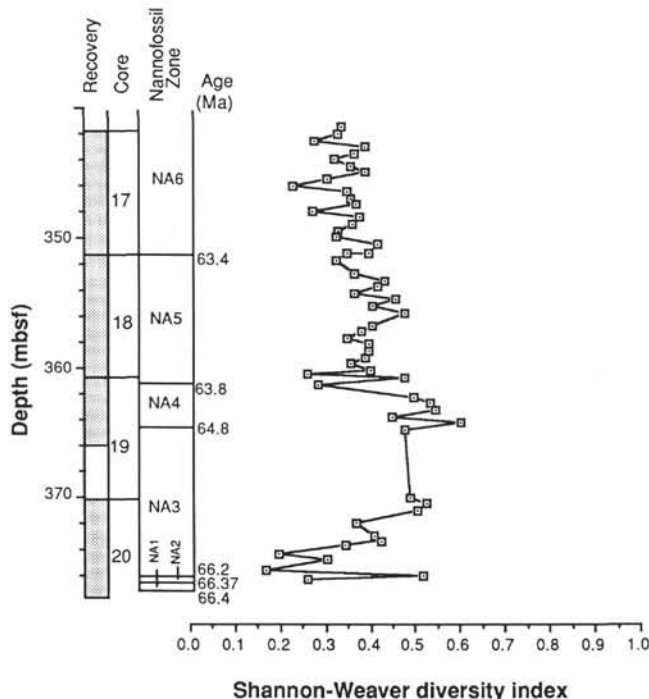


Figure 7. Changes of Shannon-Weaver diversity index through time at Site 738. See zonation section for numerical age estimates.

the paleodepth of Site 738 was relatively shallow (about 1000 m), the absence of these taxa was probably not due to preservational factors, but due to latitudinal differentiation of coccoliths in responding to the latitudinal thermal gradient in the Danian ocean. One distinct feature at Site 738 is the bloom of *Hornibrookina* shortly after the Cretaceous/Tertiary boundary extinction. This bloom appears to be a distinct southern high-latitude feature as it was also found in the Weddell Sea (Pospichal and Wise, 1990), but it has never been reported in middle or low latitudes.

SUMMARY AND CONCLUSIONS

The Danian section recovered at Site 738 is a biostratigraphically complete section divisible into six calcareous nannofossil zones. A southern high-latitude zonal scheme is proposed for the Danian, utilizing traditional markers and several nontraditional markers, i.e., FO's of *Hornibrookina*, *Prinsius martinii*, and *Chiasmolithus bidens*, and the LO of *Hornibrookina teuriensis*.

Quantitative analyses of calcareous nannofossil assemblages clearly indicate four large nannofloral changes, or steps, during the 1.6 m.y. interval after the Cretaceous/Tertiary boundary event. This rapid change of population is similar to that shown by the planktonic data from DSDP Site 577 (Leg 86) (Gerstel et al., 1987). The large and rapid changes of the planktonic biota may indicate unstable surface-water environments in the early Danian.

Taxa that bloom in low- or middle-latitude Danian sections, such as *Braarudosphaera*, *Biscutum? parvulum*, *Biscutum? romeinii*, *Cyclagelosphaera*, *Octolithus multiplus*, and *Toweius petalosus*, are absent at Site 738, as they are in the Weddell Sea. On the other hand, *Hornibrookina* bloom shortly after the Cretaceous/Tertiary boundary at the Antarctic site. Cool-water taxa, such as *Prinsius dimorphosus*, *Chiasmolithus danicus*, *Prinsius martinii*, *Chiasmolithus bidens*, and *Prinsius bisulcus* are most abundant in the Antarctic region. This indicates a distinct latitudinal differentiation of nannofloral assemblages, which may suggest significant latitudinal thermal gradients in the Danian oceans.

TAXONOMIC NOTES OF SELECTED SPECIES

Coccolithus subpertusus (Hay and Mohler) Wei and Pospichal, n. comb.

Basionym. *Ericsonia subpertusa* Hay and Mohler, 1967, p. 1517, pl. 198, figs. 11, 15, 18; pl. 199, figs. 1–3; Roth, 1973, p. 730, pl. 12, fig. 5; Haq and Lohmann, 1976, pl. 1, figs. 7, 8; Perch-Nielsen, 1977, pl. 16, fig. 3; Romein, 1979, p. 212, pl. 2, fig. 6.

Remarks. Bukry (1973, p. 676) was the first to point out that the genus *Ericsonia* is potentially a junior synonym of *Coccolithus*. Subsequently, Romein (1979, p. 105) and Wise (1983, p. 505) also presented reasons for this synonymy. Proximal views of *Coccolithus pelagicus* show the same three concentric cycles, as the type of *Ericsonia*, *E. occidentalis*, although the relative proportions of the cycles have been noted to differ. However, Wise (1983, p. 505, pl. 3) clearly showed that *C. pelagicus* changes the width of the outer proximal cycle with respect to that of the inner proximal cycle with time. It would be artificial to break the lineage between genera at some arbitrary point in the Paleogene or the Neogene based on the widths of the two proximal cycles, as proposed by Romein (1979). We agree with Wise (1983) that such a break should be made only to denote species or subspecies but not genera.

Hornibrookina teuriensis Edwards, 1973

Hornibrookina teuriensis Edwards, 1973, p. 77, figs. 22–37; Perch-Nielsen, 1977, p. 816, pl. 46, figs. 1, 4, 7–12.

Hornibrookina edwardsii Perch-Nielsen, 1977, p. 816, pl. 46, figs. 2, 3, 5, 6; p. 820, pl. 49, figs. 40–41.

Remarks. Perch-Nielsen (1977) proposed *Hornibrookina edwardsii*, which was believed to differ from *H. teuriensis* by a smaller central area and a row of protruding knobs around the central area. It is noted that

the central area of *H. edwardsii* varies considerably, and seems not to be a useful criterion for separating *H. edwardsii* from *H. teuriensis*. The protruding knobs may be the result of diagenetic effects or intraspecific variation. Furthermore, the protruding knobs usually cannot be observed in the light microscope, rendering this criterion useless. In Leg 119 material, no differentiation of *H. teuriensis* and *H. edwardsii* was possible, and *H. teuriensis* is recorded in the range chart. *Hornibrookina teuriensis* (*H. edwardsii*) bloomed shortly after the Cretaceous/Tertiary boundary at Maud Rise off East Antarctica (Pospichal and Wise, 1990) and at the Kerguelen Plateau, both extreme southern high-latitude areas. In the Brazos River section in Texas the species is present about 1 m above the Cretaceous/Tertiary boundary and is consistently rare (Jiang and Gartner, 1986).

***Prinsius bisulcus* (Stradner) Hay and Mohler, 1967**

Coccolithus bisulcus Stradner, in Gohrbant, 1963, p. 72, pl. 8, figs. 3–6. *Prinsius bisulcus* (Stradner) Hay and Mohler, 1967, p. 1529, pl. 196, figs. 10–13, pl. 197, fig. 6; Roth, 1973, p. 732, pl. 17, fig. 2; Wise and Wind, 1976, pl. 6, figs. 3–6; Okada and Thierstein, 1979, p. 540, pl. 3, fig. 5; Monechi 1985, p. 339, pl. 11, fig. 1. *Prinsius martinii* (Perch-Nielsen) Haq, Perch-Nielsen, 1977, p. 794, pl. 30, fig. 3.

Remarks. This species is elliptical, with a lens-shaped centro-distal cycle (see Romein, 1979, p. 119, for terminology). The distal shield slopes at a relatively small angle, whereas in *Toweius*, the angle is much larger. *Prinsius bisulcus* has the same structure as *P. martini* (Perch-Nielsen) Haq but is larger (see Pl. 3, Fig. 1 of this paper for a comparison of the two species). Though no clear size differentiation has been established between the two species, specimens larger than 5.5 μm are generally placed in *P. bisulcus*. This first species occurs later than *P. martinii*. The recognition of the two species is thus useful for biostratigraphic purposes.

***Prinsius dimorphosus* (Perch-Nielsen) Perch-Nielsen, 1977**

Coccosphere, in Bramlette and Martini, 1964, p. 320, pl. 1, figs. 19, 20. *Coccolithus* (?) sp. in Martini, 1964, p. 52, pl. 2, figs. 5, 6. Fossil coccosphere, in Black, 1965, p. 32, fig. 3. *Biscutum dimorphosus* Perch-Nielsen, 1972, pl. 8, fig. 2. *Prinsius dimorphosus* (Perch-Nielsen) Perch-Nielsen, 1977, p. 794, pl. 30, fig. 10–13; Romein, 1979, p. 214, pl. 3, figs. 3–6. *Biscutum? tenuiculum* Okada and Thierstein, 1979, p. 536, pl. 1, figs. 1, 2; p. 552, pl. 9, figs. 1–8. *Prinsius rosenkrantzii* Perch-Nielsen, 1979, p. 132, pl. 4, figs. 1–5. *Prinsius africanus* Perch-Nielsen, 1981b, p. 854, pl. 3, 5, 6, 8, 9. *Prinsius tenuiculum* (Okada and Thierstein) Perch-Nielsen, Wei and Wise, 1989, p. 124, pl. 1, figs. 1, 2; p. 125, pl. 2, figs. 7, 8.

Remarks. This is a small-sized coccolith species (2–4 μm) with a round or elliptical outline. Romein (1979) described the round specimens as type 1 and the elliptical as type 2 of *Prinsius dimorphosus*. *Biscutum? tenuiculum* Okada and Thierstein belongs to the round type of *P. dimorphosus*. The round and the elliptical types are gradational and usually occur together, although it appears that the elliptical types are more abundant in high-latitude sediments than in middle- or low-latitude sediments. The holotype of *Prinsius rosenkrantzii* Perch-Nielsen is round and is a junior synonym of *P. tenuiculum* (see Perch-Nielsen, 1984); the latter is in turn a junior synonym of *P. dimorphosus*. According to Perch-Nielsen (1981), *Prinsius africanus* Perch-Nielsen differs from *P. tenuiculum* by having a central netlike structure and a crown. However, the central netlike structure can only be seen in well-preserved material (isotypes of *P. africanus* do not show the netlike structure; see Perch-Nielsen, 1981, p. 854, pl. 3, figs. 6, 9) and a crown does exist in many typical *P. dimorphosus* specimens (see Pl. 1, Figs. 1 and 4). Therefore, *P. africanus* must be considered as a junior synonym of *P. dimorphosus*.

***Prinsius martinii* (Perch-Nielsen) Haq, 1977**

Ericsonia? martinii Perch-Nielsen, 1969, p. 324, pl. 32, figs. 3b, 5–7. *Prinsius martinii* (Perch-Nielsen) Haq, 1971, p. 18, pl. 5, figs. 2, 3, 5, 10, non pl. 5, fig. 1; Romein, 1979, p. 214, pl. 3, figs. 7, 8. non *Prinsius martinii* (Perch-Nielsen) Haq, Okada and Thierstein, 1979, p. 558, pl. 12, figs. 12, 14.

Remarks. As pointed out above, *Prinsius martinii* has the same structure as *P. bisulcus*, but is smaller (4–5 μm). Edwards (1971) first sug-

gested the use of the first occurrence of *P. martinii* as a datum in the Danian. This datum has proved to be very useful in subsequent studies (e.g., Perch-Nielsen, 1979; Romein, 1979; Pospichal and Wise, 1990; this study).

***Toweius pertusus* (Sullivan) Romein, 1979**

Coccolithus pertusus Sullivan, 1965, p. 32, pl. 3, figs. 5, 6. *Toweius craticulus* Hay and Mohler, 1967, p. 1530, pl. 196, figs. 7–9; pl. 197, figs. 2, 3; Edwards and Perch-Nielsen, 1975, p. 502, pl. 3, figs. 3, 6, 8, 9; Perch-Nielsen, 1977, p. 794, pl. 30, fig. 4; Okada and Thierstein, 1979, p. 560, pl. 13, figs. 1, 2; Wise and Wind, 1977, p. 324, pl. 6, fig. 2. *?Toweius eminens* (Bramlette and Sullivan) Perch-Nielsen, Perch-Nielsen, 1977, p. 794, pl. 30, fig. 6. *Toweius tovae* Perch-Nielsen, 1971, p. 359, pl. 13, figs. 1–3, 5, pl. 14, figs. 3, 9; Wise and Wind, 1977, p. 322, pl. 5, fig. 6; p. 324, pl. 6, fig. 1; Okada and Thierstein, 1979, p. 560, pl. 13, figs. 5, 6; Monechi, 1985, p. 333, pl. 11, fig. 4. *Toweius* sp. Wise and Wind, 1977, p. 326, pl. 7, fig. 1.

Remarks. As pointed out correctly by Romein (1979), *Toweius craticulus* Hay and Mohler is a junior synonym of *Toweius pertusus* (Sullivan) Romein. This species has six or more central pores. The number of central pores varies considerably, and is controlled in part by the degree of diagenetic overgrowth. Perch-Nielsen (1971) separated forms with six or seven pores from *Toweius eminens* (with four pores), naming them *Toweius tovae*. Romein (1979) placed *T. tovae* into synonymy with *T. eminens*. In our opinion, *T. eminens* is a distinct species with four central pores. Specimens of *Toweius* with five central pores are rare. On the other hand, there seems to be a continuous spectrum of variation in the central pores from six up. It is therefore appropriate to synonymize *T. tovae* with *T. pertusus*. Some specimens of *T. pertusus* are very elliptical and they are transitional to *P. bisulcus* (Pl. 3, Fig. 6, see also Edwards and Perch-Nielsen, 1975, p. 504, pl. 4, fig. 6).

ACKNOWLEDGMENTS

We are grateful to Professor Hans Thierstein, who kindly shared his information with us on the Cretaceous/Tertiary boundary studies of Site 738 and made helpful suggestions concerning this paper. He examined the core-catcher samples and samples from Section 119-738C-20R-5 aboard ship, and some of these data have been incorporated in this paper. We would also like to thank Professors Jack Baldauf, Timothy Bralower, and Sherwood Wise, and an anonymous reviewer for their most helpful reviews. This research was supported by grants from the U.S. Science Advisory Committee and a Florida State University Dissertation Fellowship to the first author. Equipment used was provided by NSF grant DPP89-17976 and a grant from the Amoco Foundation to S. W. Wise.

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Date of initial receipt: 4 October

Date of acceptance: 9 April 1990

Ms 119B-154

APPENDIX
Percentage of Individual taxa

Core, section, interval (cm)	Depth (mbsf)	C. bi/da	C. pel	Cruci	Horni	M. inv	P. mar	Towe	Z. sig	Creta
17-1, 15	341.45	14.4	0.0	0.9	0.9	0.0	76.6	6.8	0.0	0.5
17-1, 66	341.96	16.5	0.0	0.0	0.0	0.0	69.0	14.5	0.0	0.0
17-1, 115	342.45	19.1	0.5	0.0	0.2	0.0	71.2	9.0	0.0	0.0
17-2, 15	342.95	19.1	0.4	0.4	0.0	0.0	66.8	13.4	0.0	0.0
17-2, 66	343.46	23.3	0.4	0.0	0.0	0.0	67.5	8.8	0.0	0.0
17-2, 115	343.96	13.6	0.0	0.0	0.3	0.0	75.3	10.8	0.0	0.0
17-3, 15	344.45	20.4	0.0	0.0	0.0	0.0	69.8	9.5	0.4	0.0
17-3, 66	344.96	22.5	0.0	0.0	0.0	0.0	65.0	12.1	0.0	0.4
17-3, 115	345.45	17.3	0.0	0.0	0.0	0.0	75.6	7.1	0.0	0.0
17-4, 15	345.95	19.0	0.0	0.0	0.0	0.0	72.1	8.9	0.0	0.0
17-4, 66	346.46	24.7	0.0	0.0	0.5	0.0	68.4	6.3	0.0	0.0
17-4, 115	346.95	22.9	0.0	0.0	0.0	0.0	67.6	9.6	0.0	0.0
17-5, 15	347.45	22.4	0.0	0.0	0.0	0.0	66.7	10.9	0.0	0.0
17-5, 66	347.96	19.6	0.0	0.0	0.0	0.0	71.8	8.0	0.6	0.0
17-5, 115	348.45	14.6	0.0	0.0	0.0	0.6	68.4	16.4	0.0	0.0
17-6, 15	348.95	17.6	0.0	0.0	0.0	0.0	68.6	13.7	0.0	0.0
17-6, 66	349.46	17.9	0.0	0.0	0.0	0.0	72.8	9.3	0.0	0.0
17-6, 115	349.95	12.2	0.0	0.0	0.4	0.0	76.7	9.8	0.4	0.4
17-7, 15	350.45	24.2	0.0	0.0	0.5	0.0	60.3	15.1	0.0	0.0
17-CC	351.18	25.3	0.0	0.0	0.0	0.5	62.9	11.3	0.0	0.0
18-1, 35	351.25	16.8	0.0	0.0	0.0	0.0	70.6	12.6	0.0	0.0
18-1, 85	351.75	13.8	0.8	0.0	0.0	0.0	75.6	9.8	0.0	0.0
18-2, 35	352.75	27.2	2.0	0.0	0.0	0.3	52.4	17.0	0.7	0.3
18-2, 85	353.25	26.5	2.7	0.0	0.0	0.0	61.1	9.2	0.5	0.0
18-2, 135	353.75	18.3	2.3	0.0	0.0	0.0	55.0	22.9	1.5	0.0
18-3, 35	354.25	10.9	3.9	0.0	0.0	0.0	69.0	15.5	0.0	0.8
18-3, 85	354.75	20.5	0.9	0.0	0.0	0.0	52.7	25.9	0.0	0.0
18-3, 135	355.25	15.5	0.0	0.0	0.0	0.0	60.8	23.7	0.0	0.0
18-4, 35	355.75	18.2	1.8	0.0	0.0	0.0	48.2	31.8	0.0	0.0
18-4, 135	356.75	15.5	0.0	0.0	0.0	0.0	60.8	23.7	0.0	0.0
18-5, 35	357.25	18.6	1.1	0.0	0.0	0.0	59.0	20.8	0.5	0.0
18-5, 85	357.75	14.8	1.0	0.0	0.0	0.0	59.0	25.2	0.0	0.0
18-5, 135	358.25	12.1	0.4	0.0	0.0	0.0	62.4	25.1	0.0	0.0
18-6, 30	358.70	10.4	2.8	0.0	0.0	0.0	67.9	18.9	0.0	0.0
18-6, 85	359.25	10.6	2.7	0.0	0.0	0.0	71.7	14.2	0.0	0.9
18-6, 135	359.75	21.7	2.4	0.0	1.2	0.0	62.7	12.0	0.0	0.0
18-7, 35	360.25	15.8	4.0	0.0	0.0	1.0	61.4	17.8	0.0	0.0
18-CC	360.50	5.6	0.2	0.9	0.0	0.4	85.3	6.6	0.5	0.5
19-1, 35	360.85	53.7	7.8	28.3	2.0	0.4	0.0	0.0	2.0	5.9
19-1, 85	361.35	76.5	3.1	18.9	0.0	0.0	0.0	0.0	5.0	1.0
19-2, 35	362.35	60.5	10.1	21.6	1.4	0.0	0.0	0.0	2.9	3.5
19-2, 85	362.85	63.8	7.6	13.3	4.8	1.9	0.0	0.0	6.7	1.9
19-2, 135	363.35	52.4	8.1	21.5	0.6	4.0	0.0	0.0	10.7	2.7
19-3, 35	363.85	66.3	2.5	15.3	0.0	0.6	0.0	0.0	12.7	2.5
19-3, 85	364.35	8.4	12.9	48.4	3.9	0.0	0.0	0.0	23.9	2.6
19-3, 135	364.85	0.0	24.4	59.7	9.2	0.0	0.0	0.0	3.4	3.4
19-CC	370.20	0.0	49.3	31.8	1.2	0.0	0.0	0.0	16.2	1.5
20-1, 46	370.66	0.0	5.0	48.7	13.4	0.0	0.0	0.0	23.5	9.4
20-1, 98	371.18	0.0	5.2	48.1	13.0	1.3	0.0	0.0	24.7	7.8
20-2, 42	372.12	0.0	2.9	67.7	8.0	0.3	0.0	0.0	17.5	3.6
20-2, 138	373.08	0.0	2.6	73.1	7.9	0.7	0.0	0.0	11.2	4.6
20-3, 17	373.37	0.0	2.4	61.9	14.3	0.0	0.0	0.0	15.5	6.0
20-3, 70	373.90	0.0	3.5	77.8	2.1	0.0	0.0	0.0	11.1	5.6
20-3, 137	374.57	0.0	0.0	86.9	0.0	0.6	0.0	0.0	11.3	1.3
20-4, 20	374.90	0.0	0.0	77.3	0.0	1.6	0.0	0.0	16.4	4.7
20-4, 117	375.87	0.0	4.0	70.7	0.2	0.0	0.0	0.0	16.0	9.0
20-5, 2	376.22	0.0	0.0	17.4	13.0	3.1	0.0	0.0	28.6	37.9
20-5, 35	376.55	0.0	0.0	0.4	19.0	4.1	0.0	0.0	4.5	71.9
20-CC	379.90	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	100.0

Note: Depth is expressed as meters below seafloor (mbsf). C. di/da = *Chiasmolithus bidens/C. danicus*; C. pel = *Coccolithus pelagicus*; Cruci = *Crucialcolithus*; Horni = *Hornibrookina*; M. inv = *Markalius inversus*; P. mar = *Prinsius martinii/P. bisulcus*; Towe = *Toweius pertusus*; Z. sig = *Zygodiscus sigmoides*; Creta = Cretaceous taxa.

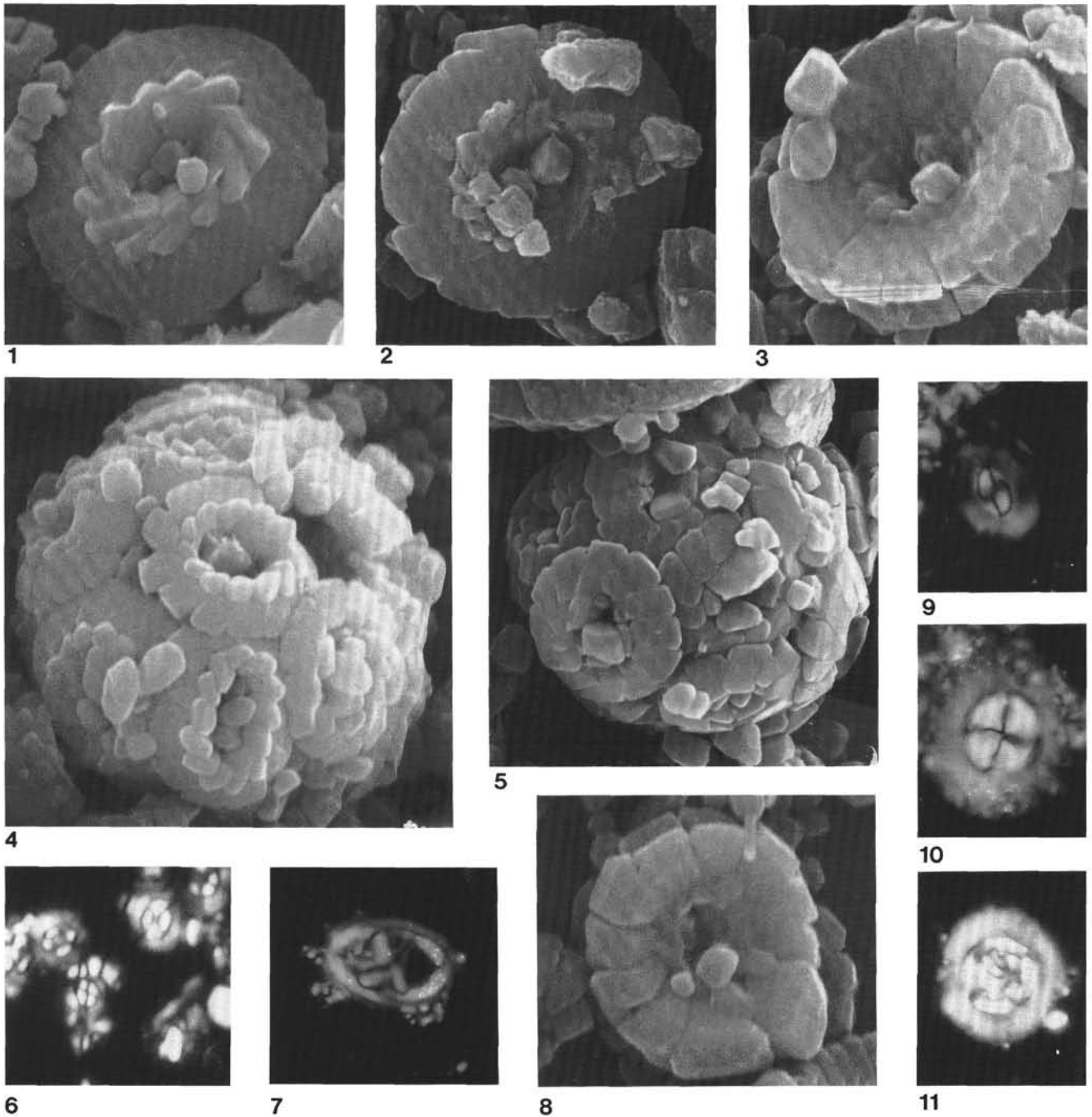


Plate 1. 1-6, 8. *Prinsius dimorphosus* (Perch-Nielsen) Perch-Nielsen. (1) Distal view, elliptical type, Sample 119-738C-19R-CC, 15,000 \times ; (2) distal view, elliptical type, Sample 119-738C-19R-CC, 13,000 \times ; (3) proximal view, round type, Sample 119-738C-20R-1, 46-47 cm, 15,000 \times ; (4) coccosphere, elliptical type, Sample 119-738C-18R-7, 55-56 cm, 10,500 \times ; (5) coccosphere, round type, Sample 119-738C-20R-4, 20-21 cm, 11,000 \times ; (6) both elliptical and round types, Sample 119-738C-19R-3, 85-86 cm, 3,000 \times ; (8) round type, Sample 119-738C-18R-7, 35-36 cm, 15,000 \times . 7. *Neochiastozygus concinnus* (Perch-Nielsen) Perch-Nielsen, Sample 119-738C-20R-5, 135-136 cm, 2,500 \times . 9. *Biscutum constans* (Gorka) Black, Sample 119-738C-20R-5, 35-36 cm, 2,500 \times . 10. *Markalius inversus* (Deflandre) Bramlette and Martini, Sample 119-738C-18R-1, 135-136 cm, 2,500 \times . 11. *Markalius apertus* Perch-Nielsen, Sample 119-738C-17R, 66-67 cm, 2,500 \times .

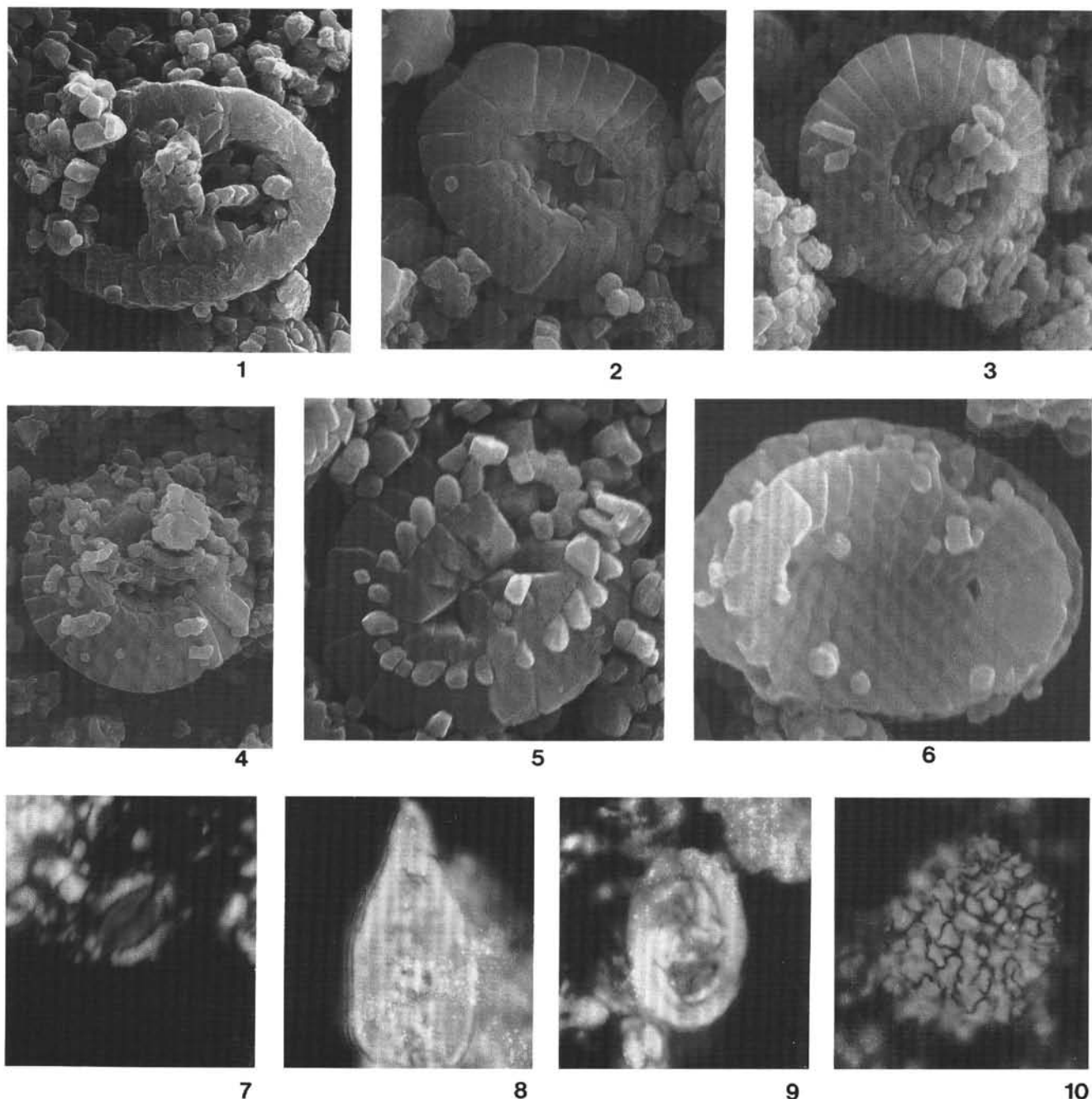


Plate 2. 1, 9. *Zygodiscus sigmoides* Bramlette and Sullivan. (1) Distal view, Sample 119-738C-20R-4, 20–21 cm, 6,000 \times ; (9) Sample 119-738C-19R-CC, 6,000 \times . 2. *Biscutum constans* (Gorka) Black, distal view, Sample 119-738C-18R-7, 35–36 cm, 7,000 \times . 3, 4. *Markalius inversus* (Deflandre) Bramlette and Martini (3) Distal view, Sample 119-738C-18R-7, 35–36 cm, 4,500 \times ; (4) Sample 119-738C-19R-1, 35–36 cm, 4,500 \times . 5–7. *Hornibrookina teuriensis* Edwards. (5) Distal view, Sample 119-20R-1, 46–47 cm, 8,500 \times ; (6) proximal view, Sample 119-738C-19R-1, 35–36 cm, 9,500 \times ; (7) Sample 119-738C-20R-5, 35–36 cm, 2,500 \times . 8. *Lapideacassis* sp., Sample 119-738C-18R-5, 85–86 cm, 2,500 \times . 10. *Thoracosphaera* sp., Sample 119-738C-19R-2, 135–136 cm, 2,500 \times .

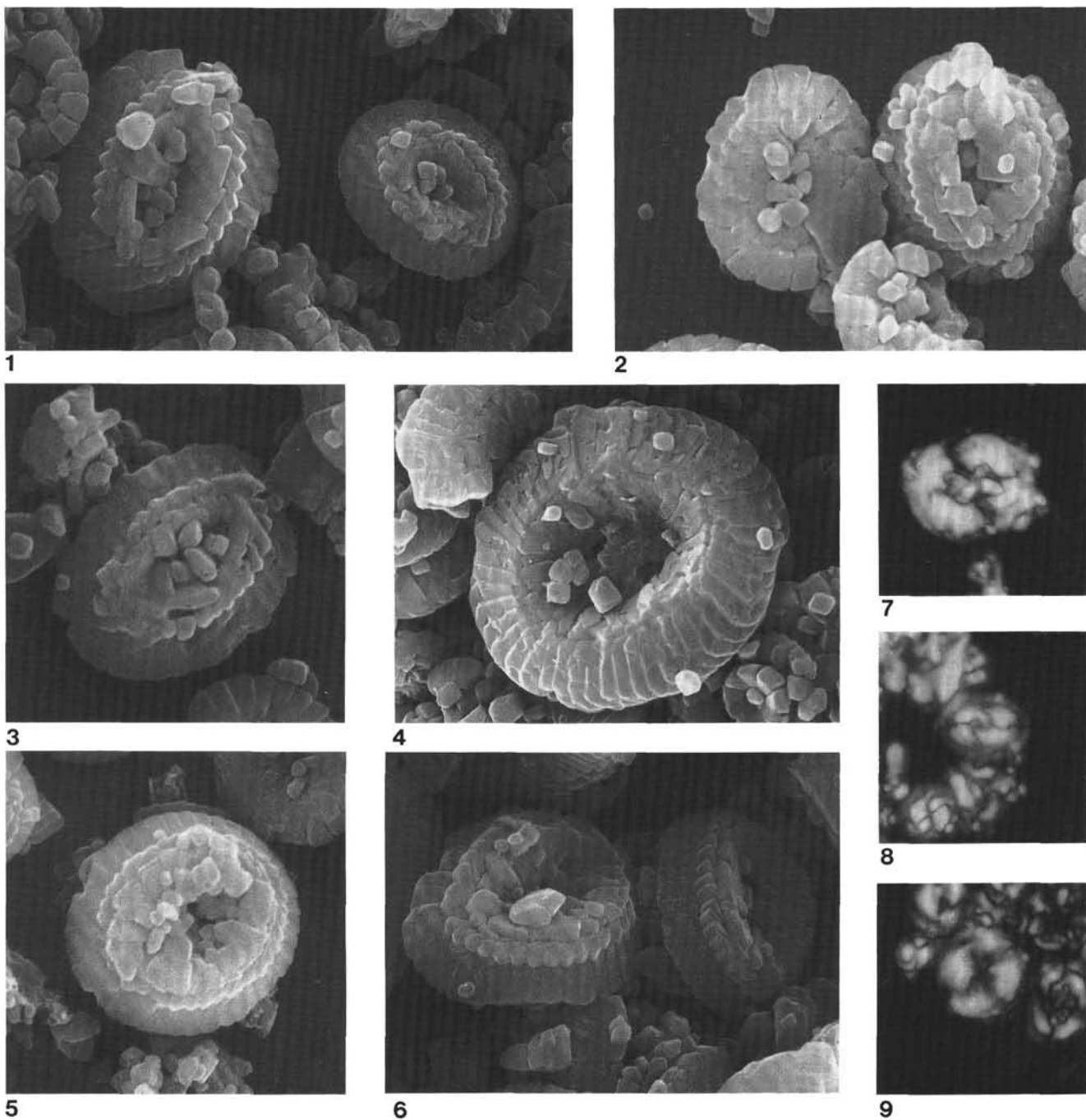


Plate 3. 1. *Prinsius bisulcus* (Stradner) Hay and Mohler on left, *Prinsius martinii* (Perch-Nielsen) Haq on right, both distal view, Sample 119-738C-17R-5, 15–16 cm, 8,500 \times . 2, 3, 8. *Prinsius bisulcus* (Stradner) Hay and Mohler. (2) Proximal view on left, distal view on right, Sample 119-738C-18R-CC, 7,500 \times ; (3) distal view, Sample 119-738C-17R-5, 15–16 cm, 9,500 \times ; (8) Sample 119-738C-18R-2, 135–136 cm, 2,500 \times . 4, 7. *Cruciplacolithus edwardsii* Romein. (4) Distal view, Sample 119-738C-19R-3, 135–136 cm, 6,500 \times ; (7) Sample 119-738C-19R-2, 85–86 cm, 2,500 \times . 5, 6, 9. *Toweius pertusus* (Sullivan) Romein. (5) Distal view, Sample 119-738C-17R-5, 15–16 cm, 6,500 \times ; (6) distal view, Sample 119-738C-17R-5, 15–16 cm, 8,000 \times , both specimens are elliptical; (9) Sample 119-738C-182, 135–136 cm, 2,500 \times .

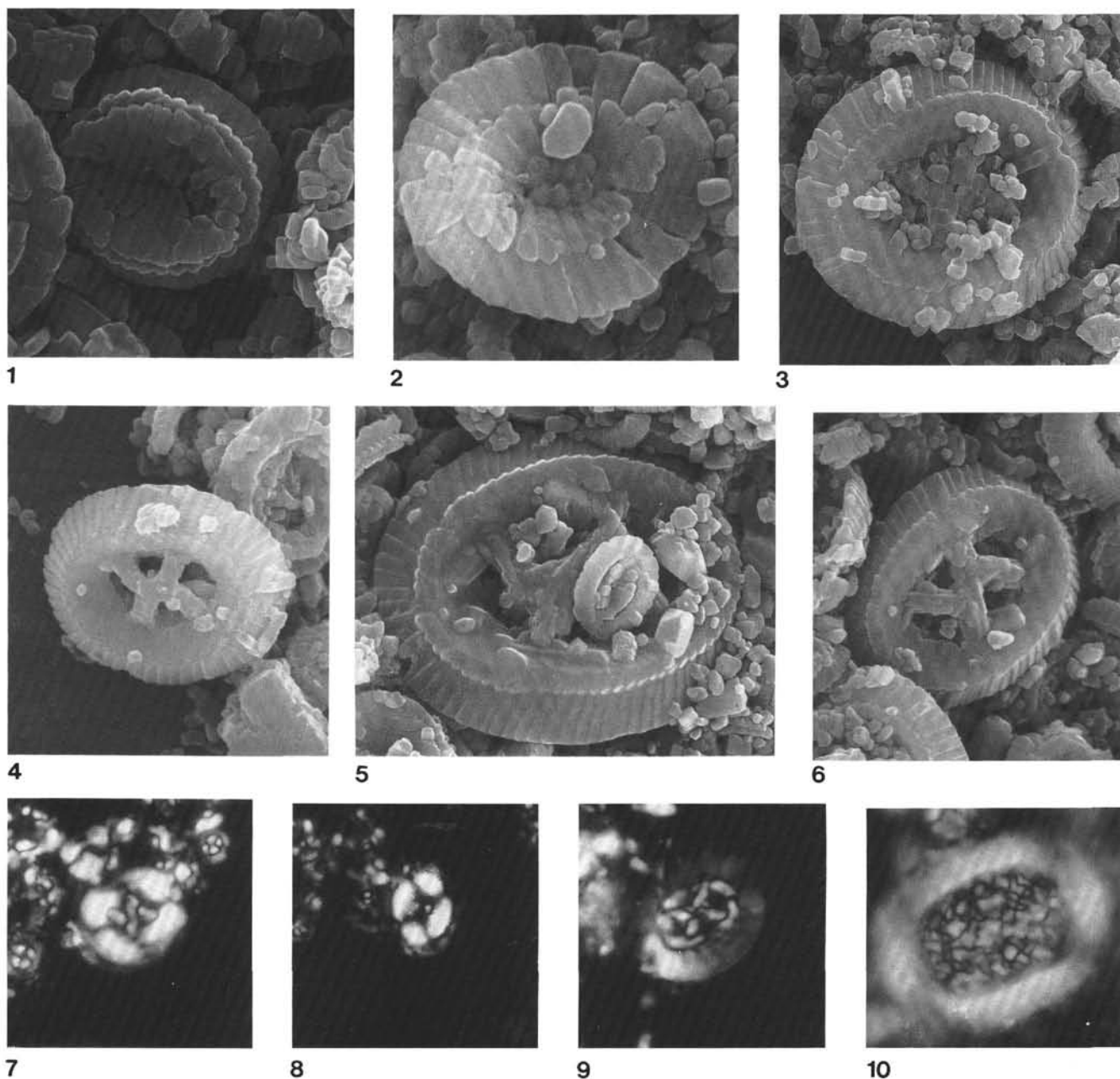


Plate 4. 1. *Toweius pertusus* (Sullivan) Romein, distal view, Sample 119-738C-17R-CC, 9,500 \times . 2. *Markalius inversus* (Deflandre) Bramlette and Martini, proximal view, Sample 119-738C-18R-7, 35–36 cm, 9,500 \times . 3, 4, 7. *Chiasmolithus danicus* (Brotzen) Hay and Mohler. (3) distal view, Sample 119-738C-18R-CC, 4,000 \times ; (4) distal view, Sample 119-738C-19R-1, 35–36 cm, 5,000 \times ; (7) Sample 119-738C-19R-1, 85–86 cm, 2,500 \times . 5, 6. *Chiasmolithus bidens* (Bramlette and Sullivan) Hay and Mohler, distal view, Sample 119-738C-17R-CC, 4,500 \times . 8. *Coccolithus pelagicus* (Willich) Schiller, Sample 119-738C-19R-2, 135–136 cm, 2,500 \times . 9. *Prediscosphaera cretacea* (Arkhangelsky) Gartner, Sample 119-738C-20R-4, 117–118 cm, 2,500 \times . 10. *Cribrosphaerella daniae* Perch-Nielsen, Sample 119-738C-20R-5, 35 cm, 2,500 \times .